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SULPHUR CREEK PLACER PROPERTY

Assessment Report on

Geophysical Exploration Program 2017

PLACER CLAIMS

JON 1-6

P515972-P 515974; P 517619-P517621

by

William LeBarge

Geoplacer Exploration Ltd.

For
Midnight Permitting Ltd.

Location of property: 63°40'22.4"N; 138°41'03"W
NTS map sheet: 115O/10
Mining District: Dawson
Date: November 9, 2017
Date of Work: May 24, 2017.

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Executive Summary

The following is an assessment report on the 2017 exploration program on the Sulphur Creek placer property, Jon 1-6 claims (P515972-P 515974; P 517619-P517621) by Geoplacer Exploration Ltd., on behalf of Midnight Permitting Ltd.

The property is located on an unnamed left limit tributary of Sulphur Creek, just upstream of its confluence with the Indian River. Access to the property can be gained by summer road from Dawson City via Hunker Creek and Sulphur Creek, a total distance from Dawson City of approximately 74 kilometres.

In 2017, an exploration program was undertaken which consisted of 630 metres of ground penetrating radar surveys, in 5 survey lines. The ground penetrating radar surveys appeared to delineate not only the bedrock contact but also the boundary between the ice-rich permafrost silt (“black muck”) and underlying ice-poor sediments (possibly sand and gravel) below. Ice-rich permafrost was evident as a zone of strong reflection in each survey profile, approximately 2 metres below the surface and extending to a depth of 5 to 6 metres below the surface. Overall, interpreted depths to bedrock varied between 2 metres and 10 metres.

Comparison of the previously-conducted resistivity surveys and the present GPR surveys showed similarities in interpreted depths to bedrock, and a commonality of the orientation and location of possible paleochannels. However, the GPR surveys tended to show bedrock deeper than corresponding nearby resistivity surveys.

Interpretations showed a distinctive paleochannel which was intersected in three of the GPR profiles and one of the resistivity surveys. This represents a highly prospective drill target.

Ground penetrating radar appeared to provide good signal response and interpretable data in the project area, although it is not recommended as a stand-alone geophysical method to interpret lithologies and depths to bedrock. Confirmation of interpretations by drilling is recommended. This would verify lithologies, sediment thicknesses and depths to bedrock, which would enable recalculation of the relative permittivity and depths of penetration and result in a more accurate interpretation of initial GPR results.

In addition, testing of the alluvial materials for placer gold is recommended. Initially, drilling methods such as auger (6 inch or larger), R/C (Reverse circulation) or RAB (Rotary Air Blast) may be used. This should be followed up by excavator test pitting and bulk processing of prospective alluvial gravels.

Introduction

The following is an assessment report on the 2017 geophysical exploration program on the Sulphur Creek placer property, Jon 1-6 claims (P515972-P 515974; P 517619-P517621) by Geoplacer Exploration Ltd., on behalf of Midnight Permitting Ltd. Previous exploration work has included resistivity geophysical surveys which were conducted by Kryotek Arctic Innovation Inc. in 2014 and 2015.

Location and Access

Sulphur Creek is a right limit tributary of the Indian River, located in central Yukon approximately 60 km by air south of Dawson City, Yukon (Figure 1). The Sulphur Creek Placer Property is located on an unnamed right limit tributary of Sulphur Creek near its confluence with the Indian River.

The centre of the property is 63°40'22.4"N and 138°41'03"W, on NTS map sheet 115O/10, in the Dawson Mining District (Figure 2).

Access to the property can be gained by summer road from Dawson City. The usual route runs from Dawson City along the Klondike Highway, then along Hunker Creek to King Solomon Dome, and down Sulphur Creek near its confluence with Indian River (approximately 74 kilometres).

Personnel and Dates of Work

The 2017 exploration program was designed and supervised by William LeBarge of Geoplacer Exploration Ltd. and Bud Davis of Midnight Permitting Ltd.

The ground penetrating radar surveys were conducted by William LeBarge (Geoplacer Exploration Ltd.) on May 24, 2017. The final report was completed by William LeBarge of Geoplacer Exploration Ltd.

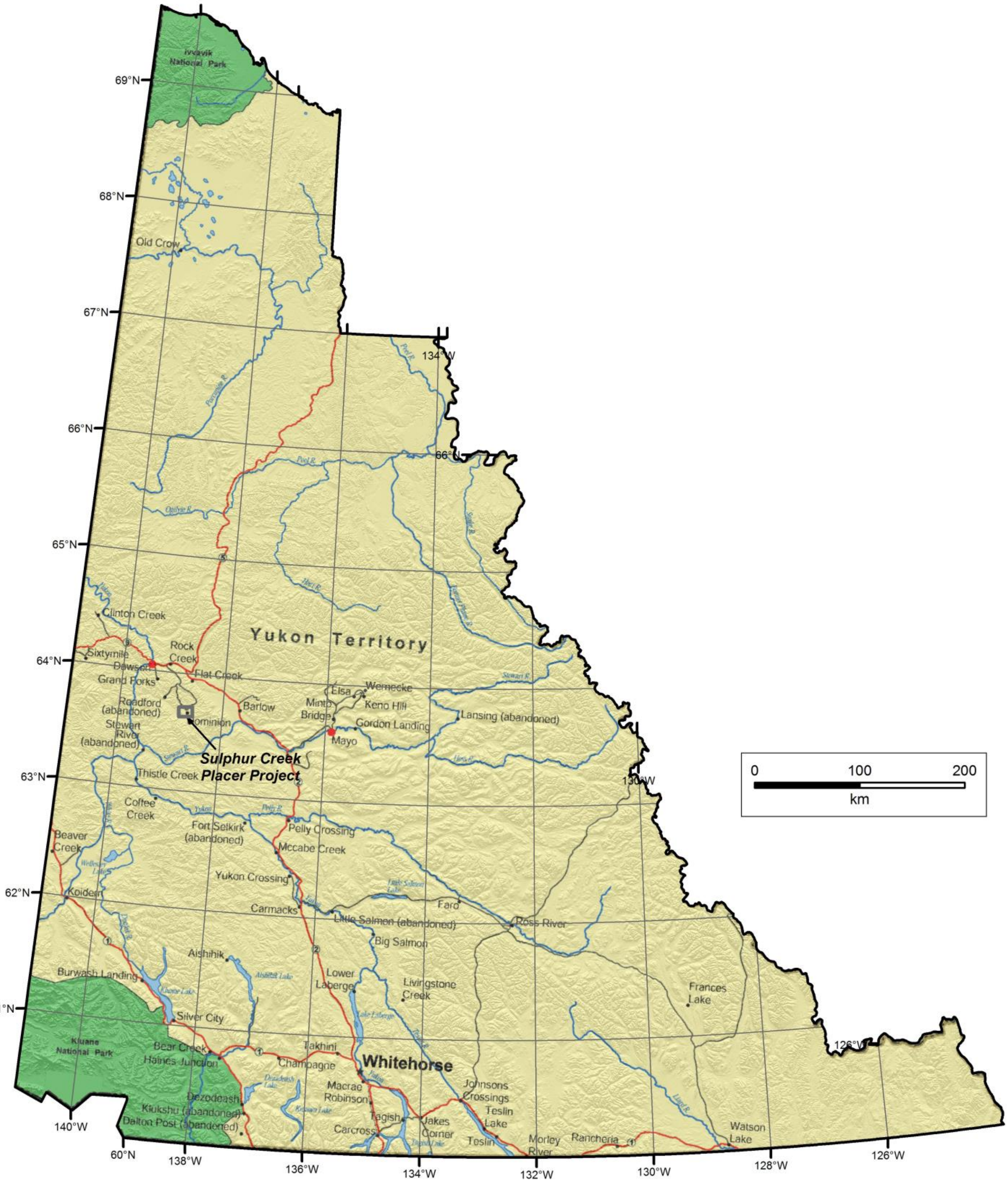

Figure 1 - General Location of Sulphur Creek Project, Yukon.

Placer Tenure

Table 1 shows a summary of the current claim status for the Sulphur Creek property.

Table 1 - Claim status, Jon Property, Sulphur Creek.

