

**THE PETROGRAPHIC ANALYSIS OF
THE HOST ROCK
AND VAESITE BEARING HORIZON ON
THE NICK PROPERTY, YUKON TERRITORY**

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An Undergraduate Thesis

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April 24, 1989

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A Petrographic Study of the Host Rocks and Mineralized Horizon, of the Nick Property, Yukon Territory.

Abstract:

The Nick Property hosts a stratigraphic unit rich in nickel, zinc, and platinum group elements. The stratigraphy, nature and textures of the Nick Property are similar to sedimentary exhalative deposits of the Selwyn Basin. The spacial relationships of the metallic phases will be discussed in conjunction with assay data from drill core and handsamples from outcrop to provide a better understanding of the nature of the mineralized zone. Discussions will be included concerning the distribution of metalliferous phases, such as: (1) the distribution of lead, zinc and nickel contained in the diagenetic pyrite, (2) the nickel and platinum group element associations of this deposit type, and its application to further study and exploration, and (3) similarities between the geological relationship of the Nick Property to other deposits. The results of this study should produce a better understanding of the geology of the property and the ore forming system.

INTRODUCTION:

This geologic research report is based on petrographic and x-ray powder diffraction analysis carried out on samples taken from the Nick Property, Yukon Territory. The analysis of the petrographic thin sections and polished sections provides a useful tool for the interpretation of textural and mineralogical relationships between the metallic phases of a deposit. The use of X-ray diffractometry is useful for the identification of relatively large, pure samples. However, its usefulness diminishes with the complexity of the mineralogy of the sample. For example, the intergrown nature of the fine grained

sulfides (10-30 um) could not be separated. The identification of clay and silicate phases can be done using this method.

A literature search will be undertaken to compare the Nick deposit to a deposit of a similar nature. Additional information regarding the nature of the deposit will be utilized from prior field work completed by Archer, Cathro & Associates Ltd. during the 1988 field-season. The mineralogy of the Nick deposit will be defined and compared to various types of sedex and stratiform sedimentary mineral deposits.

REGIONAL GEOLOGY:

The Nick property is located in the Northern extremity of the Selwyn Basin, in an outlier of basinal sediments on the MacKenzie Carbonate Platform. The property is located at 64 degrees 43' N latitude, 135 degrees 13' W longitude, on the northeast margin of the Selwyn Basin. (see map, figure .1)

The stratigraphy of the region is composed of a series of clastic and chemically deposited sedimentary rocks. The top of the sequence is comprised of a shale unit which is composed of two parts. The upper unit is a brown weathered shale with minor fine grained chert pebble conglomerates that is equivalent to Green's unit 14 (Green, 1972). The upper shale unit is thought to be equivalent to the Upper Earn Formation. The lower unit is a thinly bedded black to silvery black weathered shale with underlying black chert and limestone, equivalent to unit 13 (Green, 1972). The lower shale unit is equivalent to the Lower Earn Formation. These two units can be correlated with the Canol and Imperial Formations in the Selwyn Basin. Underlying the shale units in the area is a black, thinly-

bedded limestone unit which is locally known as the Limestone Ball Member. Underlying the Limestone Ball Member is a calcareous shale unit. This unit mapped by Green (1972), is known as unit 12, and is equivalent to the Road River Formation.

The structure of the area may be controlled by post depositional normal faults which are thought to be responsible for the preservation of this shale outlier. The underlying limestone unit has a strong control on the synform in the region. The Devonian to Permian age shales in the region overlie the thick Ordovician to middle Devonian limestones. The remapping of the area has determined that the two outliers to the north and east of the main shale unit 14, (Green, 1972), are actually a parallel basin and an extension of the main shale unit, respectively (Cathro, 1988) (see fig. 2 & 3). The Paleozoic sediments of the syncline are in fault contact with the Precambrian MacKenzie Carbonate Platform to the east of the Nick Property. This fault is post depositional and may be responsible for the preservation of the shale outlier.

PROPERTY GEOLOGY:

Stratigraphy

The nickel bearing strata is overlain by interbedded brown argillaceous shale, silvery black siliceous shale, black carbonaceous and phosphatic cherts. This unit is divided into two subdivisions: the upper division is dominated by clastic sedimentary rocks, such as, brown argillites and occasional fine detrital cherty quartz sandstones. The upper clastic division is more recessive than the lower division. The lower division is composed mainly of chemically and biochemically precipitated sediments, such as, phosphatic and carbonaceous cherts, siliceous shales and minor

LEGEND

QUATERNARY

- 26 Unconsolidated glacial and alluvial deposits

CARBONIFEROUS TO PERMIAN

- 14 Black- and silvery-weathering shale; minor chert-pebble conglomerate

DEVONIAN TO CARBONIFEROUS

- 13 Black shale, chert and minor limestone

SILURIAN (?) TO MIDDLE DEVONIAN

- 12 Dark grey-weathering, black, thin-bedded, platy limestone

DEVONIAN

- 10 Light grey limestone and dolomite

ORDOVICIAN AND SILURIAN

- 9 Road River Formation: Black chert and black argillite

- 8 Grey- and buff-weathering dolomite and limestone

PROTEROZOIC

- 2 Dolomite, slate, minor phyllite and quartzite

- 1 Argillite, slate and phyllite

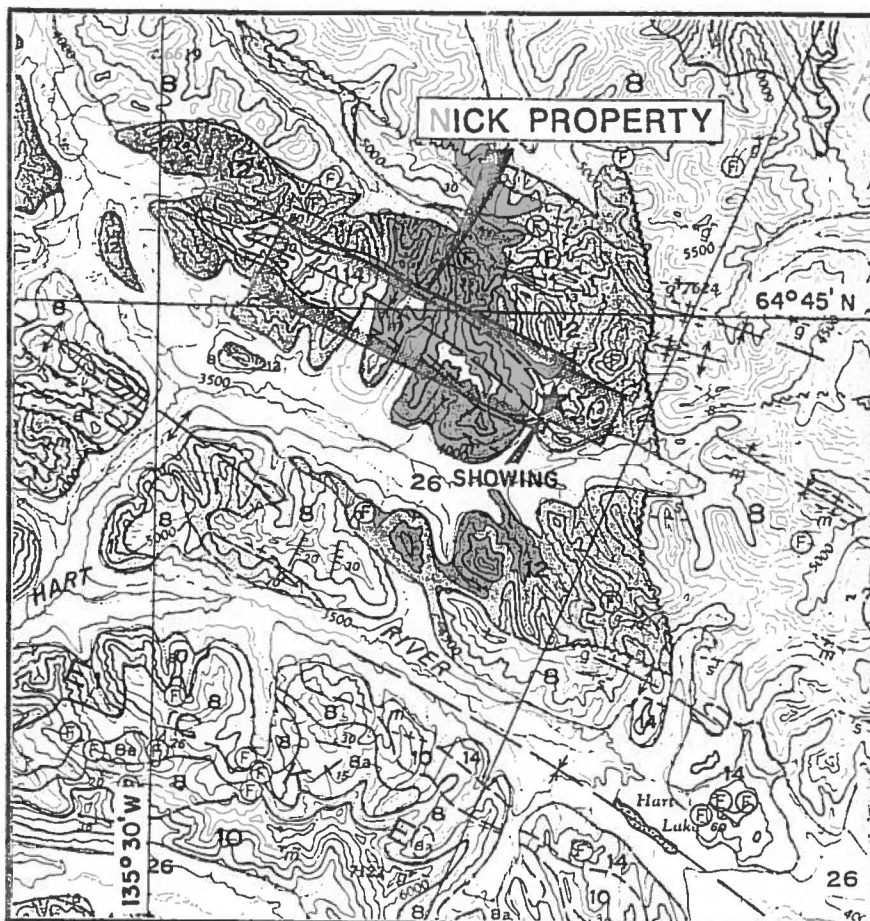


Figure 2

REGIONAL GEOLOGY

NICK PROPERTY, YUKON

NDU RESOURCES LTD.

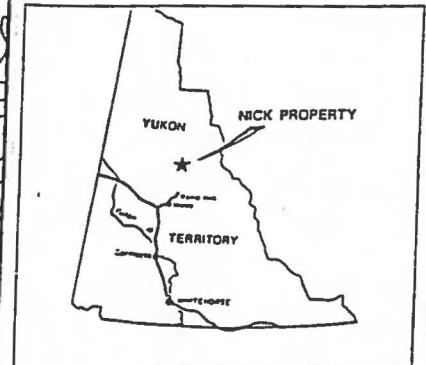
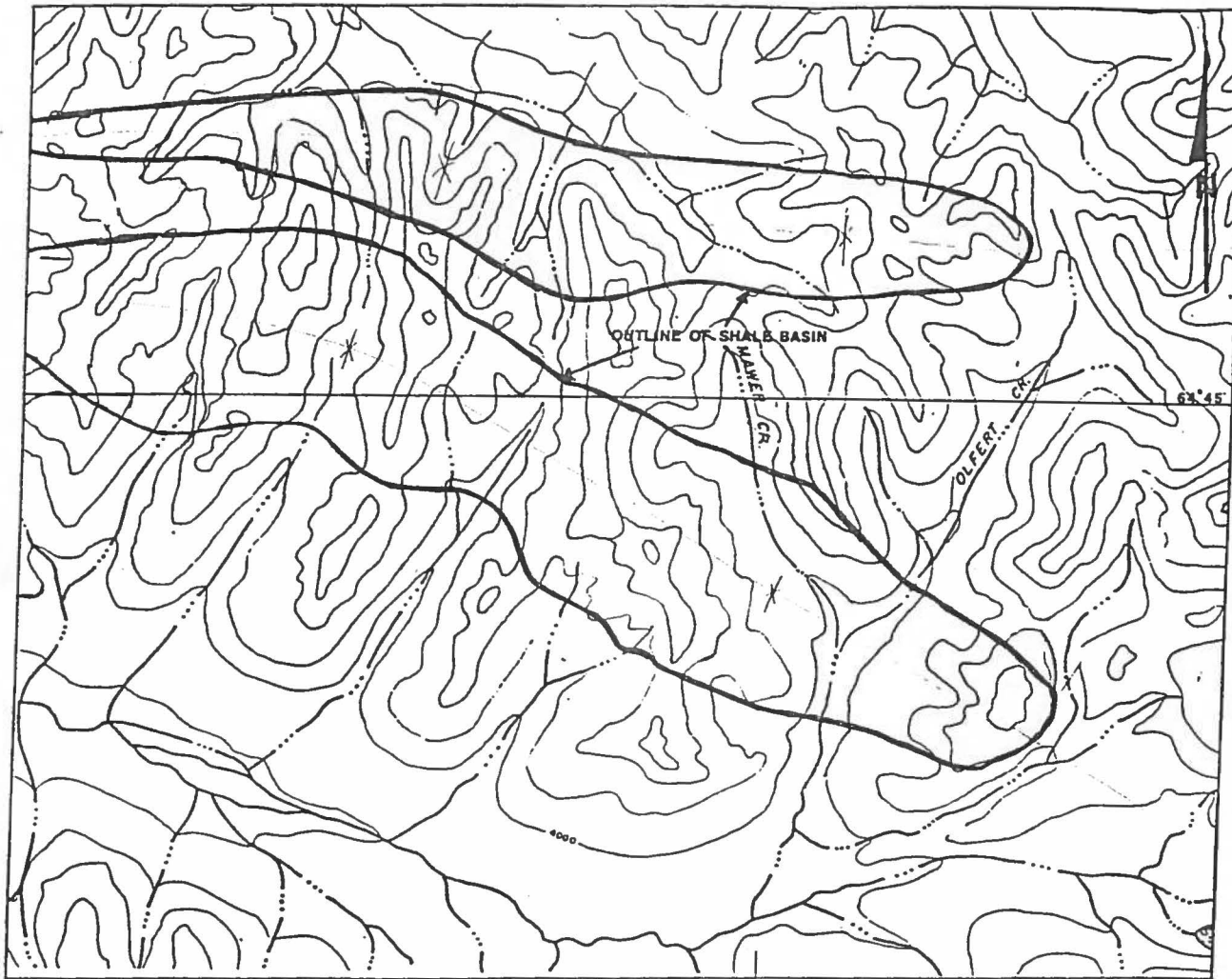
SCALE 1:250,000

0 5 10 15 Km

0 5 10 Miles

From GSC Map I282 A and Memoir 364
by L.H. Green, 1972

(report by R.J. CATHRO dated Aug./88)



OUTLINE OF MINERAL
SHALE OUTLIERS

Figure: 2.

ARCHER, CATHRO & ASSOCIATES (1991) LIMITED

NICK PROPERTY, YUKON

NDU RESOURCES LTD.

PAK-MAN RESOURCES LTD.