GRANT NUMBER	STATUS	CLAIM NAME	OWNER NAME	STAKING DATE	RECORDED DATE	EXPIRY DATE
P 515972	Active	Jon 1	Midnight Permitting Ltd. - 100%	5/27/2014	5/29/2014	11/29/2019
P 515973	Active	Jon 2	Midnight Permitting Ltd. - 100%	5/27/2014	5/29/2014	11/29/2019
P 515974	Active	Jon 3	Midnight Permitting Ltd. - 100%	5/28/2014	5/29/2014	11/29/2019
P 517619	Pending	Jon 4	Midnight Permitting Ltd. - 100%	5/21/2016	5/27/2016	11/27/2019
P 517620	Pending	Jon 5	Midnight Permitting Ltd. - 100%	5/21/2016	5/27/2016	11/27/2019
P 517621	Pending	Jon 6	Midnight Permitting Ltd. - 100%	5/21/2016	5/27/2016	11/27/2019

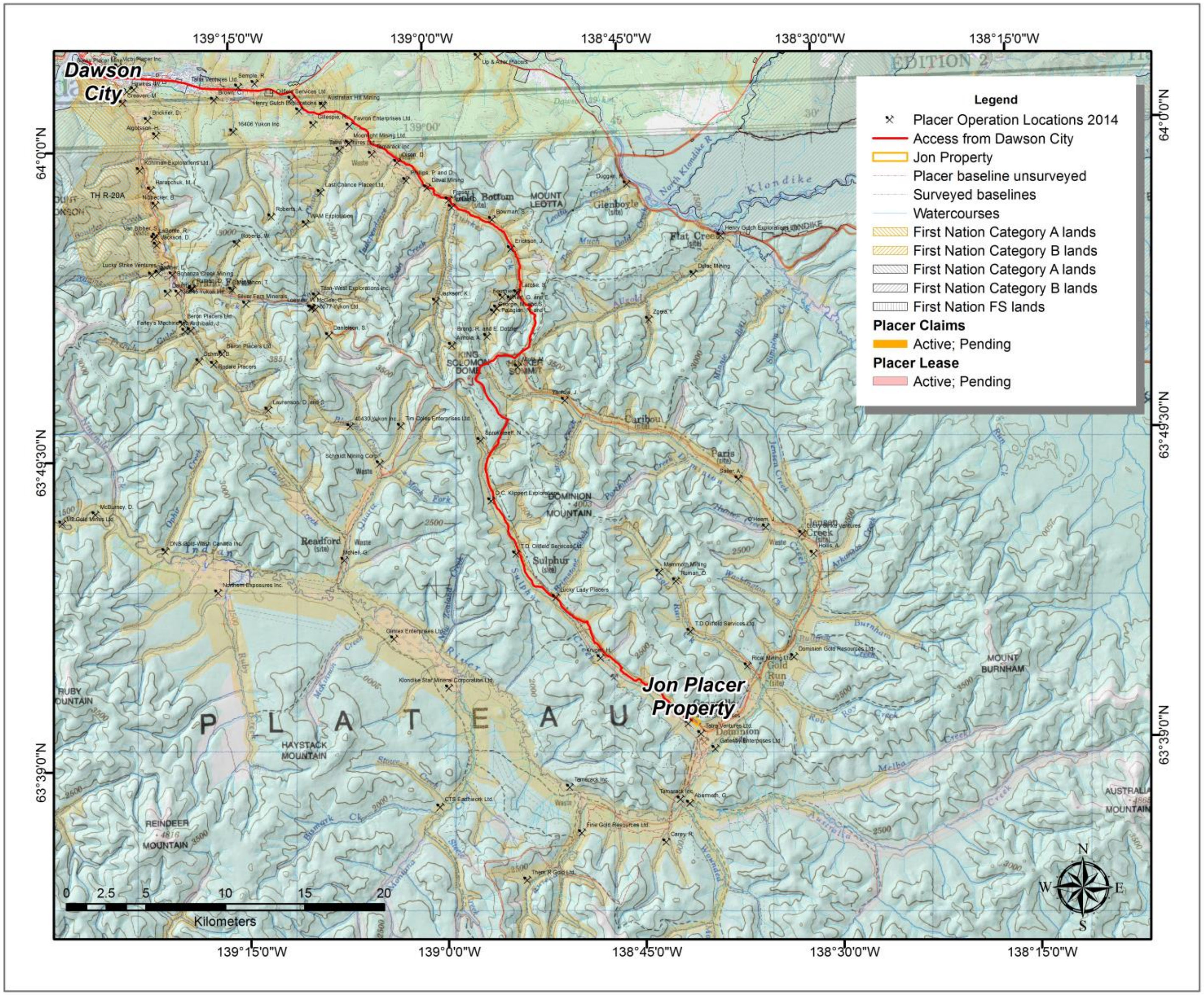

Figure 2 – Location of Sulphur Creek Placer Project and Dawson region placer tenures.

History of Exploration and Mining – Sulphur Creek

Sulphur Creek has been mined since the beginning of the Klondike Gold Rush in 1898, first by hand methods, and then by dredging. Green (1977) notes that three dredges mined on Sulphur Creek beginning in 1936. YCGC (Yukon Consolidated Gold Corporation) Dredge #6 mined 148,000 ounces between 1936 and 1966; YCGC Dredge #8 mined 212,000 ounces between 1937 and 1966 and YCGC Dredge #9 mined 113,000 ounces between 1938 and 1966.

Mechanical mining replaced the dredges after 1966 and dozens of operations have mined on Sulphur Creek from then up to the present day. Much of the activity is documented in LeBarge (2007) with more recent mining documented in LeBarge and Welsh (2007), LeBarge and Nordling (2011), and Van Loon and Bond (2014). Gold production from these sources and Yukon Government royalty records shows a total of over 352,000 ounces produced from Sulphur Creek between 1940 and 2015. This does not include the hand mining from the 40+ years previous.

The nearest active operation to the Jon Property is Favron Enterprises Ltd., who mined less than 1 km away on Sulphur Creek between 2010 and 2013. Just downstream from that operation was Tatra Ventures Ltd., who mined from 2010 to 2013.

Regional Bedrock Geology

The project area is situated within the Yukon-Tanana terrane, an accreted pericratonic sequence that covers a large part of the northern Cordillera from northern British Columbia to east-central Alaska (Gordey and Ryan, 2005; Colpron and Nelson, 2006). The Yukon Tanana Terrane consists of Paleozoic schist and gneiss that were deformed and metamorphosed in the late Paleozoic, and intruded by several suites of Mesozoic intrusions that range in age from Jurassic to Eocene (Colpron and Nelson, 2006). The Paleozoic rocks are pervasively foliated with at least two overprinting fabrics (MacKenzie and Craw, 2010; MacKenzie et al, 2008). During Late Permian to Early Jurassic time these rocks were tectonically-stacked along thrust faults which were parallel to regional foliation. Later tensional-extensional tectonics occurred during the mid-Cretaceous, and this resulted in brittle fracture of the Paleozoic rocks, which is likely responsible for structurally-controlled gold mineralization in the south Klondike area including the White Gold exploration camp (MacKenzie et al, 2008; MacKenzie and Craw, 2010; MacKenzie and Craw, 2012).

Major units in the Klondike area include: the Snowcap (Nasina) Assemblage, the Klondike Series, the Slide Mountain (Moosehide) Assemblage, upper Cretaceous Carmacks Group volcanics/volcanoclastics, and Eocene intrusives (Figure 3). The basement unit is the Snowcap (Nasina) Series, consisting of metamorphosed schist and quartzite. It is overlain by the Klondike Series, a dominantly quartzofeldspathic schist of Early Permian (280 m.y.) age. Mid-Permian Sulphur Creek orthogneiss cuts the Klondike Schist extensively along Sulphur Creek. In the south and west Klondike, the Klondike Series is in contact with Late Devonian to Mississippian Simpson Range orthogneiss. Structurally overlying the Klondike and Nasina Series are greenstone and altered ultramafic of the Slide Mountain (Moosehide) Assemblage. In the east and south Klondike, upper Cretaceous andesitic volcanics and clastic sediments occur. These units are intruded by Eocene age rhyolite and diorite dykes and sills. Significant lode gold has been found throughout the Klondike and south Dawson areas (Chapman et. al., 2011 and others). The precise relationship between lode gold sources and local placer gold deposits is enigmatic and has been the subject of many scientific studies.

Local Bedrock Geology

Figure 3 shows the bedrock in lower Sulphur Creek as Sulphur Creek orthogneiss (map unit PqS). This is bounded on the east by Snowcap (Nasina) assemblage quartzite and schist (map unit PDS1) and on the west and north by Klondike Schist (map units PK1 and PK2).

Quaternary History

Most of the Klondike region has not been glaciated (Duk-Rodkin, 1999; Jackson et al., 2001). However, the marginal effects of a pre-Reid glaciation deposited glaciofluvial gravel along Australia Creek and Indian River. These were sourced from meltwater channels which breached the divide in the headwaters to the east. There is no evidence that glacial ice advanced into the drainage, although the pre-Reid glaciofluvial terraces covered pre-existing Tertiary White Channel gravels. These are especially evident in downstream reaches above Indian River (Froese and Jackson, 2005).

Surficial Geology

The surficial geology of the project area was mapped by Froese and Jackson (2005). Along Sulphur Creek are surficial units of several ages and types, shown in Figure 4. These include: CEaP/AtT (Pleistocene colluvial-aeolian sediments overlying Tertiary alluvial terrace sediments), CEaP (Pleistocene colluvial-aeolian sediments), AtP (Pleistocene alluvial terrace), ACxP (Pleistocene alluvial/colluvial complex), Ax (alluvial complex), Cx (colluvial complex), Cl (landslide) and Cb-v (colluvial blanket-veneer). In general, the AtT (Tertiary alluvial terrace) units are more prevalent downstream, whereas upstream reaches are dominated by ACxP (Pleistocene alluvial/colluvial complex) and Cx (colluvial complex). The Jon Property is mapped as Ax (Alluvial Complex) in the valley centre, which is flanked on the left limit by Cb-v (colluvial blanket-veneer) and O (organic).

Placer Geology

Placer gravels in Dominion Creek and its tributaries (Gold Run and Sulphur) can be characterized by 5 types of deposits: Pliocene White Channel gravel; Pleistocene terraces; early Pleistocene incised-valley gravel (Ross gravel); Pleistocene Dominion Creek gravel; and creek and gulch deposits (Froese et al., 2001).

Nearby Tatra Ventures Ltd. in Sulphur Creek valley is described by Van Loon and Bond (2014) as having stratigraphy as follows: Chlorite schist bedrock overlain by 0 to 3.8 m (0 to 12.5 ft.) of massive, imbricated clast-supported gravel with subrounded to rounded cobble and boulder-sized clasts (predominantly quartz), overlain by 3.8 to 7.0 m (12.5 to 23.0 ft.) of crudely-stratified, moderately-oxidized pebble-cobble gravel, overlain by 7.0 to 8.5 m (23.0 to 27.9 ft.) of black muck. The entire stratigraphic section in this part of the main valley is up to 20 metres thick. On the left limit tributary (Jon property), the thickness of sediment overlying bedrock is likely significantly thinner.

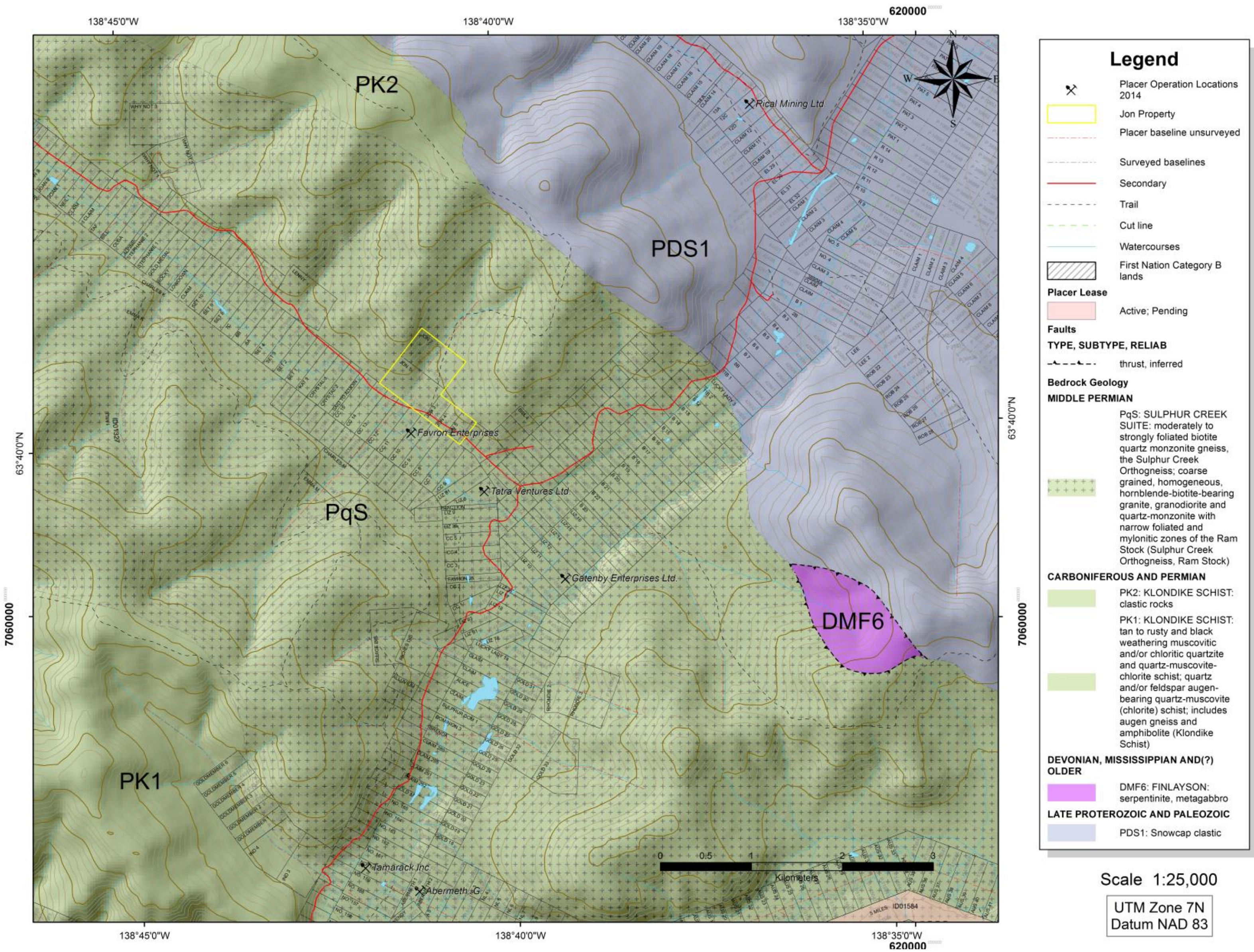

Figure 3 - Bedrock Geology of Jon Property, lower Sulphur Creek area, after Yukon Geological Survey (2016).

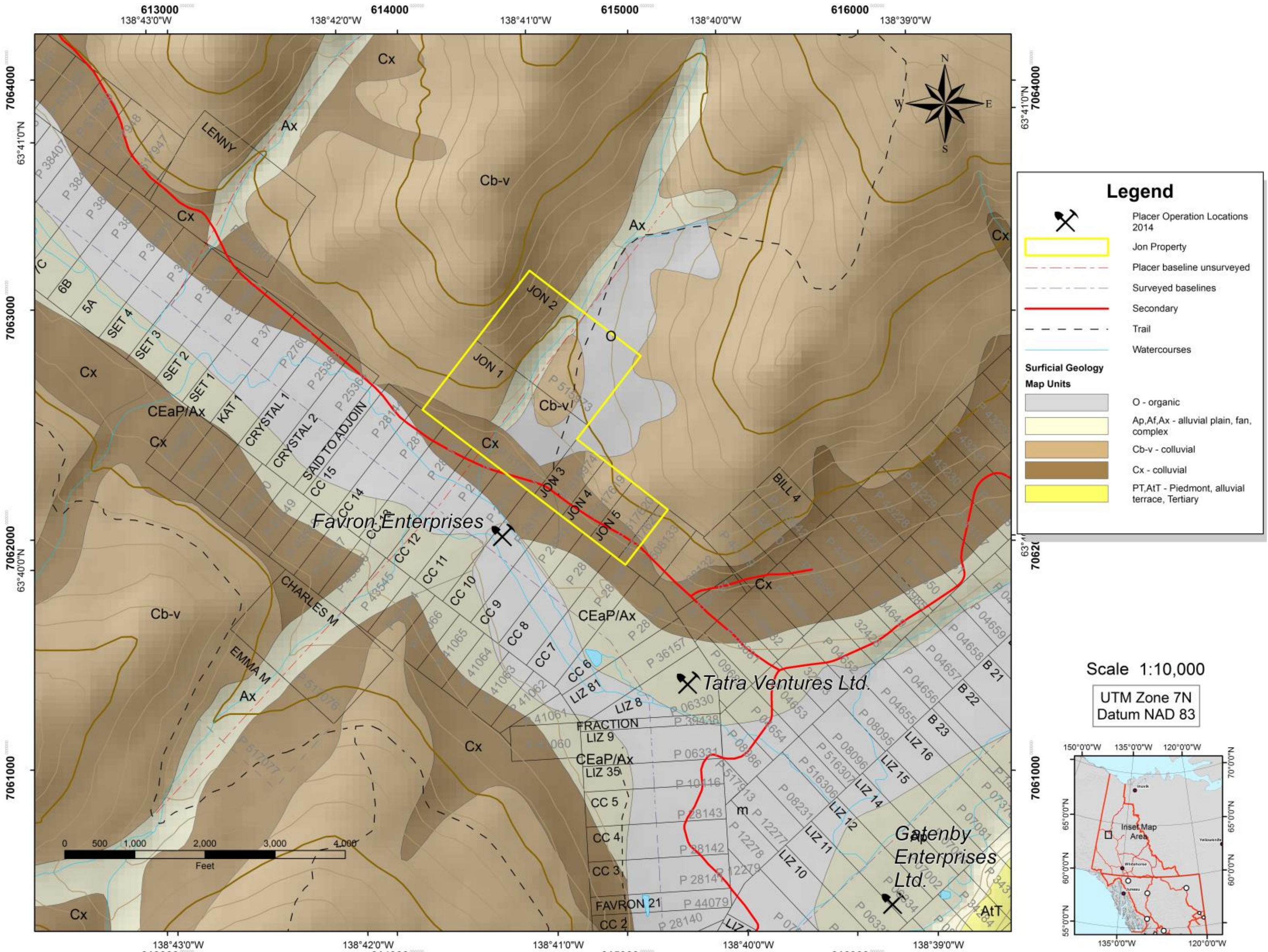

Figure 4 - Surficial Geology, Jon Property, lower Sulphur Creek, after Froese and Jackson (2005).