2001 RESOURCE INDUSTRIES INC.

SCALE 1:45,890



135°15'

calcareous shales. The lower division is up to 250 metres thick in the syncline (Cathro, 1988).

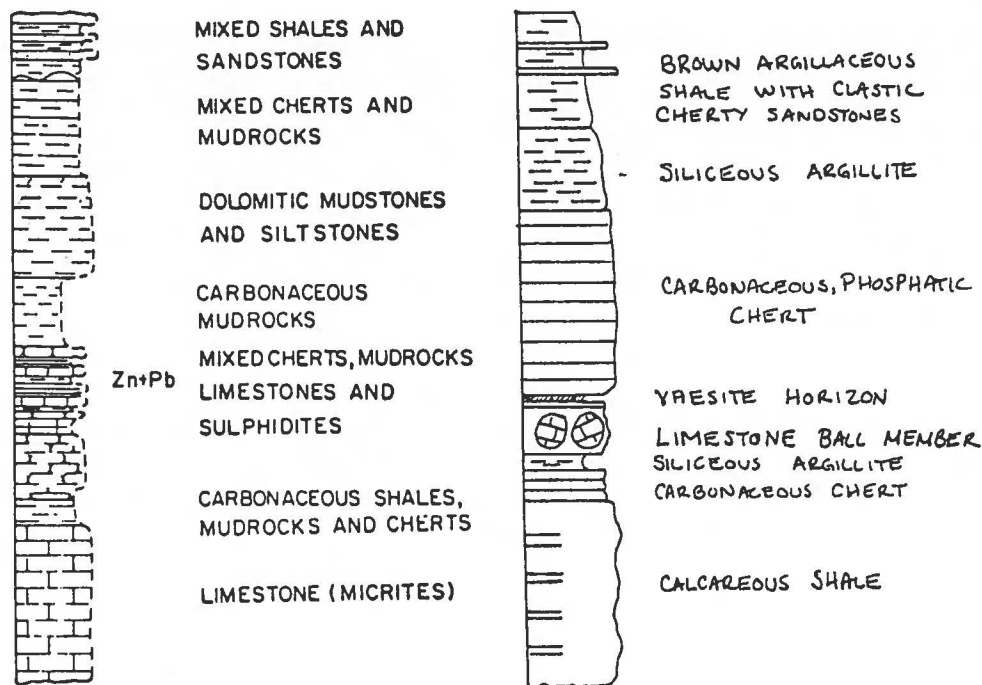
The nickel bearing strata is located stratigraphically above a thinly bedded limestone member. Locally this limestone unit is known as the "Limestone Ball Member". The Limestone Ball Member is a marker horizon due to its relative resistance and ease of definition in outcrop. The Limestone Ball Member derives its name from the boulder size, oval to round clasts, of grey limestone in a black shaley chert matrix. The boulders vary in size up to two metres long and one metre high. This unit is known to underlie the nickel mineralized bed from the northwest of the headwaters of Mawer Creek to the southeast extremity of the syncline east of Olfert Creek (Cathro, 1988). The northeast and southwest boundaries are unproven. Throughout the basin, as seen in outcrop and drill core, the thickness of the Limestone Ball Member is less than five metres.

The limestone unit is underlain by a dominantly siliceous thin shale, below which is a five to ten metre thick black limestone unit conformably overlying a calcareous shale member, up to 100 metres thick according to Cathro (1988). The thin siliceous shale is thought to be the Road River Formation. However, it may be confused with the (Canol) Lower Earn Formation, Green's units 10 and 13 respectively (1972). As mentioned by R.J. Cathro (1988), the clarification of this could be done by using diagnostic graptolite fossils used to separate the older Silurian to Middle Devonian from the younger Middle Devonian to Mississippian shale.

This stratigraphic sequence is similar to that of the marginal platform deposits of Morganti (1981). The figure below shows the generalized stratigraphic columns of the Nick location and the typical Platform Marginal deposit of Morganti (1981).

PLATFORM MARGINAL DEPOSITS

NICK DEPOSIT



NOT TO SCALE

D. PARRY

fig. 4

Structure

The nickel mineralization is located between the siliceous black chert and the limestone Ball Member in the limbs and trough of a gently folded synform. The direction of thickening of the vaesite horizon has not been determined from structural mapping. However, it has been determined that the nickel horizon either pinches out or is truncated by faulting between the drill holes situated in Olfert Creek and the single hole in Mawer Creek, 1600 metres to the northwest. It is also likely that the vaesite horizon exists in the outlier to the north of the main synform (see fig. 2 & 3).

The texture of the vaesite horizon is described by Cominco (1981) and Cathro (1988) as having been formed through soft sediment deformation. According to Davis, (1984) "*Wet sediments may slump, flow, fold or fault*

where inadequately supported or, where triggered to do so by earthquake shocks. Resulting structures are confined to a very narrow stratigraphic interval overlain by strata that are not themselves internally deformed..".

This explanation parallels the textural relationships found in the chert, vaesite and limestone horizons. The clasts of the Limestone Ball Member are noted to lie with bedding slightly oblique to the underlying shales. The formation of the limestone balls have been interpreted as a result of dissolution of calcite along joints (Cathro, 1988). The orientation of the limestone balls may well be due to rolling, or sliding down a paleoslope. This may provide evidence to show that the environment for deposition was a slope and not a flat lying third order basin (Ross, personal communication, 1989). The fact that the sulfide horizon shows these slumping textures, may provide an impetus to study the structure and stratigraphy to determine the direction of movement of the vaesite layer and Limestone Ball Member and thus a direction of possible origin. As discussed in Morganti (1981), the importance of the study of these slump and flow structures preserved in the rock record may help define the source and possible origin of structurally thickened horizons.

Host Rock Mineralogy

The host rock mineralogy is described in Appendix A, pages 1-27. In discussing the host rock mineralogy, the rock types will be grouped according to lithology. These types are: (1) fine grained clastics, (2) siliceous shales and carbonaceous to phosphatic cherts, and (3) micritic limestones. The order of these is equivalent to descending through the stratigraphic column (see fig. 4, page 5).

The fine grained clastics are at the top of the sequence observed at the Nick Property. They consist of non-siliceous argillites with a few thin beds of coarser, fine grained chert pebble sandstone and conglomerates. The individual beds are relatively homogeneous. The sequence generally coarsens upwards with argillites becoming less siliceous and resistive to weathering, and sandstones and fine pebble conglomerates becoming more numerous. Sample B is representative of the quartz, chert sandstone within this unit. (see page 1 of Appendix A) This sample is a detrital quartz-chert fine grained sandstone. The quartz grains are subrounded to rounded indicating a relatively distant source. The chert clasts are angular to subangular suggesting a relatively proximal source. Detrital magnetite was deposited with the quartz and chert, and formed detrital grains, and where oxidized formed the hematitic cement. The rock has undergone burial metamorphism which is responsible for recrystallized quartz infilling joints forming veinlets. The goethite and limonite which coats some of the grains and is pore filling, is likely an effect of diagenesis and supergene weathering, respectively. This unit is comprised of clastics and is named the Upper Earn Formation.

The Lower Earn Formation is composed of siliceous shales, carbonaceous and phosphatic cherts. The siliceous silvery black shales are, in fact, slaty and they form the upper portion of the Lower Earn Formation. The bedding is defined by laminations of varying concentrations of carbonaceous material. The mineralogy of the samples A-1 and C is dominantly illite with various amounts of silica in the form of chert, microcrystalline quartz and chalcedonic quartz flame structures. The slates were likely derived from colloidal clays which were regionally metamorphosed at low grade and imparted with an axial planar cleavage.

PLATE: 1

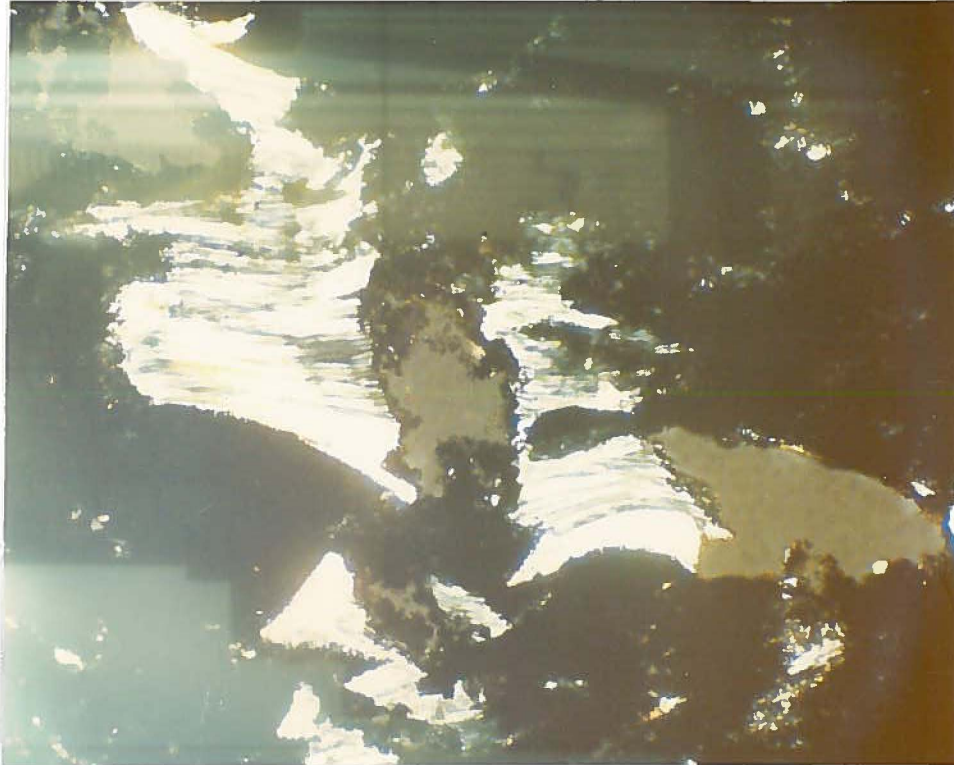


← 1.25 mm →

Transmitted Plane Polarized Light (TPPL)

Silvery black shale, Lower Earn Formation. (Sample A-1)
Showing the conduits of carbonaceous material in the illite rich shale.

PLATE: 2

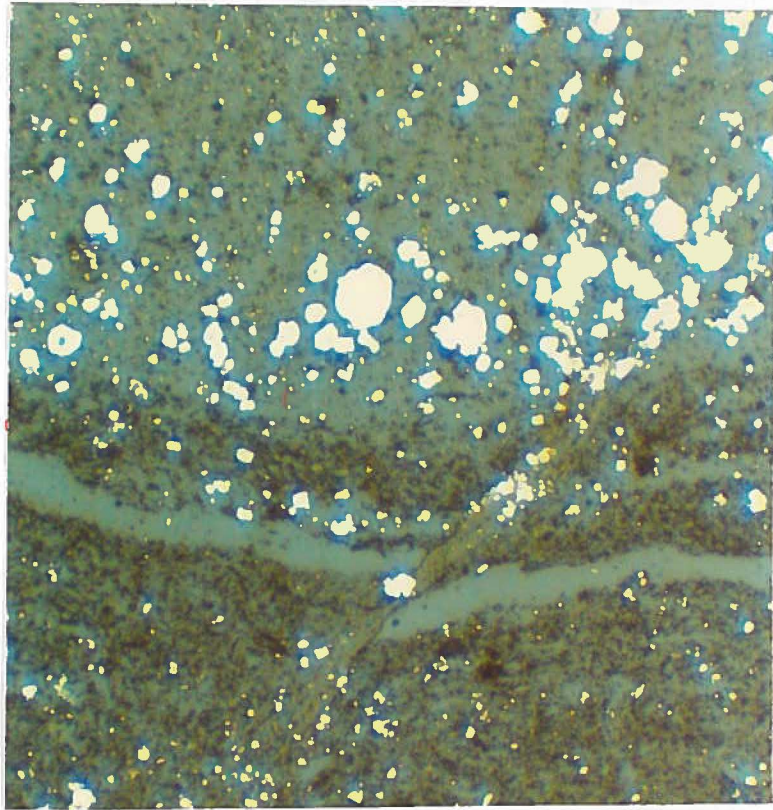


— 90 μm —

Transmitted Cross Polarized Light (TXPL)

Pyrite (grey) which is semi-homogenized and from which have grown siliceous flame structures, or "pressure shadows" of chalcedonic quartz white in black chert. (Sample 9049)

PLATE: 3



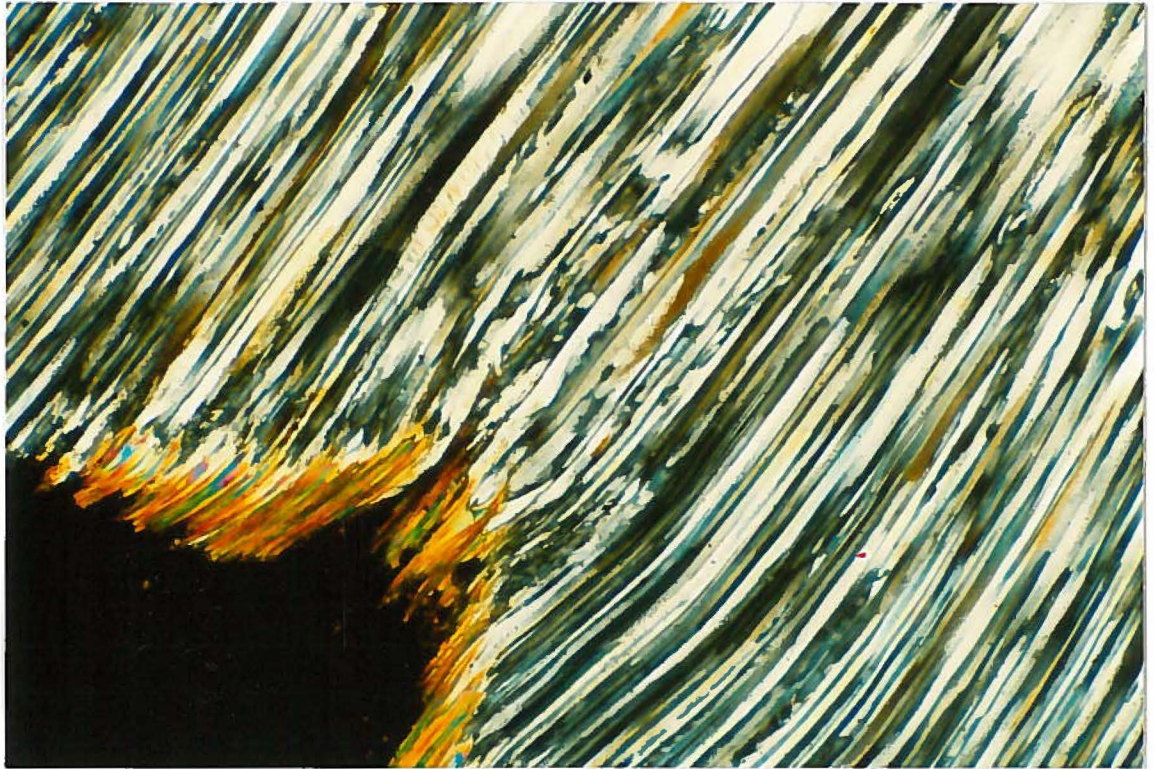
1 mm
Reflected Plane Polarized Light (RPPL)

The photomicrograph shows the base of a pyrite (white) layer. The band becomes more densely packed with semi-homogenized pyrite towards the top. The band of (grey) recrystallized quartz at centre shows the slight soft sediment deformation. Plane polarized reflected light (polished thin section). (Sample 9048)

This axial planar cleavage is a conduit for the transport of carbonaceous material within the slate (see plate 1). Descriptions of these samples are in Appendix A, pages 3-7.

The cherts form the lower portion of the Lower Earn Formation. They are represented by samples 9048, 9049, R-1, R-3 and R-6 in Appendix A, pages 8-16. These chert samples are all very uniform in appearance. They typically are finely laminated with pyrite in the form of framboids. This framboidal nature is an indication that these sulfides formed during diagenesis (Udubasa, 1984, Chauhan, 1984 and Nair and Ray, 1984). The pyrite framboids have altered in some cases to subhedral pyrite cubes. The samples are enriched in pyrite to approximately five percent, with varying amounts of chalcedonic quartz as flame structures of up to four percent (see plate 2). The presence of the wider bands of pyrite, such as observed in sample 9048, are common throughout the chert units in drill core (see Appendix C) and is indicative of a greater source of iron and sulfur. The sample 9048, which contains the band of pyrite, has also undergone a minor amount of deformation. The sample observed in plate 3 shows a minor amount of soft sediment deformation which may be responsible for the chalcedonic fibres in sample A. The pyrite in this band has (20 P.P.B.) trace platinum group elements, background arsenic levels and only slightly higher than normal estimated lead, zinc, iron and gold concentrations (Cathro, 1988, Assay Data). The source for the pyrite bands in the chert is unlikely to be the same as that which produces the vaesite-pyrite horizon at the base of the chert unit. The arsenic path finder element for the vaesite horizon is typically much higher than was observed in the pyrite bands in the chert unit.

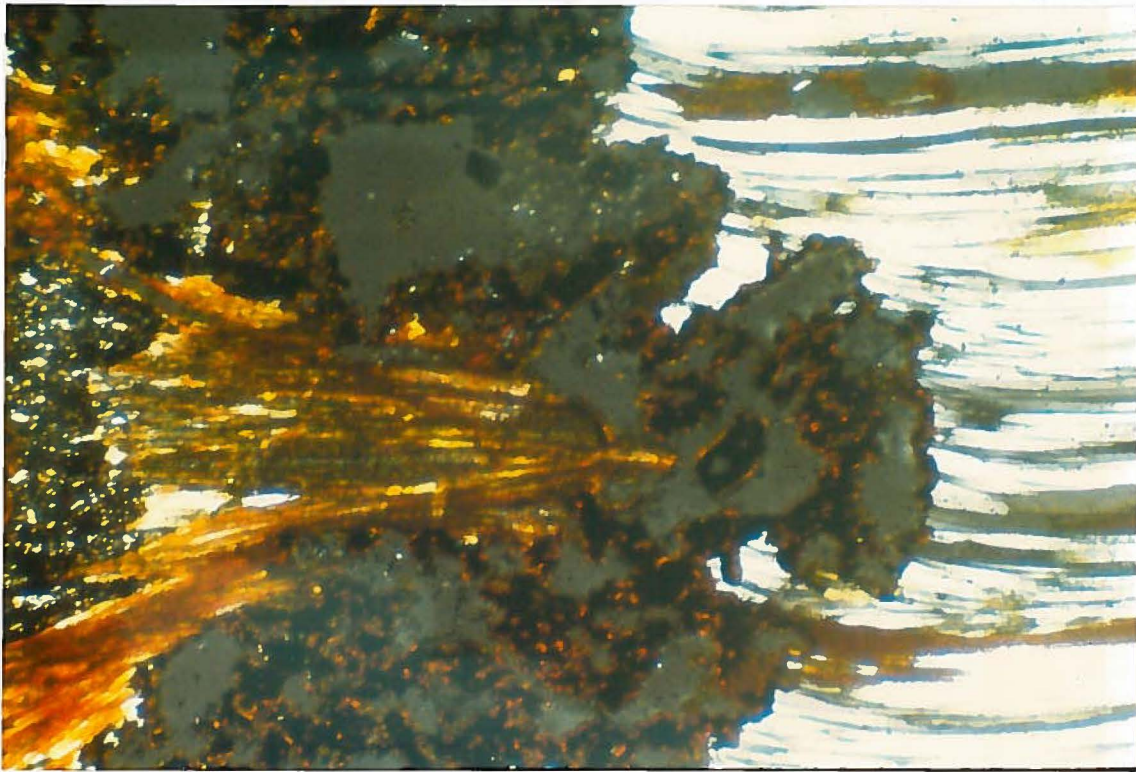
PLATE: 4



1.77 mm
Transmitted Cross Polarized Light (TXPL)

This photomicrograph shows chert host (black) with chalcedonic quartz (white-grey) fibres which have grown towards the direction of extension in this sample. (Sample A)

PLATE: 5



1.77 mm

Transmitted Cross Polarized Light (TXPL)

This photomicrograph shows the chalcedonic fibres (white to grey) on the right growing from the side of the cavity which is composed of limonite (reddish brown) and chert speckled yellow and black far left. The dark grey-green color represents vesicles in the sample probably due to weathering. (Sample A)

Within the cherty portion of the Lower Earn Formation are siliceous fibres formed in fractures, developed parallel to bedding. These structures take the form of sheared veinlets. These structures are found in association with both the chert and/or pyrite. They are commonly composed of quartz, hematite and limonite (see plates 4 and 5). Usually these fibrous masses form at the top of a pyrite band, however, they do form between chert beds which show slight soft sediment deformation (see Appendix A, pages 17-18). The structures are the result of soft sediment deformation and produce the fine white bands of siliceous material which range in size from submicroscopic to 18 cm² in area and 2 cm in depth.

Within the chert unit are minor thin beds of limestone with fine carbonaceous lamellae and pyritic euhedral crystals. The limestone units commonly have recrystallized sparry calcite in veinlets along fractures, probably the joints in the limestone (see Appendix A, pages 19-20).

The Limestone Ball Member which underlies the vaesite horizon, is composed of boulders of micritic limestone. This unit has recrystallized sparry calcite forming veinlets along probable joints in the limestone (see Appendix A, pages 19-20) and is underlain by a thin black chert and a thick calcareous shale.

A breccia sample was found which had penetrated upwards through the stratigraphic column, and may be of interest in the study of later deformation and hydrothermal fluid movement. It is unique in that it is the only hydrothermally altered sample taken from the property which has been analyzed in this study. The sample, R-5, found at 80.74 m depth in drill hole 88-4 is comprised of quartz, calcite, epidote, carbonaceous shale fragments and chert fragments (see plates 6 & 7). This sample is described in detail on pages 23-25 in Appendix A. It was taken from the drill core

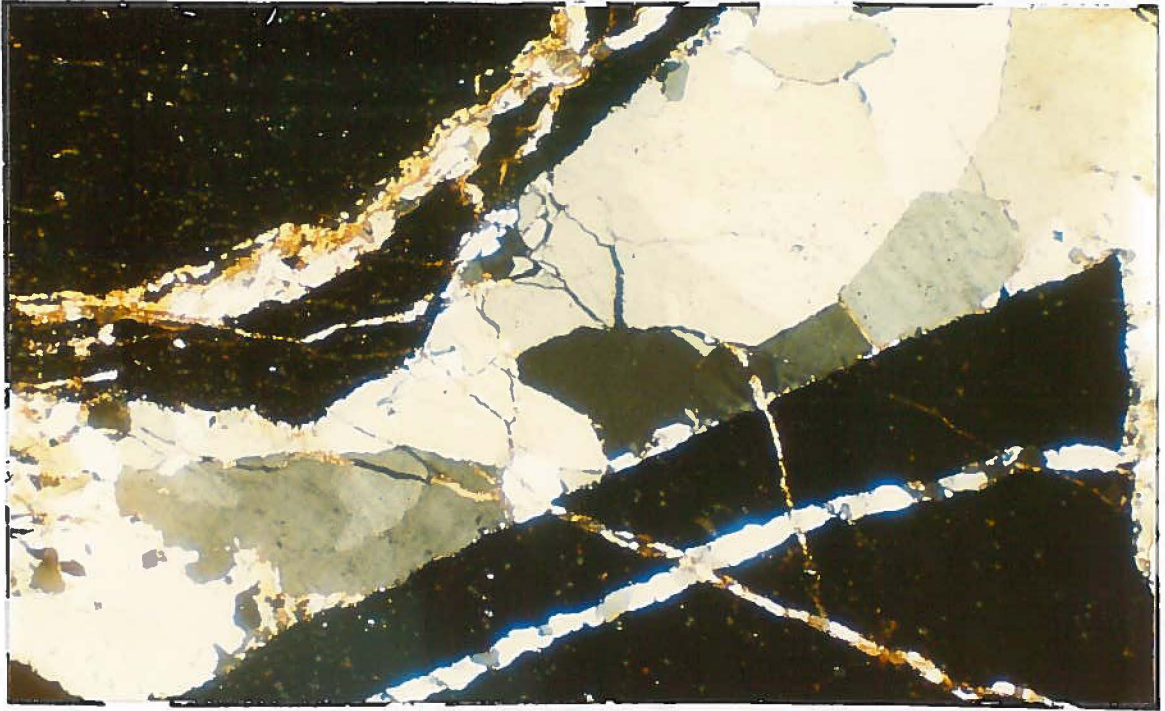
approximately 60 metres above the Limestone Ball Member. The vaesite horizon directly overlies the Limestone Ball Member to the southwest and, therefore, this latent breccia sample cannot be linked to the formation of the vaesite. However, fluid inclusions from this breccia sample may be of use in determining the nature of the hydrothermal fluids which circulated through this sequence after deposition. This small breccia pipe contained two-phase fluid inclusions of primary (negative crystals), pseudosecondary and secondary nature. Several of the inclusions contained H₂O and CO₂. At room temperature the CO₂ bubble was very active (see plates 8 & 9). Many of the inclusions of the secondary type were formed from the necking down of elongate inclusions. Fluid inclusion studies on an insufficient population of inclusions is of little scientific value because of the existence of error introduced by only one sample.