Previous Exploration

Resistivity Geophysical Surveys – 2014 and 2015

Kryotek Arctic Innovation Inc. conducted one resistivity survey on the Jon placer property in July 2014 (Coates, 2014), and two resistivity surveys in 2015 (Coates, 2015). The surveys were conducted using a Lippmann 4-point Resistivity System. Interpreted profiles from those surveys are shown as Figures 5 to 7, and the locations of those surveys are shown on Figure 8.

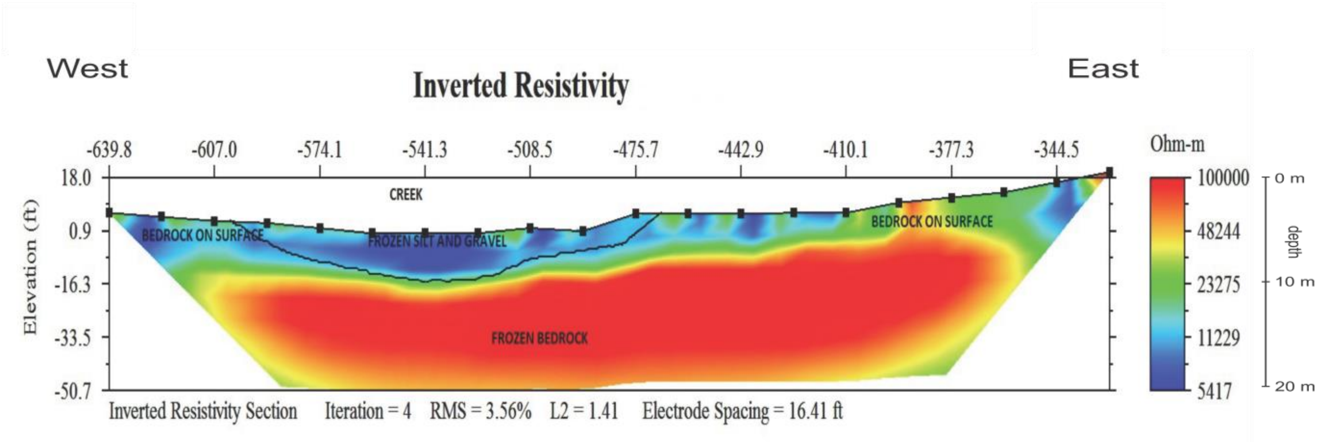

Figure 5 - Resistivity profile JCA, Kryotek Arctic Innovation Inc., July 2014. Vertical scale shown in feet and metres.

Figure 5 shows resistivity profile JCA conducted by Kryotek in July 2014. Bedrock was exposed at the surface on both ends of the profile. Permafrost was present throughout, with electrodes encountering frozen ground 20-30 cm below the surface. High resistivity bedrock was encountered at depths of 0-16 feet (0 – 5 m). Below the current creek, bedrock was interpreted at a depth of 16 feet (5 m) below the surface in a channel approximately 60 feet (20 m) wide.

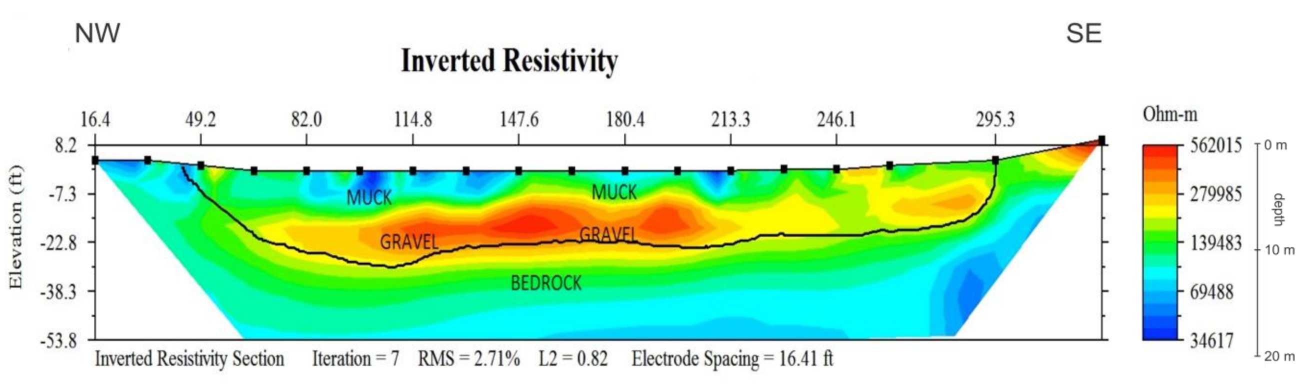

Figure 6 - Resistivity profile JCB, Kryotek Arctic Innovation Inc., April 2015. Vertical scale shown in feet and metres.

Figure 6 shows resistivity profile JCB conducted by Kryotek in April 2015. Bedrock is exposed on the surface near the start of the line in the NW. Permafrost was present throughout. Interpreted bedrock is between 15 and 20 feet (4.6 – 6.1 m) below the surface. Overlying this is interpreted to be 8-10 feet (2.4 – 3 m) of sand and gravel followed by 5-7 feet (1.5 – 2.1 m) of frozen black organic "muck".

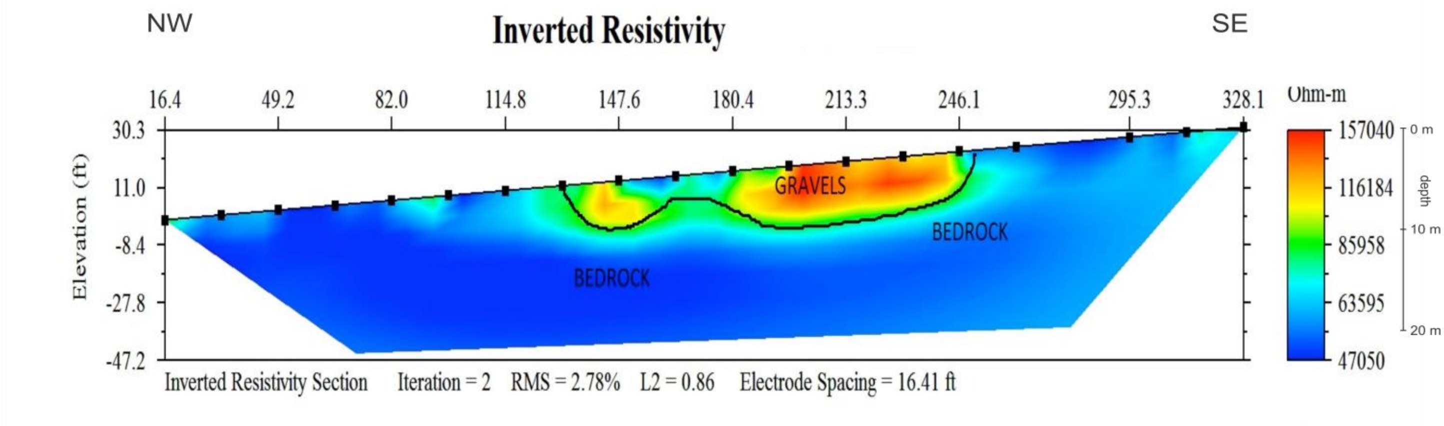

Figure 7 - Resistivity profile JCC, Kryotek Arctic Innovation Inc., April 2015. Vertical scale shown in feet and metres.

Figure 7 shows resistivity profile JCC conducted by Kryotek in April 2015. Interpretation by J. Coates shows a possible channel with depths to bedrock of 7-12 feet (2.1 – 3.7 m). Alluvial gravels are interpreted at depths of 3-12 feet (0.9 – 3.7 m) and are covered by colluvium. Permafrost was encountered throughout the survey line.