Mineralization

The mineralization occurs in a thin bed of stratiform sulfides between the overlying black cherts and underlying Limestone Ball Member. The mineralogy of the vaesite (NiS₂) horizon is very complex. The minerals present are pyrite, a variety of pyrite called melnikovite (Harris, J.F., 1988), vaesite, sphalerite, various forms of chert and bitumen. The association of minerals and textures in this stratiform deposit are unusual. The study of the mineralogy was conducted using petrographic reflected and transmitted light microscopy x-ray diffraction with additional identification procured from a report by J.F. Harris (1988) in which he analyzed several of the finer grained minerals by Scanning Electron Microscope.

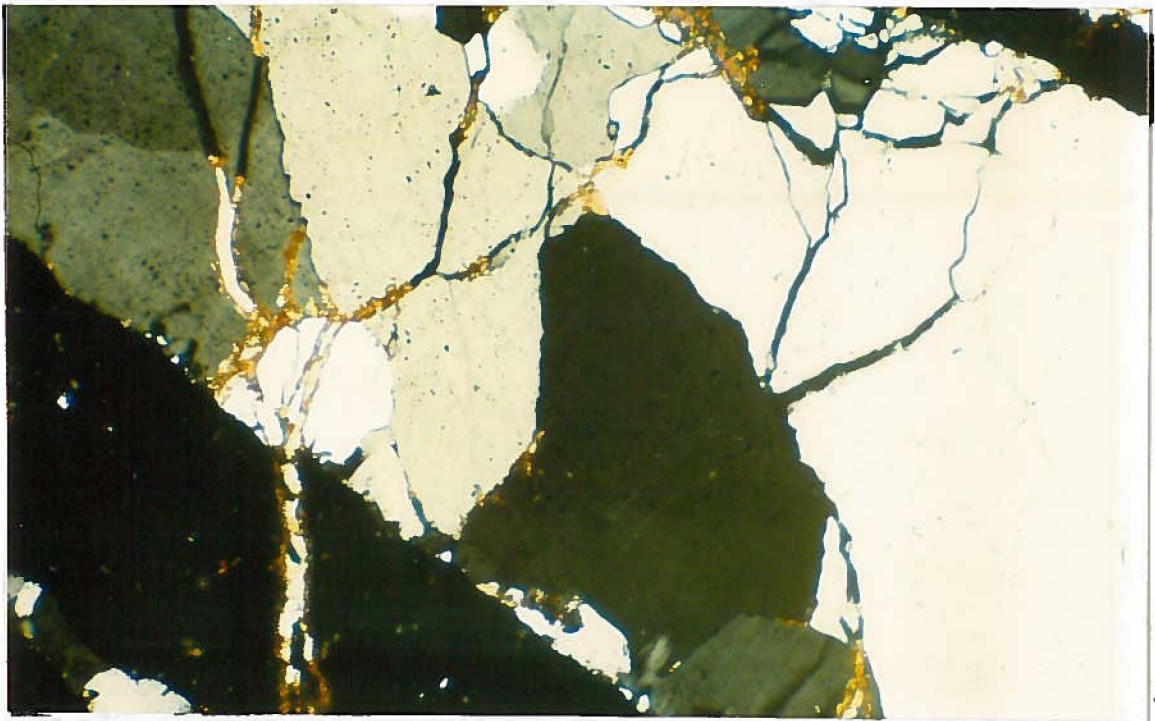
The Nick mineralization is present in disturbed, rhythmically deposited lamellae which form individual layers up to 3 mm thickness. The

PLATE: 6



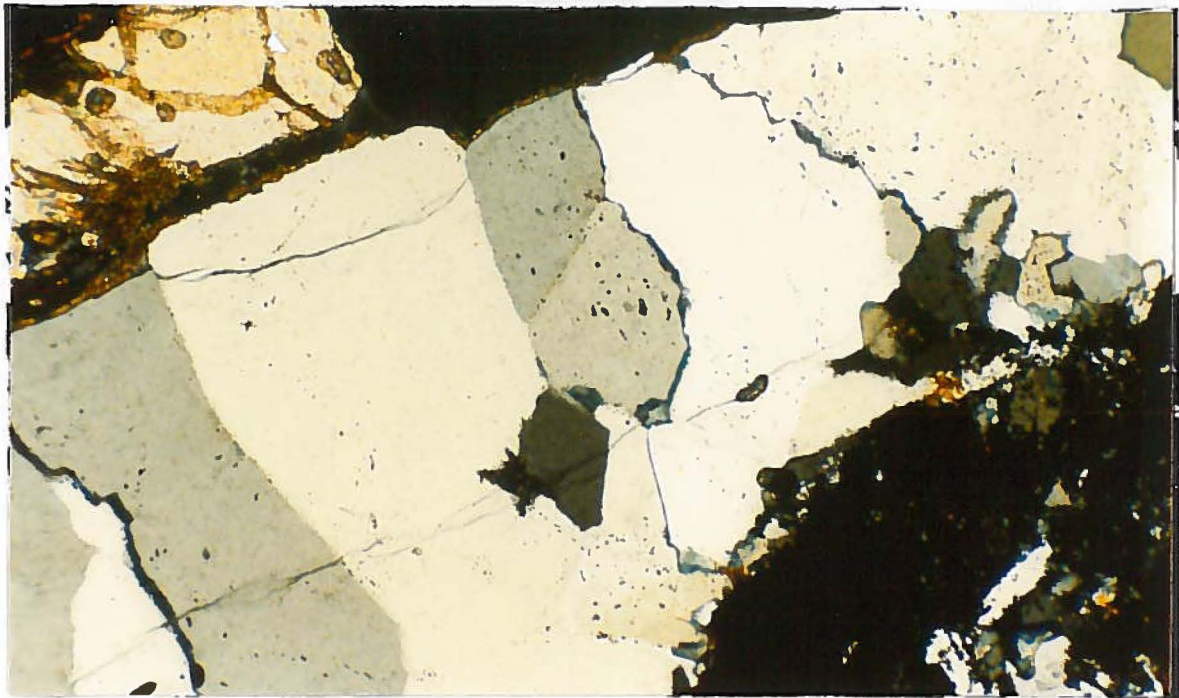
— 2.48 mm —
Transmitted Cross Polarized Light (TXPL)
Cross cutting relations in the quartz breccia sample (R-5).

PLATE: 7



— 900 μ m —
Transmitted Cross Polarized Light (TXPL)
Higher magnification of above. (R5, DDH88-4, 80.74m)

PLATE: 8

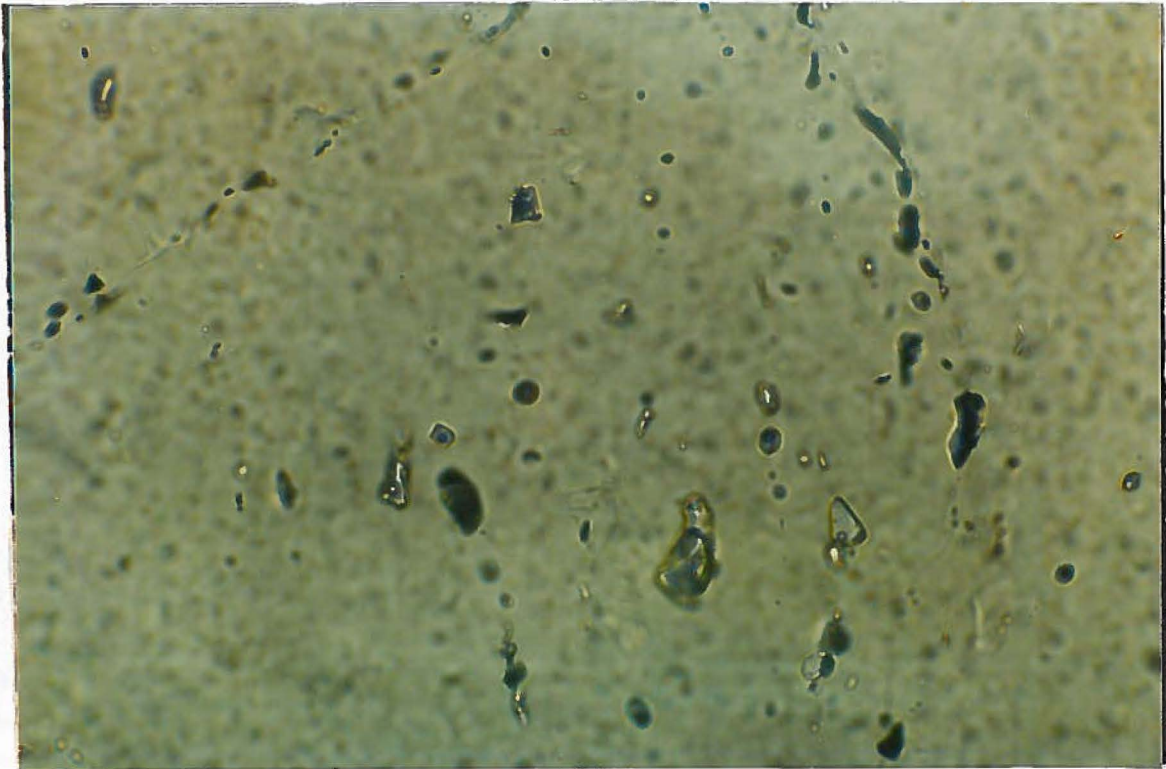


800 μm

Transmitted Cross Polarized Light (TXPL)

Multiple fields of fluid inclusions in quartz-epidote breccia sample.

PLATE: 9



160 μm

Transmitted Cross Polarized Light (TXPL)

A view of several primary double phase fluid inclusions in quartz from the sample above. (Sample R5, DDH 88-4, 80.74m)

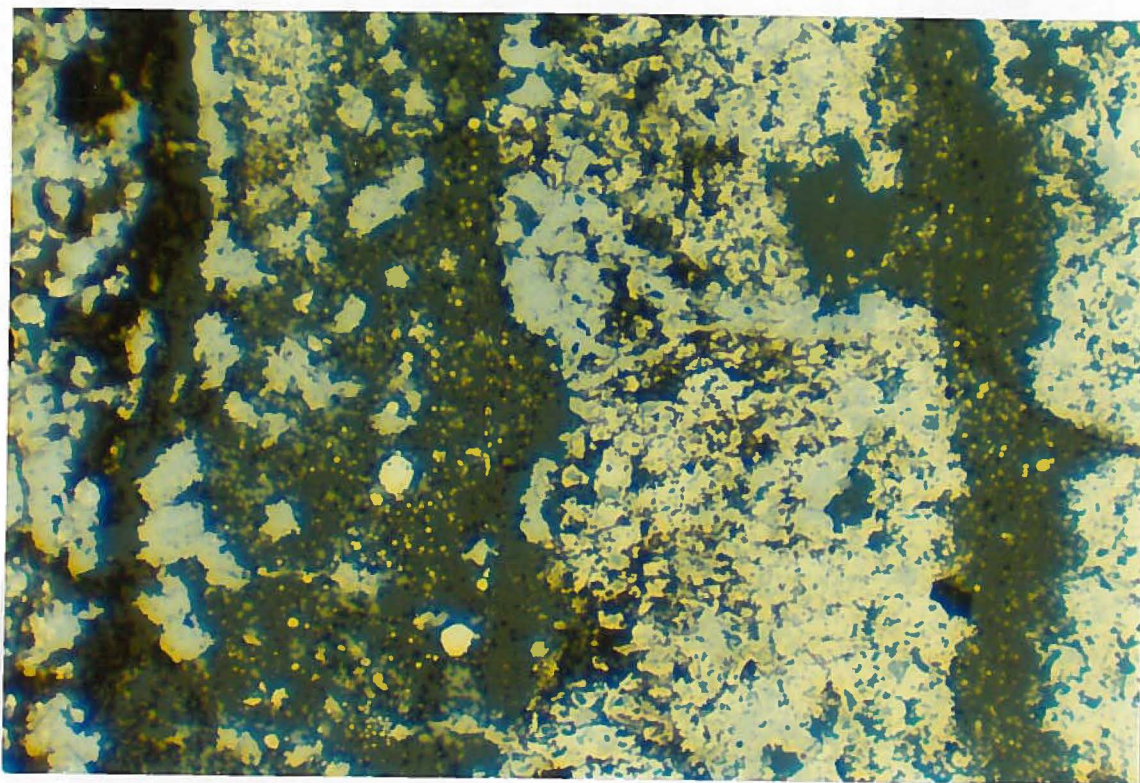
nickel mineralized horizon varies from 2 cm to 10 cm in thickness. Each lamellae is unique. No layered sequence of mineralization is apparent, although there exists an affinity for the deposition of sphalerite after that of pyrite, melnikovite and vaesite. This apparent sequence of mineralization may be due to the solubility of zinc, iron and nickel in exhaled thermal brines, or the late incorporation of base metals. The mineralized layers are typically interlaminated with carbonaceous and phosphatic chert and nodules of probably biogenetic chert. The appearance of the laminae are expressed in plate 10.