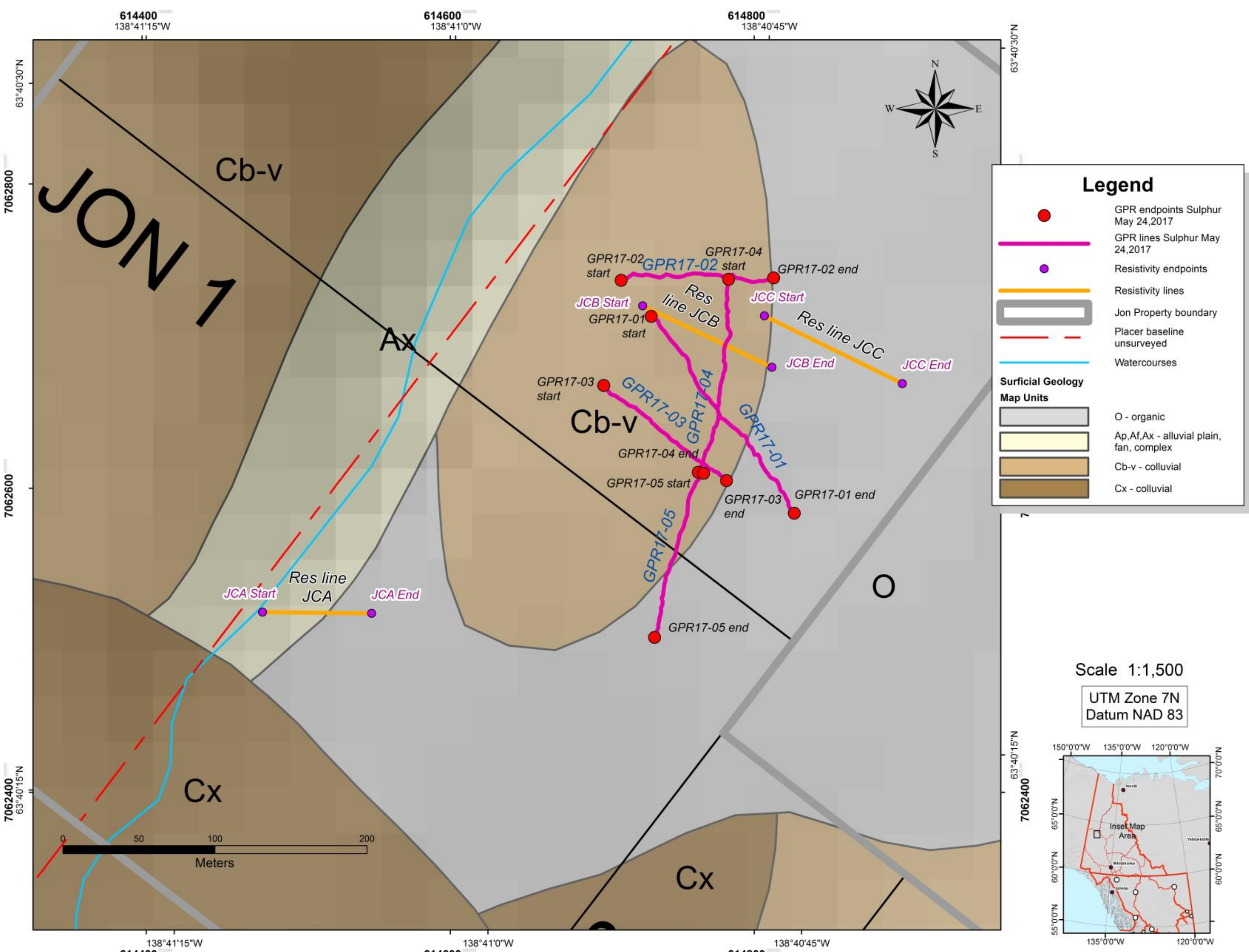

Figure 8 - Map of Jon property showing surficial geology (after Froese and Jackson, 2005), 2014 and 2015 resistivity lines and 2017 ground penetrating radar lines.

2017 Placer Exploration Program – Ground Penetrating Radar

General

The ground penetrating radar surveys were conducted on the property by William LeBarge (Geoplacer Exploration Ltd.) on May 24, 2017, with field assistance provided by Bud Davis and Jon Davis. Five (5) lines were surveyed. The combined total length of the 5 lines is 630 metres. The locations of the surveyed lines are shown on Figures 8 and 14, the interpreted profiles are shown in Figures 9 to 13 and the coordinates of the endpoints are given in Table 3.

Methodology

The ground penetrating radar survey was conducted with the “EasyRad Pro” instrument, an above-ground (non-contact) GPR with two antennas (dipole-type), having an antenna separation of 1.22 m. The instrument operated at a frequency of 100 MHz, and data was gathered in real-time on an Android tablet running the EasyRad software application. Results were saved in the proprietary EasyRad format, as well as in the universal SEG-Y format. The results of the survey were analyzed using Prism 2.1 GPR software. Survey data was corrected for distance to ground, move-out (distance between antennae), background noise and data outliers. Amplification was applied for weak signal returns (a factor occurring with increasing depth). The GPR lines were georeferenced in the field using a hand-held Garmin GPS, which recorded the tracks of the lines as well as the start and endpoints of the surveys.

Scientific Basis

Relative Permittivity is the defined as the ability of a material to store and then permit the passage of electro-magnetic energy (e.g. ground-penetrating radar) when a field is imposed on the material (Baker et al, 2007). Air has a relative permittivity of 1, while water has a relative permittivity of 80 (Davis and Annan, 1989). All other earth materials fall between these values. Table 2 illustrates relative permittivity values for selected geologic materials.

Table 2 - Relative Permittivity (ε_r) of selected geologic materials, after Davis and Annan, (1989).

Material	Relative Permittivity (ε_r)
Air	1
Water	80
Fresh water ice	3–4
Sea water ice	4–8
Snow	8–12
Permafrost	4–8
Sand, dry	3–5
Sand, wet	20–30
Sandstone, dry	2–3
Sandstone, wet	5–10
Limestone, dry	7
Limestone, wet	8
Shales	5–15

Material	Relative Permittivity (ε_r)
Shale, wet	6–9
Silts	3–30
Clay, dry	2–6
Clay, wet	15–40
Soil, sandy dry	4–6
Soil, sandy wet	15–30
Soil clayey dry	4–6
Soil, clayey wet	10–15
Coal, dry	3.5
Coal, wet	8
Granites	4–6
Salt, dry	5–6

The interpreted contacts in the GPR profiles are reflections from interfaces caused by contrasts in relative permittivity, and they are affected by variations in permeability, porosity, saturation, permafrost and ice content (Baker et al., 2007). These are related to differences in mineralogy, lithology and grain size, but it is the interaction of these factors which may complicate any stratigraphic interpretations from GPR results (Martinez and Byrnes, 2001).

The depth of penetration of GPR into permafrost is usually greater than that for thawed sediments. For example, Bristow et al., (2000) report less than 20 m of penetration in sand deposits in temperate or tropical environments, while Arcone et al. (2002) report a penetration depth of nearly 80 m in marginally frozen and stratified alluvial sands in Fairbanks, Alaska. Arcone et al. (2002) also report that the relative permittivity at their sites in Fairbanks ranged between 4 and 5.5, with the lowest permittivity values indicating higher ice contents in excess of 70%. This is comparable with the average ice content of 70% for frozen silt ("black muck") in the Klondike region including the project area (Kotler, 1998). In addition, since the dielectric permittivity of ice (3.2) is appreciably different from that of the surrounding soil or permafrost matrix (4–5.5, as discussed above), massive deposits of ground ice will give strong radar reflections in the GPR profiles (Lawson et al, 1999).

Limitations and Disclaimer

Interpretations are limited by the accuracy of the system and methods employed. Interpreted results are always improved with calibration by drill holes or other geophysical methods. Data is more accurate toward the surface and becomes more uncertain with increasing depth. Interpretation of GPR data should be performed with caution and verified by other methods whenever possible. The accuracy of the information presented is not guaranteed and all mine development is the client's responsibility. William LeBarge of Geoplacer Exploration Ltd. accepts no liability for any use or application of these data by any and all authorized or unauthorized parties.

Table 3 – Coordinates and Lengths of Ground Penetrating Radar Lines, Sulphur Creek.

Ground Penetrating Radar Lines- Sulphur Creek, October 2017										
Date	Line Name	Length (m)	Start Point				End Point			
			Latitude	Longitude	y	x	Latitude	Longitude	y	x
2017-05-24	GPR17-01	164	63.67349	-138.68095	7062713	614732	63.67230	-138.67915	7062583	614826
2017-05-24	GPR17-02	108	63.67371	-138.68134	7062737	614712	63.67369	-138.67931	7062738	614813
2017-05-24	GPR17-03	105	63.67310	-138.68162	7062668	614701	63.67251	-138.68003	7062605	614782
2107-05-24	GPR17-04	135	63.67370	-138.67991	7062737	614783	63.67257	-138.68041	7062611	614763
2017-05-24	GPR17-05	118	63.67256	-138.68034	7062610	614767	63.67160	-138.68107	7062502	614734

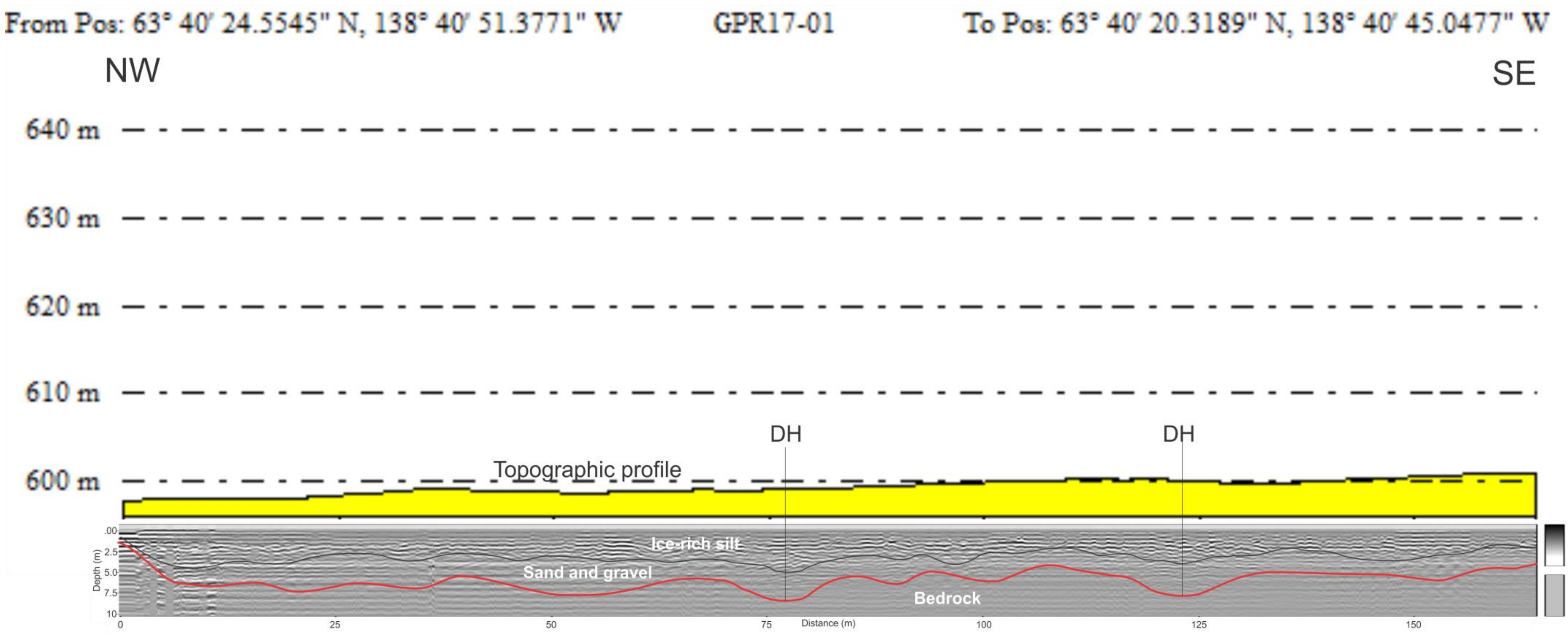

Figure 9 - Ground Penetrating Radar Line GPR17-01. Horizontal and vertical scales are 1:1. Topographic profile is shown above the radar profile, which was not topographically corrected as the topography is gentle and flat. The red line on the profile represents the interpreted bedrock contact, while the black line on the profile indicates the interpreted boundary between the permafrost/black muck and the possible sand and gravel below. Two possible drill targets are shown on the profile.

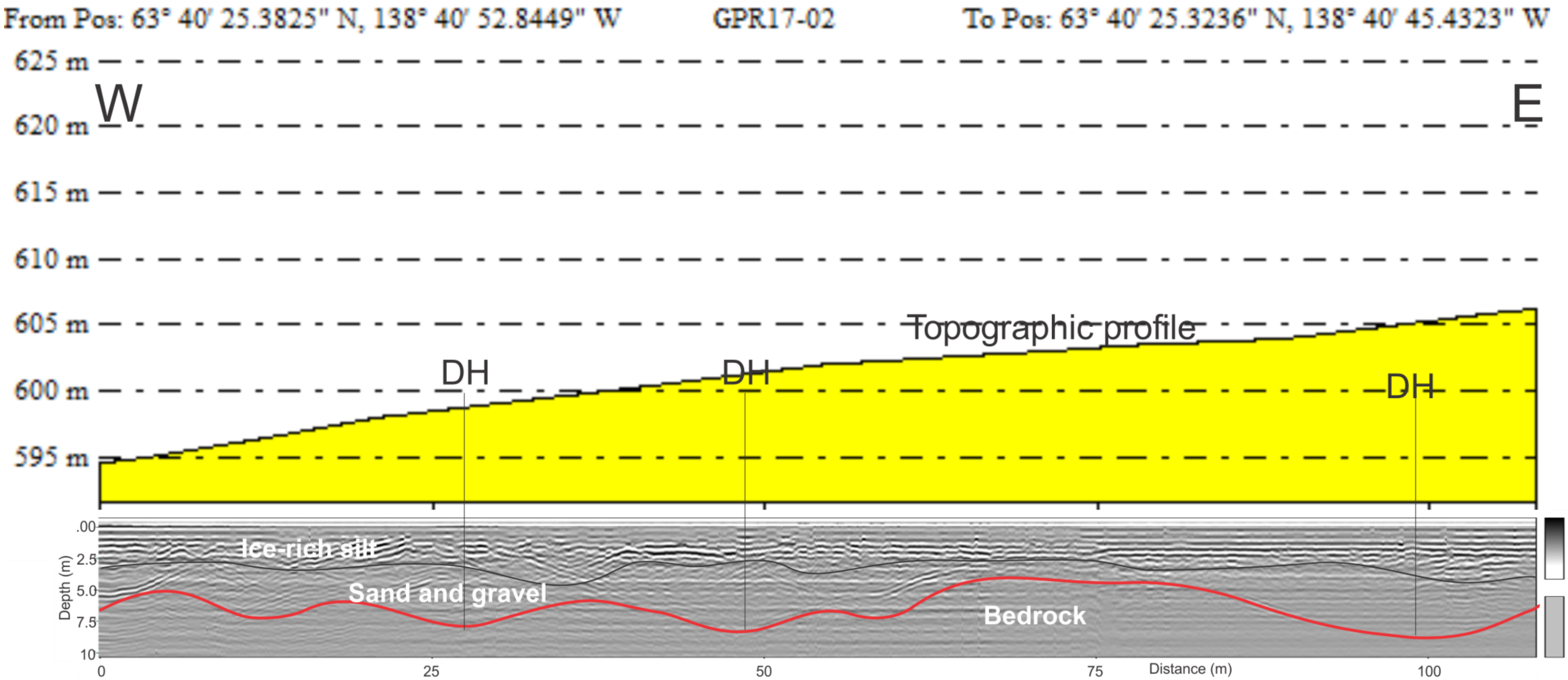

Figure 10 - Ground Penetrating Radar Line GPR17-02. Horizontal and vertical scales are 1:1. Topographic profile is shown above the radar profile, which was not topographically corrected as the topography is relatively flat. The red line on the profile represents the interpreted bedrock contact, while the black line on the profile indicates the interpreted boundary between the permafrost/black muck and the possible sand and gravel below. Three possible drill targets are shown on the profile. The easternmost target may be shallower than shown if topographically corrected.

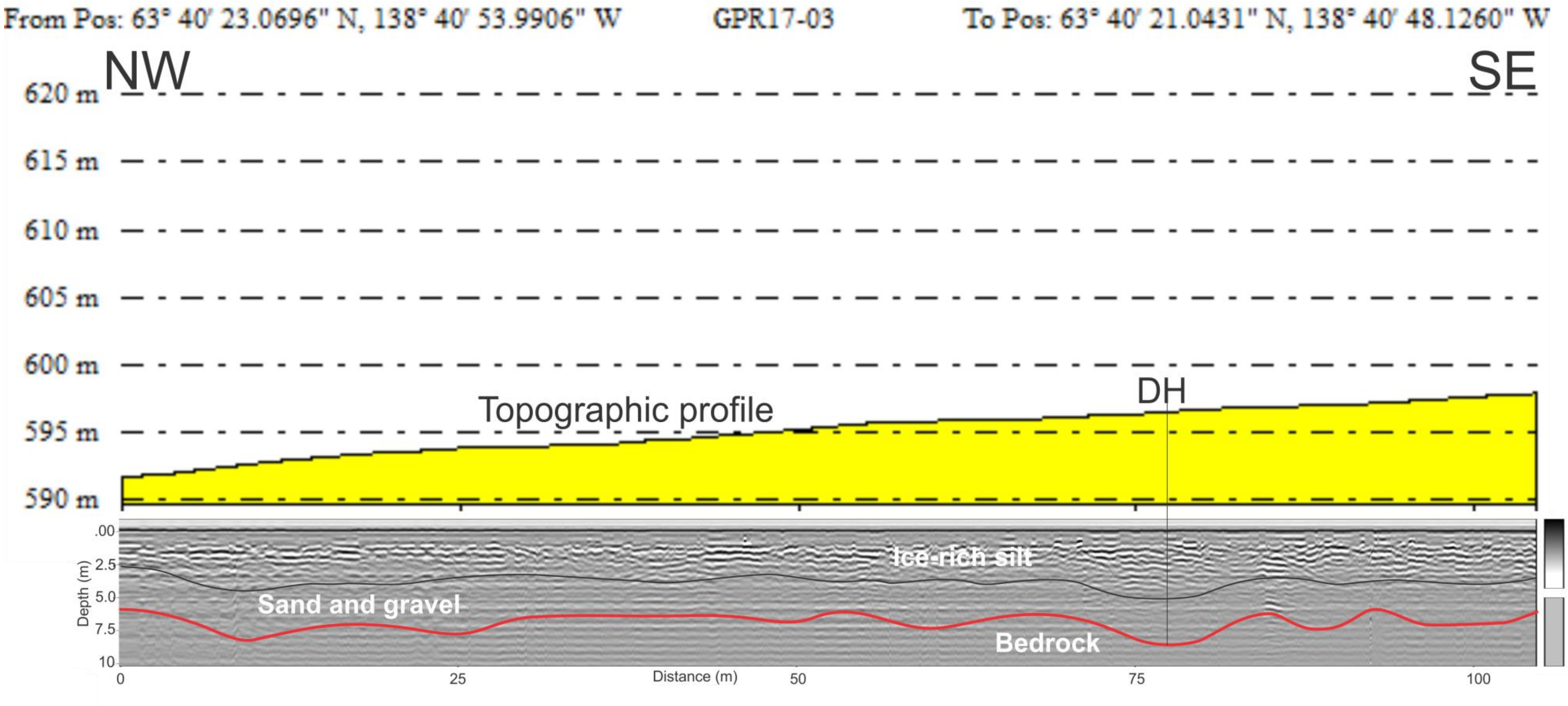

Figure 11 - Ground Penetrating Radar Line GPR17-03. Horizontal and vertical scales are 1:1. Topographic profile is shown above the radar profile, which was not topographically corrected as the topography is relatively flat. The red line on the profile represents the interpreted bedrock contact, while the black line on the profile indicates the interpreted boundary between the permafrost/black muck and the possible sand and gravel below. One possible drill target is shown on the profile.