The deposition and formation of the mineralized zone is thought to occur through a process of diagenetic recrystallization from colloidal sulfides (Chauhan, 1984). The pyrite which occurs in these layers as globular bands may be intergrown or rimmed by arsenic (plates 11 & 12) similar to the spheroidal pyrite grains of two Romanian examples described by Udubasa, (1984). Furthermore, Udubasa outlined a genetic link between framboidal pyrite and an accessibility to carbonate in the host rocks.

The bands which formed in the cherts and shales above the mineralized zones contain much higher concentrations of framboidal pyrite. This may be a result of increased carbonate in both the rocks and marine waters. The lack of framboidal textures in the mineralized strata may be due to hydroplastic dewatering of the pyrite which would cause a homogenization of the pyrite (Chauhan, 1984). This process would also produce the siliceous flame structures in the pyrite of the upper shales and cherts, and the fibrous siliceous material within the vaesite horizon.

The chalcedonic siliceous fibres which are found in association with the pyrite layers have also been reported in the Ordovician greywacke from southwest Scotland and in Precambrian cherts in the algal zone of the

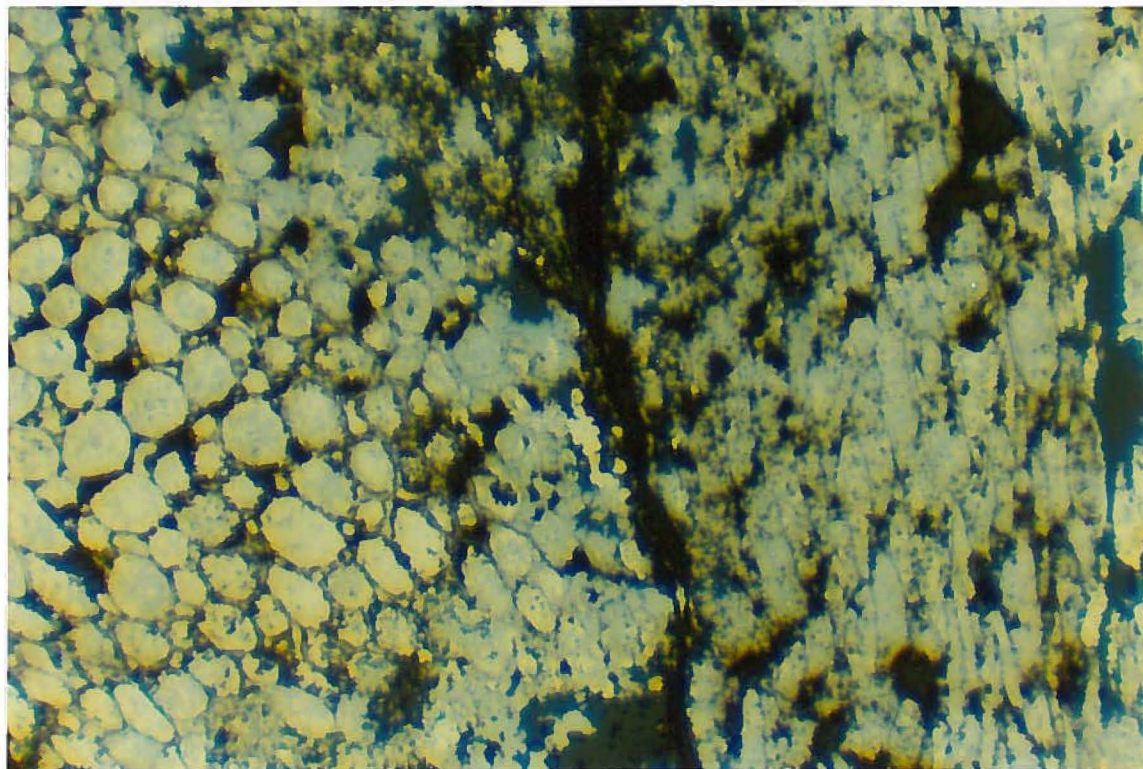
PLATE: 10



2.0 mm
Reflected Plane Polarized Light (RPPL)

Typical field showing the intense variation in textural relations and mineralogy. The wide band right of centre is composed of pyrite (yellow), melnikovite (brownish yellow, interior of band), vaesite (pale grey) and sphalerite (dark green-grey). This laminar band is suspended in a framboidal pyrite rich chert. The interlamination and fine intergrowths are typical of the vaesite horizon. (Sample Nick outcrop A-1)

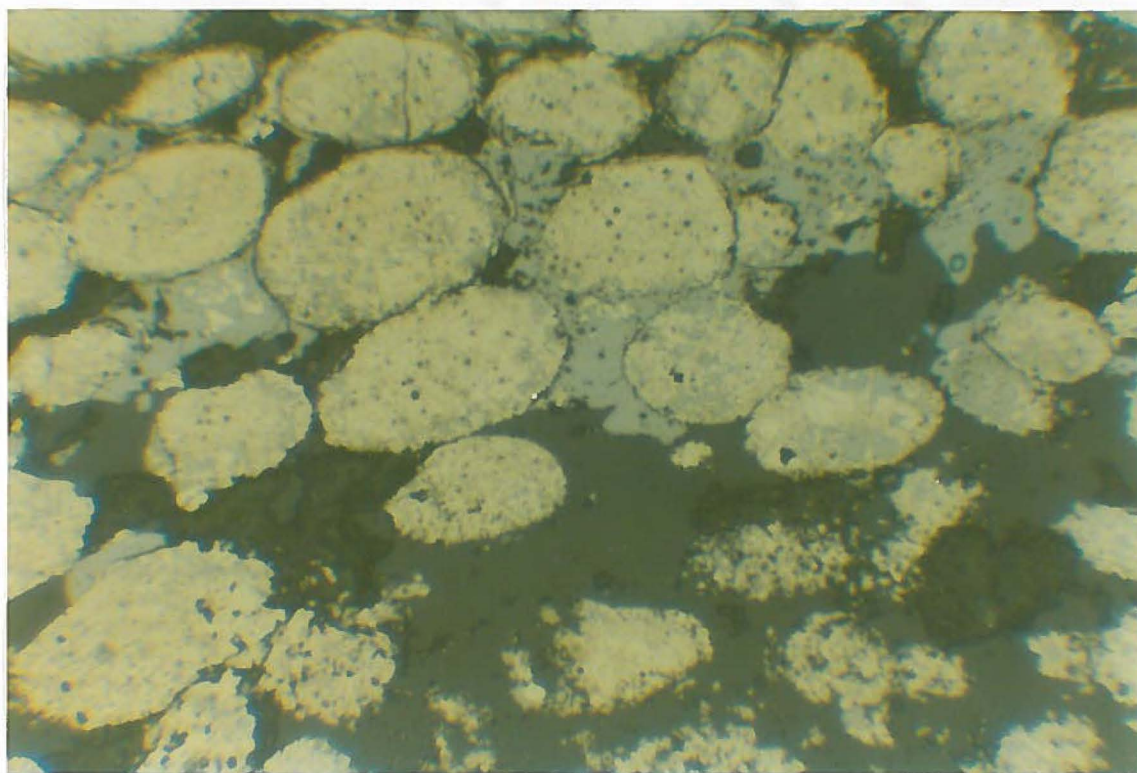
PLATE: 11



—500 μ m—
Reflected Plane Polarized Light (RPPL)

Similar to plate 12, the homogenized framboidal pyrite (creamy white) and vaesite (blue grey) has been compressed and attenuated in the band to the right by soft sediment deformation. The band to the left shows a little deformation near the interface between the bands. (Sample Nick outcrop A-1)

PLATE: 12

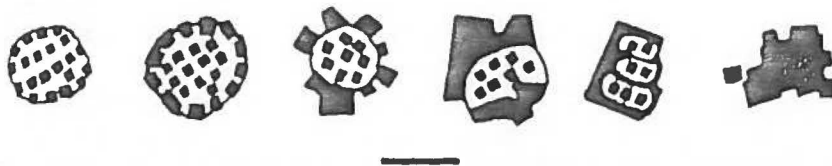


— 250 μm —
Reflected Plane Polarized Light (RPPL)

Spheroidal masses of homogenized pyrite (creamy white) and melnikovite (brownish cream) with intergrown vaesite (blue grey). Carbonaceous chert (dark green) and vaesite (blue grey) with (yellow) pyrite and intergrown pale green colored sphalerite (1-10 μm) in the interstitial vaesite. Note the rim on the spheroids which may be due to arsenic. (Sample Nick outcrop A-1)

Gunflint, and Biwabik Formations (Kelling, 1971, Tyler and Barghoom, 1963 in: Nair and Ray, 1984). "Such zones or rims are produced either by contraction as a result of conversion of the amorphous iron sulfide into crystalline pyrite and their filling by diagenetic silica/carbonate (or pyrite/vaesite) or, by the development of stress shadows during epimetamorphism and the filling of spaces" (Nair and Ray, 1984).

Chauhan (1984), suggested two processes which produce homogenized pyrite. The first *sammelrekristallisation* which results from the coalescence of the microcrysts of pyrite, expelling the trapped material, producing a welded texture. The second process introduces a later pyritic fluid which produces a second generation of pyrite infilling in the voids and replacing trapped material, similar to figure 5, from Udubasa, (1984)



Gradual transformation of framboidal pyrite to euhedral pyrite grains in a hydrothermally affected black shale, Rotunda-Strimbu area, Baia Mare region, N. Romania. Scale bar is 10 μm

G. Udubaşa

The *sammelrekristallisation* method of homogenization is in keeping with the homogenized and semi-homogenized pyrite of the cherts and shales of the overlying strata, and to a lesser extent may play a role in the homogenization of the vaesite horizon. The infilling of diagenetically formed framboids by a second generation of nickel, zinc, iron and sulfur is conceivable as a process for the formation of intergrown vaesite, sphalerite, melnikovite in the banded layers and at the margins of the banded layers.

The diagenetic events which produce the mineralization in the vaesite layer are a variation on that proposed by Chauhan, (1984). (See Table 1.)

Table 1. Generalized diagenetic events during crystallization of metallic sulfides

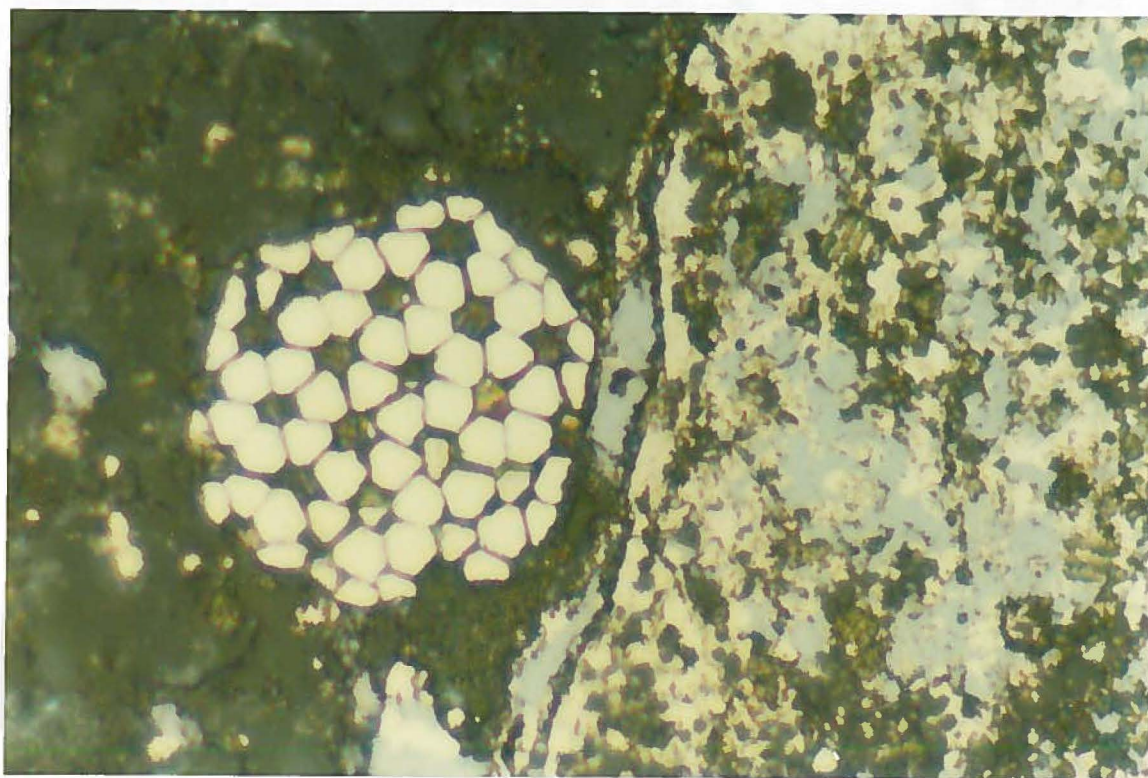
	1st diagenetic stage (depositional)	2nd diagenetic stage (shallow burial)	3rd diagenetic stage (cementation – premeta- morphic)
Pyrite	Tiny pyrite grains and distinct framboids appear. Tiny grains of pyrite assemble in cubic outlines	Cementing pyrite appears; <i>Sammelrekristallisation</i> takes place	Either grains or framboids cemented together or with one another to form bigger crystals
Sphalerite	–	First appears in small amounts	Mainly crystallizes during this period
Galena	–	–	Makes appearance during early part of this stage and mainly crystallizes through this stage

D.S. Chauhan

The Nick Property mineralization varies with respect to the introduction of nickel in the same stage as that of pyrite *sammelrekristallisation* (stage 2). An example of the effects of mineralization of this type is illustrated in plate 13. Samples from this horizon do not show any intergrowths of galena, and assays show very low concentrations of lead. Therefore, this horizon has not undergone the third stage of diagenetic crystallization.

The pyrite vaesite horizon is composed of the diagenetically semi-homogenized pyrite intraclasts (Nair and Ray, 1984). These intraclasts maintain the initial structure and mineralogy of the lamellae prior to the penetrative deformation. The maintenance of these intraclasts in the deformed unit provide a relative timing for the deformation. The deformation may have ended the second stage of diagenesis or occurred afterwards (see table 1, Chauhan, 1984 and plate 14).

PLATE: 13

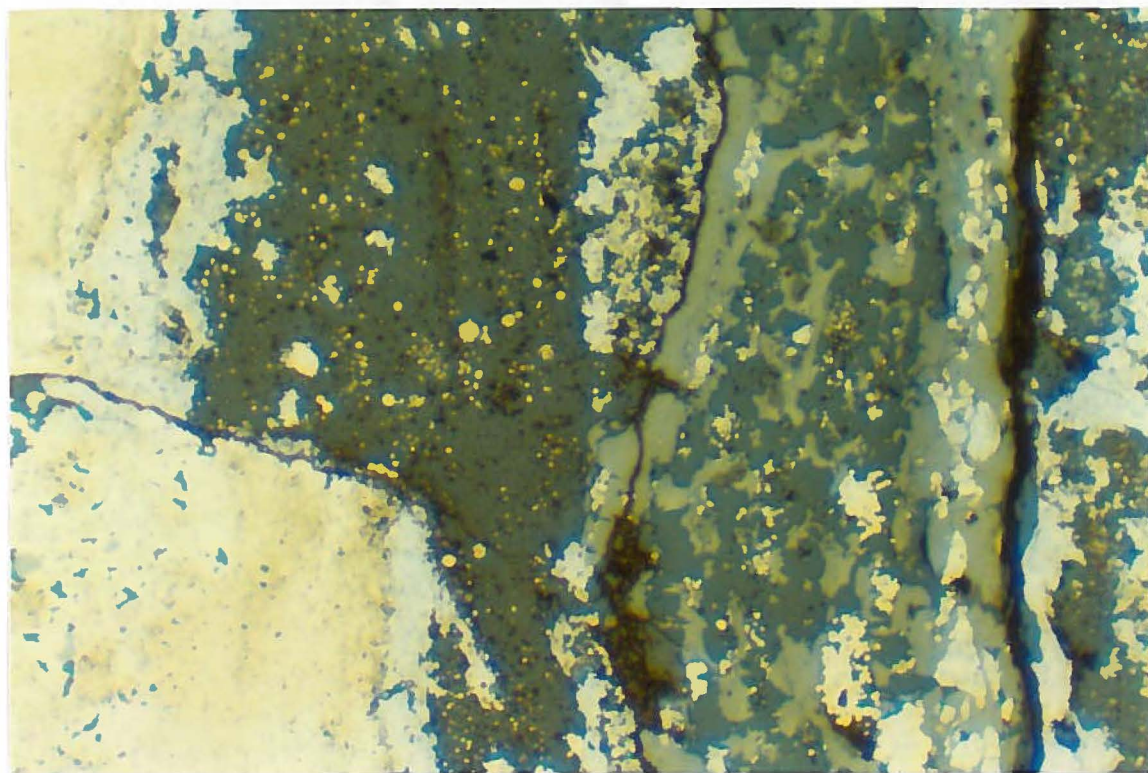


— 60 μ m —

Reflected Plane Polarized Light (RPPL)

Pyritohedron (pale creamy yellow) formed from smaller pyritohedron contained within the chert band (dark grey-green). The chert is interspersed with pyrite framboids and amalgamations of vaesite, pyrite, melnikovite and minor sphalerite. The pyrite band to the right is partially homogenized (near stage 2) and is admixed with sphalerite (blue-grey color). The dark green-grey material is carbonaceous phosphatic chert which was trapped in the original pyrite framboids. Several pyrite framboids are still visible in the band. (Sample Nick outcrop A-1)

PLATE: 14



————— 200 μ m —————
Reflected Plane Polarized Light (RPPL)

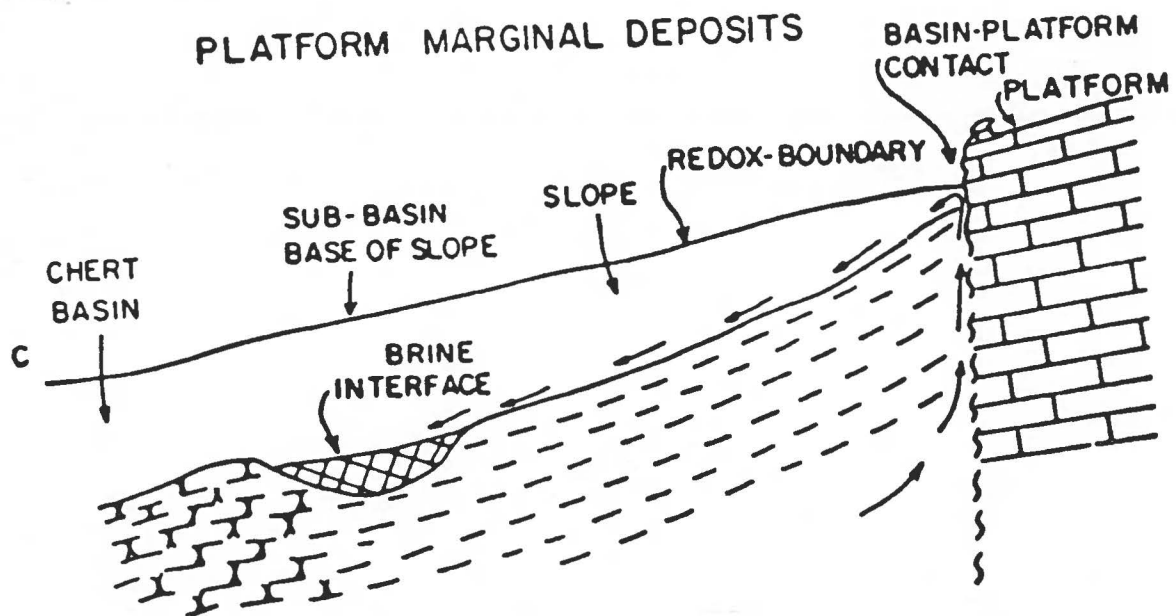
The photomicrograph above shows the intraclasts of pyrite (yellow) and vaesite (blue grey) deformed with framboidal chert conforming to the intraclasts (dark grey with yellow flecks and blebs). Intergrown sphalerite and chert show a later development yet, with near straight layer boundaries. The brown bands which are positioned on either side of the sphalerite chert lamellae may be bitumen. (Sample Nick outcrop A-1)

The deformation of the finely laminar sulfide layers at the nodular cherty boundary is complex. The sulfides show deformation around the chert nodules, but blebs of sulfides contained within the nodule are underformed (see plates 15 & 16).

DEPOSITIONAL MODEL:

As alluded to previously, the model for the transport and deposition of metals closely resembles that of sedimentary exhalative deposits. Sedex deposits of the Selwyn Basin, and other regions, are differentiated by the source of the fluid, the method of transport and depositional location. Morganti (1981) suggested three tectonstratigraphic environments; epicratonic deposits, flysch deposits and platform marginal deposits.

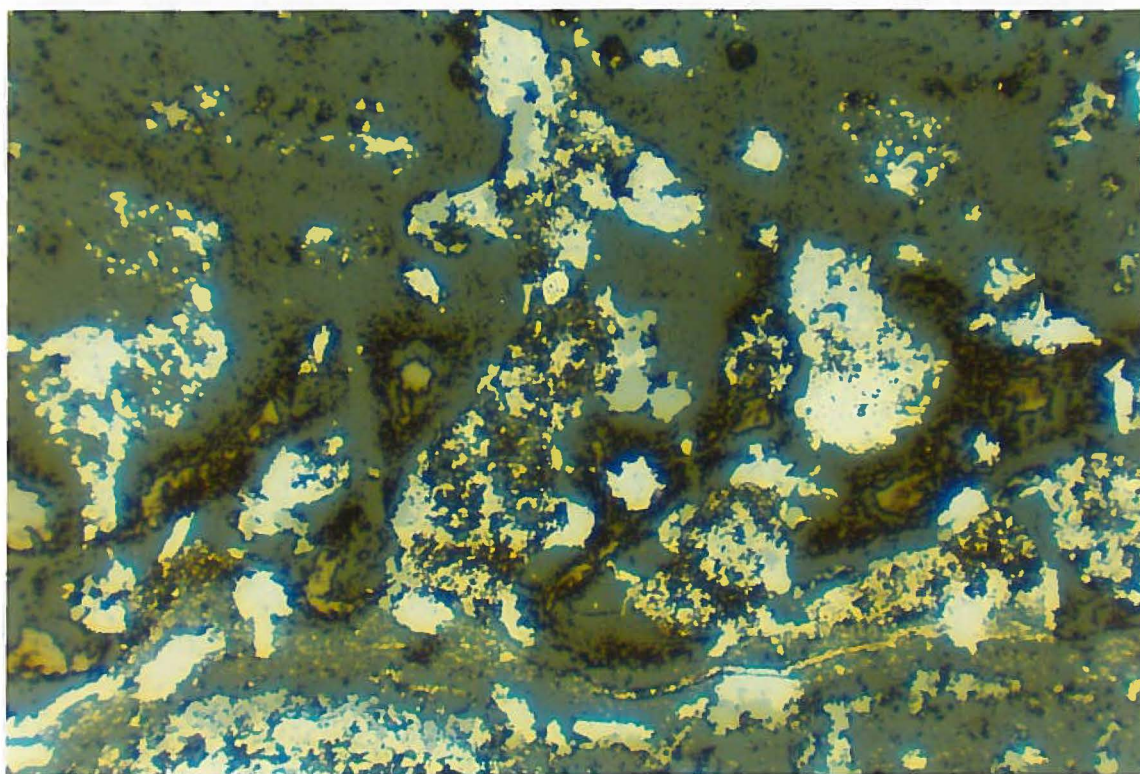
Using the Morganti model for platform marginal deposits, as seen in fig. 6, the factors which are present at the Nick Property show a strong resemblance to this model.



(Morganti, 1977).