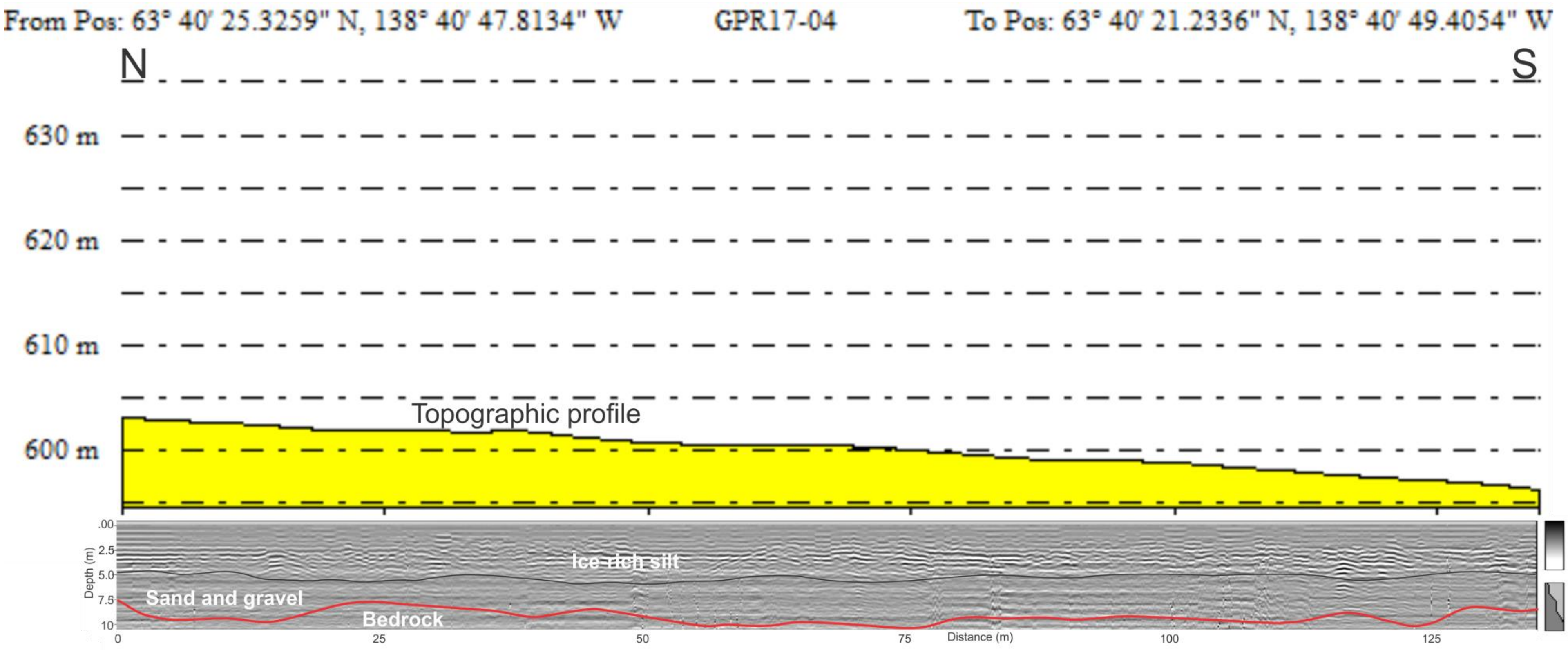

Figure 12 - Ground Penetrating Radar Line GPR17-04. Horizontal and vertical scales are 1:1. Topographic profile is shown above the radar profile, which was not topographically corrected as the topography is relatively flat. The red line on the profile represents the interpreted bedrock contact, while the black line on the profile indicates the interpreted boundary between the permafrost/black muck and the possible sand and gravel below. The bedrock is relatively flat (following topography) and no obvious drill targets are noted.

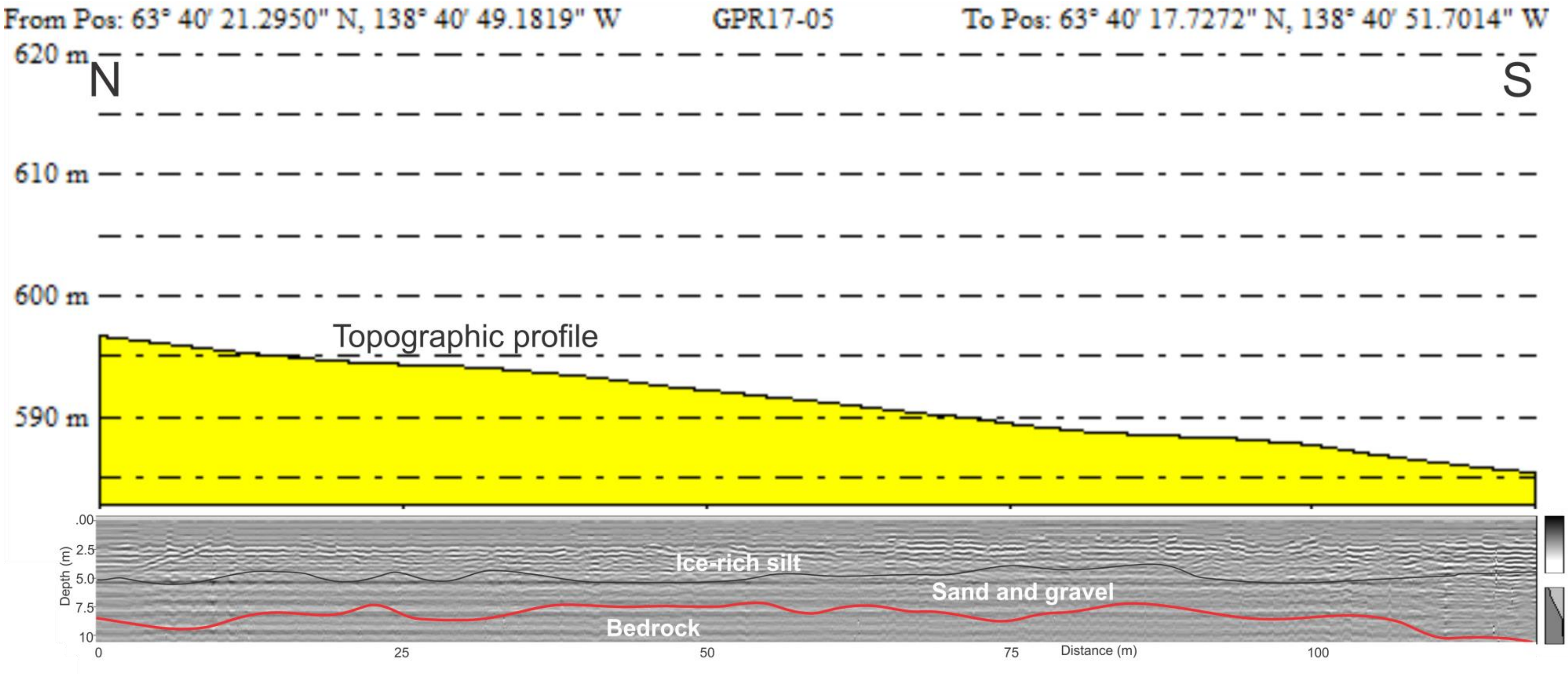

Figure 13 - Ground Penetrating Radar Line GPR17-05. Horizontal and vertical scales are 1:1. Topographic profile is shown above the radar profile, which was not topographically corrected as the topography is a gentle continuous slope. The red line on the profile represents the interpreted bedrock contact, while the black line on the profile indicates the interpreted boundary between the permafrost/black muck and the possible sand and gravel below. The bedrock is relatively flat (following topography) and no obvious drill targets are noted. Sediments appear to thicken towards the valley centre to the south.

Discussion of Results

For the reasons discussed in the previous section, an average permittivity of 4 was assigned to the GPR data for the purposes of processing and interpretation. This resulted in an interpreted depth of penetration of 10.5 metres in the GPR profiles. It should be emphasized that drill hole verification of lithologies, sediment thicknesses and depths to bedrock would enable recalculation of these factors and allow a more accurate interpretation of GPR results.