Fig. 6

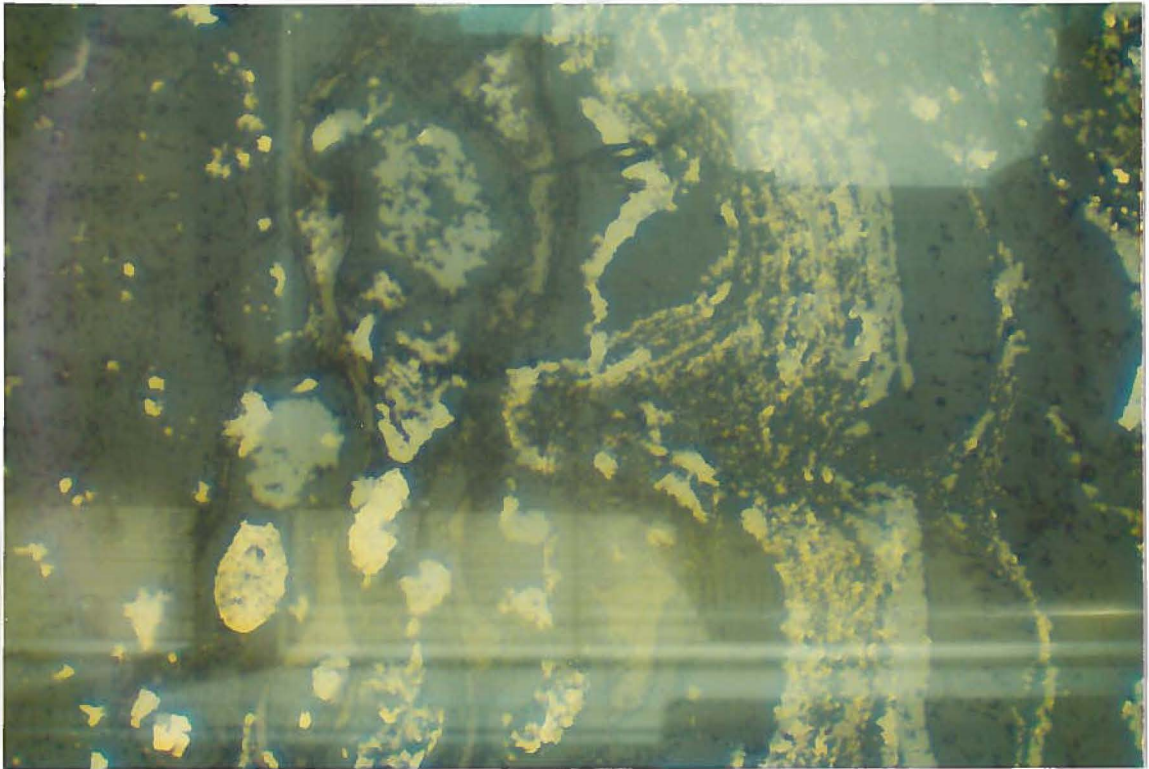
PLATE: 15



500 μm
Reflected Plane Polarized Light (RPPL)

The photomicrograph above shows a chert nodule, sulfide lamellae boundary. The sulfides: vaesite (blue grey), pyrite (yellow), melnikovite (brownish yellow) and sphalerite (light greenish grey) are intergrown into chert (dark green-grey) nodule which contains blebs of bitumen (?) at the boundary (dark brown). Notice the compaction and conforming of pyrite to the chert nodule during soft sediment deformation and the 40 μm pyrite framboid in the nodule. (Sample Nick outcrop A-2)

PLATE: 16



— 80 μm —
Reflected Plane Polarized Light (RPPL)

The above photomicrograph was taken along the same boundary where the sulfide was deformed more readily by the confining pressure and the rigid chert nodule. (Sample Nick outcrop A-2)

The Nick mineralization is rich in nickel and platinum group element, with lesser amounts of molybdenum. This varies from the typical lead, zinc sulphide occurrences of the OP, ANNIV and XY deposits of the Western Selwyn Basin. The absence of the high barium concentrations at Nick compared to their presence at the Flysch - Epicontinental type deposits, attests to a dissimilarity in the metal source, transport and depositional system. The above sketch of the model proposed by Morganti (1977 in: Morganti, 1981) shows the pooling of sulphide rich brines at the lithological interface between shales and cherts. However, at the Nick Property, the mineralization occurs above the Limestone Ball Member and a thin layer of chert. The formation of this mineralization occurs below the redox boundary, probably in an euxinic region. The concentration of nickel and PGE's may have been encouraged by the activity of micro-organisms or bacteria in this environment.

The structural parameters such as the fault bound carbonate margin, paleoslope indicators and host rock lithology can be shown to parallel those of the Western Selwyn Basin deposits (see fig. 7).

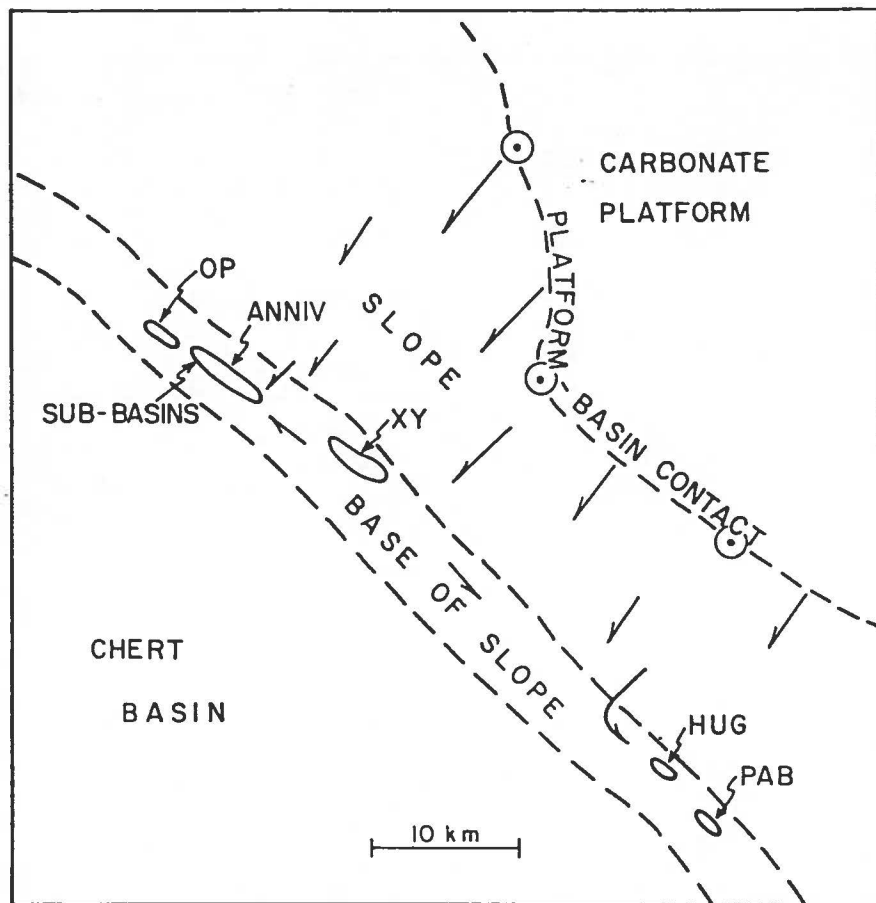


Figure 7 General model for the formation of the platform-marginal deposits showing the geometry of the platform-slope-base of slope and chert basin facies. Arrows show surface movement of ore fluids which migrate up to the sediment-water interface, down the basin slope and are trapped in the sub-basins at the base of the slope.

The framboidal and recrystallized forms of pyrite serve as good guides in the tracing of economically important base metals which form intergrowths with the pyrite (i.e. framboidal pyrite will contain little zinc and nickel, whereas recrystallized or overgrown pyrite will usually indicate high base metal concentrations). The determination of diagenetic pyrite helps in discovering of the probable time of base metal fixation and behavior of post depositional soft sediment deposition.

DISTRIBUTION OF METALS:

The assay results of samples from the vaesite layer are tabulated in Appendix B. These results provide information regarding the distribution of nickel, zinc, barium, platinum group elements and various other metals contained within this horizon. The vaesite layer maintains a respectable grade of approximately 4.4% nickel throughout the explored region. The other base metal values for concentrations vary, zinc ranges from approximately 770 to 10,000 grams/tonne, and values for lead concentration are appreciably lower. The molybdenum values range between 1454 and 3930 grams/tonne. The element arsenic consistently parallels the variation in the percentage of nickel and therefore is a useful tracer element for the nickel bearing zone. The precious metals; platinum, palladium, gold and silver also closely parallel the nickel concentration.

From the assay data and field relations it become apparent that the source for both the nickel and base metals is probably the same one that supplies the precious metals. The variations in the concentration may be caused by a number of reasons. Firstly, deposition may involve a micro-organism which is responsible for nickel precipitation which itself is concentrated in certain environments and secondly, the paleo-terrain may effect the concentration of the hydrothermal brine fluid, in channelways on slope and third order basins at the break-in-slope of the basinal margin.

THE COMPARISON OF THE NICK PROPERTY WITH OTHER SIMILAR EXAMPLES:

The Nick mineralization is unique although the components are similar to several different metal bearing regions world wide. In Southern China, the Lower Cambrian is an important ore genetic epoch. The black shales

series of this region contains phosphorite, stone coals and Ni-Mo polyelemental rich beds (Fan, 1983). At the Kasompi Mine in Zaire, vaesite and pyrite have been reportedly found in occurrences with dolomite (Cathro, 1988). Several Chinese locations exist according to Coveney and Chen (1988).

The Songlin occurrence, for example, contains a nodular sulfide lense 5 to 15 cm thick within a black shale which contains up to 4% nickel, 4% molybdenum, 2% zinc, 0.7 ppm Au, 50 ppm Ag, 0.3 ppm Pt and other platinum group elements. Vaesite is present with other nickel and molybdenum sulfides in microspheroids of organic matter in thin beds up to 5 cm thick (Fan, 1983). Several other localities exist which contain this similar nodular Ni-Mo sulfide occurrence.

At a location in the Hunan Province, Cambrian aged shales contain these Ni-Mo rich nodules and are associated with P.G.E.'s (Fan, 1983). The nodular textures exhibited together with other factors imply that submarine hydrothermal springs related to basement faulting may have been responsible for the metals (Chen, 1988 in: Coveney and Chen, 1988).

In the Lower Cambrian, in the Hunan Province, a series of strata resembles the stratigraphy and mineralization of the Nick Property. The Niutitang Formation in Tianmenshan is compared with the Nick Property stratigraphy (in descending order):

- a) Dark grey and black illite shale, intercalated with grey dolomitic lenticles. This shale is 35 to 40 metres thick, and is characterized by fine lamellae of pyrite.
- b) Black argillaceous siliceous rock (approximately 18 metres thick).
- c) Black dolomitic shale and silty shale 1.2 to 1.5 metres thick, containing fine discontinuous pyrite laminae of 2 to 3 mm thickness.

- d) Black shale with phosphatic concretions and Ni-Mo polyelement beds commonly the phosphatic concretions overlie the Ni-Mo horizon (2 to 5 cm thick).

The units (a-d) closely resemble the sequence at the Nick Property which is composed of a clastic fine grained shale horizon equivalent to (a), a siliceous black shale unit equivalent to (b), a black chert with intercalations of calcareous shale equivalent to unit (c), and the vaesite horizon which is equivalent to unit (d) (Fan, 1983).

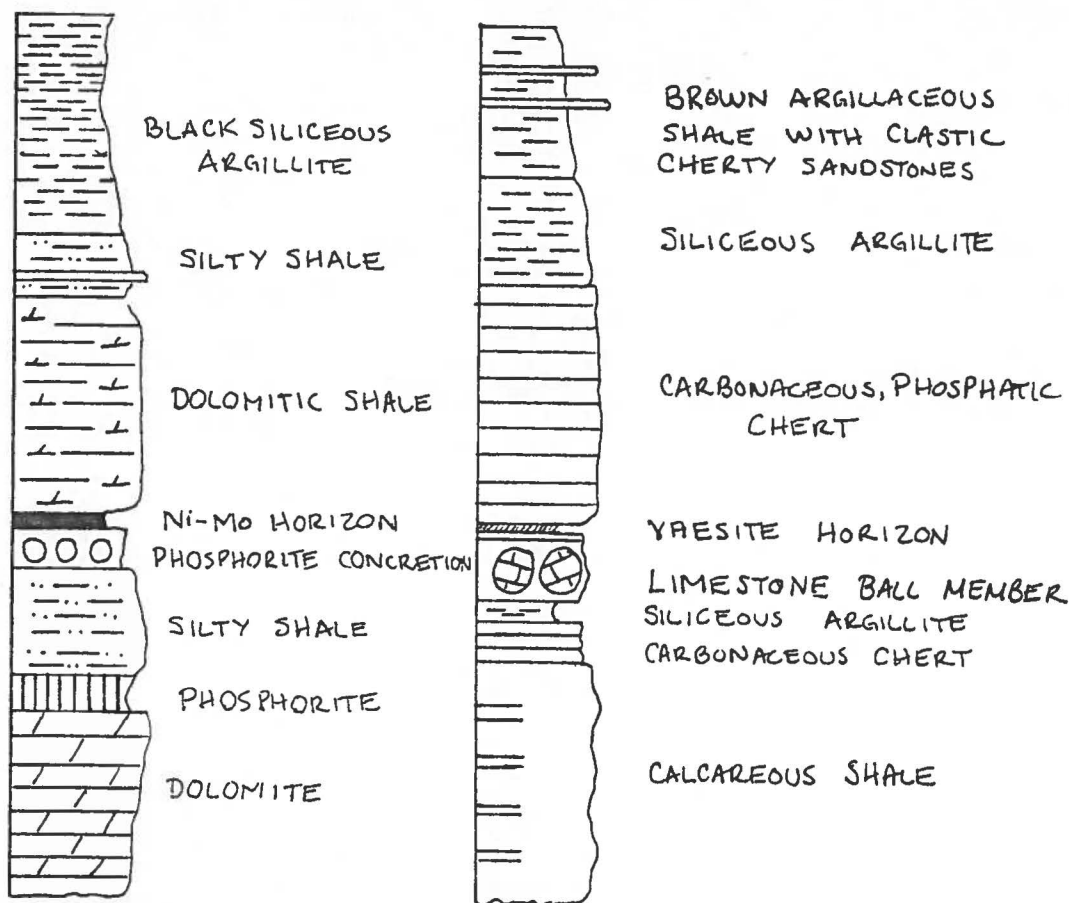
The Chinese sequence which contains the polymetallic beds has several striking similarities to the Nick Property:

- 1) The Lower Cambrian black shale series is located between two thick layers of argillaceous carbonates.
- 2) The Chinese example and the Nick Property both show generally shallowing upwards sequences.
- 3) Both host units are composed of dark shales, and siliceous black rocks with the mineralized zone at the base.
- 4) Both contain the nickel-molybdenum mineralization suites.