Figures 9 to 13 show the interpreted GPR profiles along with the overlying topography. Overall, interpreted depths to bedrock varied between 2 metres and 10 metres.

The ground penetrating radar surveys appeared to delineate not only the bedrock contact but also the boundary between the ice-rich permafrost silt (“black muck”) and underlying ice-poor sediments (possibly sand and gravel) below. Ice-rich permafrost was evident as a zone of strong reflection in each survey profile, approximately 2 metres below the surface and extending to a depth of 5 to 6 metres below the surface.

The profiles which were perpendicular to the dip of the slope (GPR17-01 to GPR17-03) showed the most variations in bedrock topography and thus had more possible paleochannel targets. There was a strong correlation between the easternmost paleochannel target on GPR17-02 and GPR17-03; and the centre paleochannel target on GPR profile GPR17-01. This distinctive paleochannel is the most prospective of all targets, and the trend of it is shown on Figure 14.

Comparison of the previously-conducted resistivity surveys and the present GPR surveys also showed some similarities. For instance, the above-mentioned paleochannel which appears on GPR profiles GPR17-02, GPR17-01 and GPR17-03 may be coincident with a paleochannel identified on resistivity profile JCB. However, the GPR surveys tended to show bedrock deeper than corresponding nearby resistivity surveys.

The profiles which were surveyed down-slope (GPR17-04 and GPR17-05) showed little variation in bedrock topography. Bedrock in those profiles appeared to descend correspondingly with the overlying topography, although sediments thickened towards the valley centre.

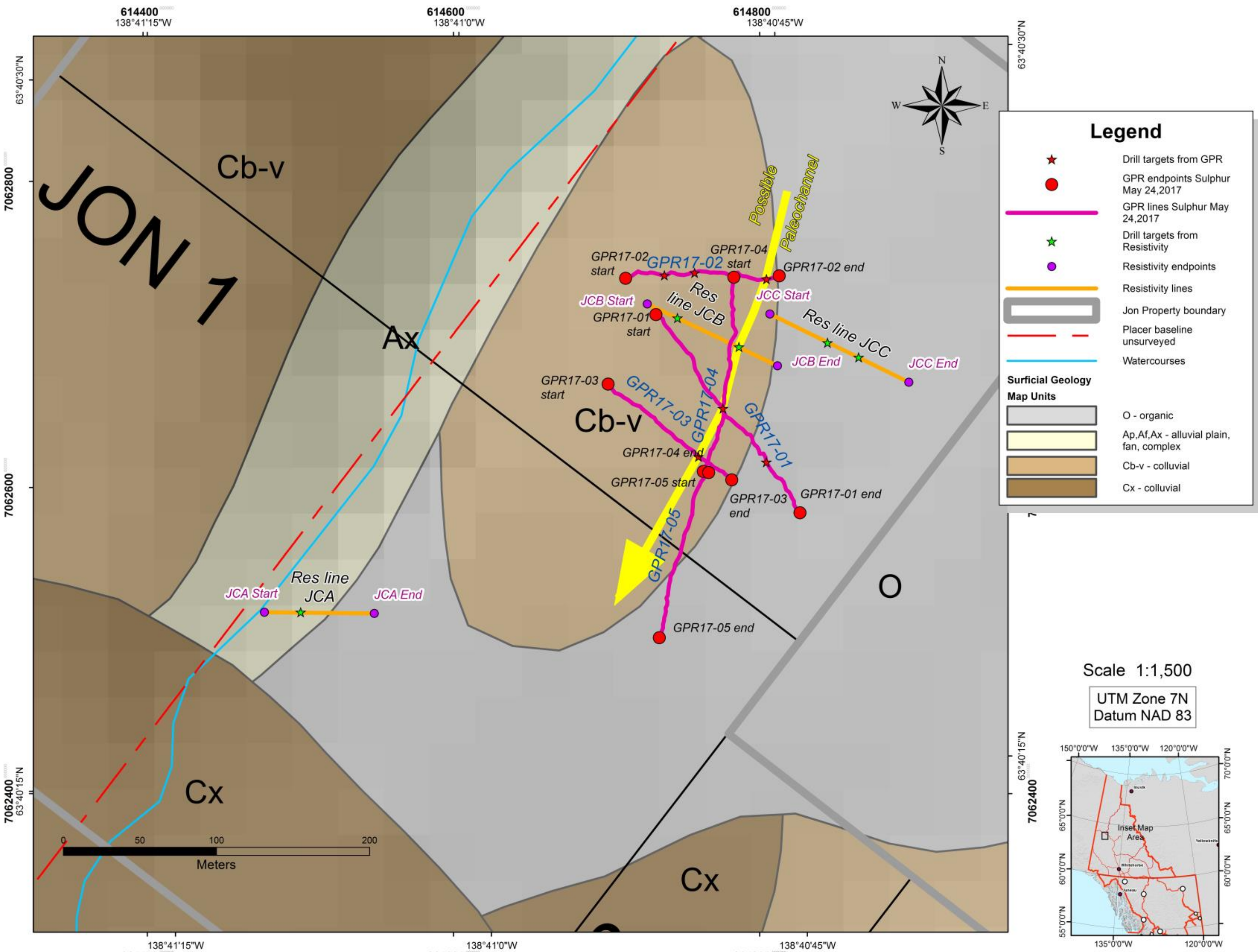

Figure 14 - Map of Jon property showing proposed drill targets from resistivity lines and ground penetrating radar lines, as well as a possible paleochannel (in yellow).

Conclusions and Recommendations

An interpretive map showing possible drill targets generated from both the GPR and Resistivity surveys, as well as a possible paleochannel is given as Figure 14. Coordinates for the drill targets noted on profiles GPR17-01, GPR17-02, and GPR17-03 and resistivity profiles JCA, JCB and JCC are shown in Table 4 below.

Table 4 - Coordinates for drill hole targets generated from GPR and Resistivity profiles.

Target Name	Line	Type	Latitude DD	Longitude DD	UTM_N	UTM_E
GPR17-01-1	GPR17-01	GPR	63.67293	-138.68	7062652	7062652
GPR17-01-2	GPR17-01	GPR	63.67261	-138.68	7062617	7062617
GPR17-02-1	GPR17-02	GPR	63.67372	-138.681	7062739	7062739
GPR17-02-2	GPR17-02	GPR	63.67373	-138.68	7062740	7062740
GPR17-02-3	GPR17-02	GPR	63.67368	-138.679	7062736	7062736
GPR17-03-1	GPR17-03	GPR	63.67265	-138.68	7062620	7062620
JCA-01	JCA	Resistivity	63.67183	-138.686	7062519	614500.2
JCB-01	JCB	Resistivity	63.67347	-138.681	7062711	614746.3
JCB-02	JCB	Resistivity	63.67329	-138.68	7062692	614786.5
JCC-01	JCC	Resistivity	63.67329	-138.679	7062695	614844.3
JCC-02	JCC	Resistivity	63.6732	-138.678	7062685	614864.7

Overall, ground penetrating radar appeared to provide good signal response and interpretable data in the project area, although it is not recommended as a stand-alone method to interpret lithologies and depths to bedrock. Confirmation of interpretations by drilling is recommended. This would verify lithologies, sediment thicknesses and depths to bedrock, which would enable recalculation of the relative permittivity and depths of penetration and result in a more accurate interpretation of initial GPR results.

Testing of the alluvial materials for placer gold is recommended. Initially, drilling methods such as auger (6 inch or larger), R/C (Reverse circulation) or RAB (Rotary Air Blast) may be used. This should be followed up by excavator test pitting and bulk processing of prospective alluvial gravels.

Statement of Costs, 2017 Placer Exploration Program, Jon Property

Table 5 - Statement of Costs, Placer Exploration, Jon Property

2017 Placer Exploration Program Statement of Costs	Item	Rate	Subtotal	GST	Total	Invoice #
GPR data acquisition and compilation including report, Sulphur Creek	0.630 line km	$1500/line km	$945.00	$47.25	$992.25	2017-08
Mobilization/Demobilization to property from Whitehorse	1200 km	0.60	$720.00	$36.00	$756.00	2017-08
Accommodation in Dawson, 2 nights	2 days	$129.00	$258.00	$12.90	$270.90	2017-08
Total					$2019.15	

Statement of Qualifications

William LeBarge

I, William LeBarge, of 13 Tigereye Crescent, Whitehorse, Yukon, Canada, DO HEREBY CERTIFY THAT:

1. I am a Consulting Geologist with current address at 13 Tigereye Crescent, Whitehorse, Yukon, Canada, Y1A 6G6.
2. I am a graduate of the University of Alberta (B.Sc., 1985, Geology) and the University of Calgary (M.Sc., 1993, Geology – Sedimentology)
3. I am a Practicing Member in Good Standing (#37932) of the Association of Professional Engineers and Geoscientists of British Columbia (APEGBC).
4. I have practiced my Profession as a Geologist continuously since 1985.
5. I am President and sole shareholder of Geoplacer Exploration Ltd., a Yukon Registered Company.

Dated this 8th day of November, 2017

William LeBarge, P. Geo.

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