See figure below.

HUNAN DEPOSIT

NICK DEPOSIT



NOT TO SCALE

DISCUSSION:

The Nick Property may be characterized as a non-magmatic hydrothermally derived deposit type related to rifting and basin development. Kissen, (1988) suggests that a suite of five elements: nickel, cobalt, arsenic, bismuth and silver is related to shale hosted deposits, which contain in addition, Mo Tl, Se, Cd. According to Pereira (1977), Ni, Pt, Co and Cr metals are plate margin indicators, and can be linked to certain localities, for instance, Ni is associated with the Mekong fold belt sediments. According to Coveney and Chen, (1988) the Pennsylvanian age shales of the central U.S. may have similar ore grade enrichments associated with

basement faults associated with the Central North American Rift System and other structures. The Nick deposit, like many of the Chinese deposits, has a strong affinity for Pt and Ni with lesser affinities for Co and Cr. In Eisbacher (1985), conclusions were drawn which correlated the Southern Chinese, Eastern Australian and Western North American Precambrian strata. These regions were later rifted away from one another during the Early Paleozoic. On two of these three continents black shale related Ni, Mo, Pt type deposits have been discovered near the paleo-rift margins.

Kissen (1988), provided geochemical evidence for the transport of Ni and Co in hydrothermal systems. However, the Nick mineralization has a low Co:Ni ratio. Kissen (1988), conceived highly saline, oxidizing carnate fluids mobilizing Ni, Co, As, Bi, and Ag from black shales, sulfide rich metavolcanics, and other continentally derived strata in a rifting environment. The transport of this fluid through open structures near the basin margin mixing with diluting and cooling ground waters may reduce the Ag, Co, and Bi concentration in such a way that a more concentrated nickel, arsenic, molybdenum and P.G.E. rich brine would exhale. Localized factors such as carbonaceous sediments or, pyritic sulfides in fahlbands could be the location for deposition of these base and precious metals (Kissen, 1988).

CONCLUSION:

This study is important because of the understanding one can draw from it pertaining to the local geology and constraints on the ore forming system. The distribution of nickel and platinum group elements was shown to be enriched in a horizon which is characterized by the homogenization of pyrite with intergrowths of sphalerite and minor nodules of black chert. The

nature of the Nick deposit, strongly favors the platform marginal model for ore formation.

The nickel, molybdenum, P.G.E. metal association in sedimentary black shale environments is shown to occur at only one other location worldwide, in the Hunan Province, China. These metal associations may also be linked to rifting of the Proto-Pacific Ocean. This type of deposit should be given more attention in both the research and the exploration fields because of its unusual metal associations and correlation with rifting continental margins of the Proto-Pacific.

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APPENDIX A
PETROGRAPHIC DESCRIPTIONS

Petrographer: D. Parry

Date: January 23, 1989

Thin Section: B Hole: -

Depth: -

Rock Name: Quartz-Chert Sandstone

Location: Upper Earn, Fine Clastic Unit

Transparent Minerals:	%
<u>polycrystalline quartz</u>	<u>40%</u>
<u>quartz</u>	<u>~4%</u>
<u>chert</u>	<u>50%</u>
<u>hematite altered to goethite and limonite (iron stain and cement)</u>	<u>5%</u>
<u>sphene (titanite)</u>	<u><0.1%</u>

Texture:
detrital grains rounded to subrounded; pressure solution, silica suturing & silica cement as cements
recrystallized as veinlets up to 1 mm thick detrital angular - subangular (250-500 μm)
anhedral, subhedral and euhedral cubes found coating grains and filling pores only observed as nucleus of chert grains also occasionally found within chert

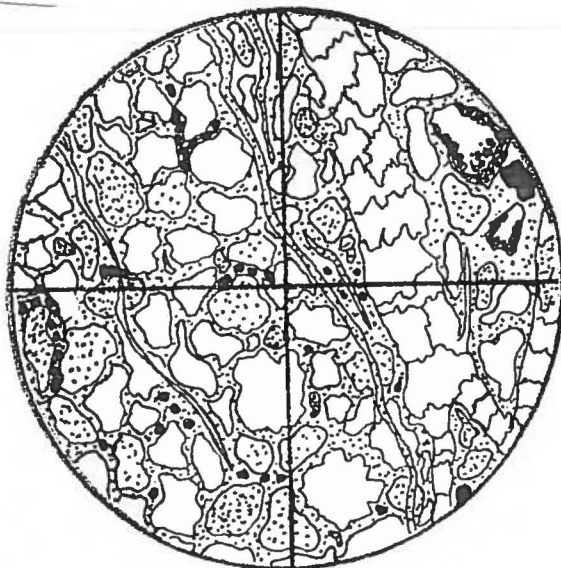
Opaque Minerals:	%
<u>hematite</u>	<u>1%</u>
<u>_____</u>	<u>_____</u>
<u>_____</u>	<u>_____</u>
<u>_____</u>	<u>_____</u>

Texture:
detrital magnetite oxidized to hematite
sorting is poor to moderate
grains are anhedral ranging in size from 0.1 mm to 0.25 mm







Diagram: PPL: X

XPL: _____

Scale Bar: 4.25 mm



4.25 mm

-  hematite/limonite
-  chert grain
-  quartz grain
-  veinlet quartz (secondary)
-  chert and iron cement
-  quartz filled fractures

Alteration Minerals:

limonite + goethite

Notes:

Formed from the hematite and magnetite

during hydrous oxidation in post

depositional stage of alteration probably

due to supergene weathering effects.

Stages of Alteration:

- (1) Authigenic chert formation during diagenesis.
 - (2) Alteration of pyrite and/or magnetite - to hematite during diagenesis.
 - (3) Supergene alteration produces goethite and limonite.
- _____
- _____
- _____

Paragenesis:

- (1) Detrital chert and quartz with sphene as nucleus of cherty grains deposited in situ.
 - (2) Magnetite and hematite coat some polycrystalline quartz and chert crystals.
 - (3) Introduction or remobilization of iron and silica, forms secondary chert - hematite
cement.
 - (4) Fracturing occurs probably along jointing planes and later introduction of quartz and
hematite produces veinlets.
- _____

Petrographer: D. Parry

Date: February 3, 1989

Thin Section: A-1

Hole: -

Depth: -

Rock Name: Silvery Black Shale

Location: Earn Outcrop A

Transparent Minerals:	%	Texture:
<u>illite with varying amounts</u>	<u>90%+</u>	<u>microcrystalline, poorly laminated, mixed</u>
<u>of carbon</u>		<u>in lamellae with carbonaceous material</u>
<u>microcrystalline quartz</u>	<u>~6%</u>	<u>bundles of acicular crystals</u>
<u>(chert) and chalcedonic</u>		<u><0.1 mm in length, perpendicular to</u>
<u>quartz</u>		<u>bedding probably diagenetic quartz</u>
		<u>growths as pressure shadows</u>

Opaque Minerals:	%	Texture:
<u>carbon</u>	<u>3-4%</u>	<u>disseminated, microcrystalline anhedral</u>
<u>hematite</u>	<u><1%</u>	<u>disseminated, subhedral, subcubic,</u>
		<u>probably oxidized pyrite</u>

Diagram: PPL: X XPL: _____ Scale Bar: 1.77 mm

Transmitted Plane Polarized Light (see plate 1)

2 figures
low concentration
of carbon + hematite

←----1----> zones of high concentration of magnetite + carbon

(i)

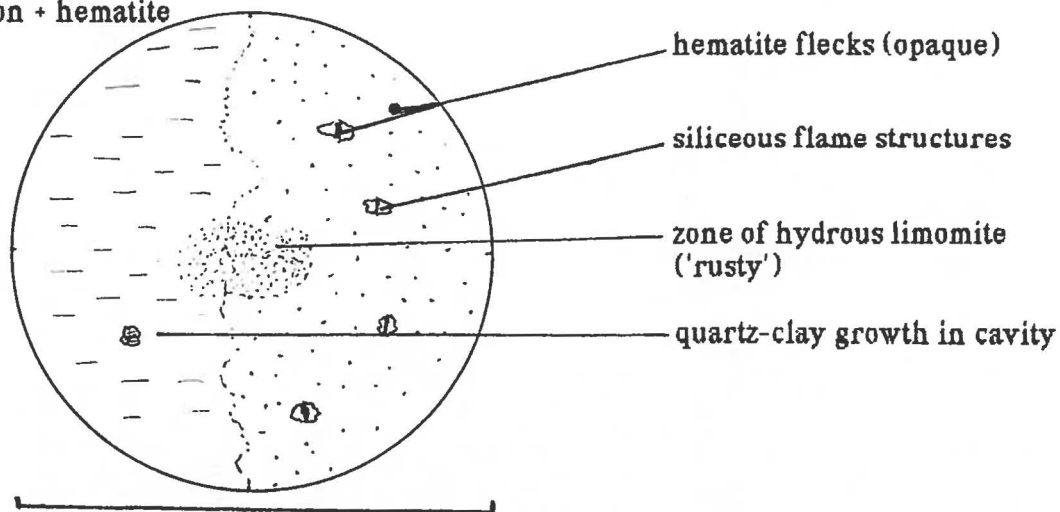
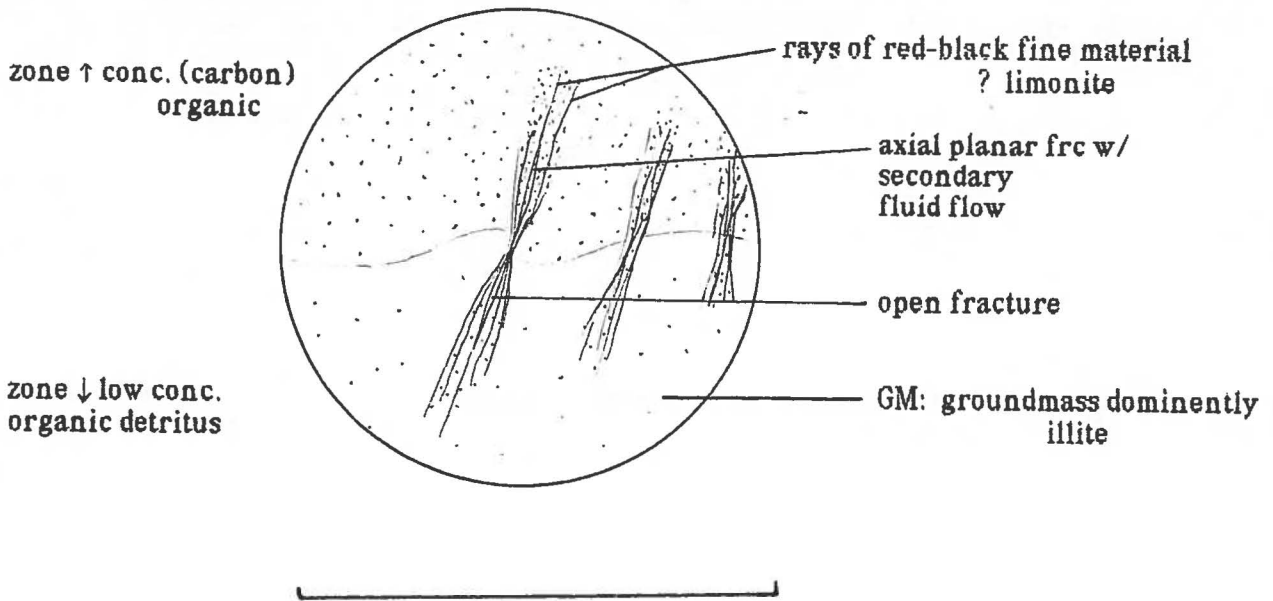


Diagram: PPL: X XPL: _____ Scale: 4.5 mm

figure (ii)



Alteration Minerals:	%	Texture(s):
<u>poorly defined carbon/iron</u>	<u>1-3%</u>	<u>fracture filling, semiplanar</u>
<u>in rays</u>	<u>_____</u>	<u>perpendicular to bedding</u>
<u>_____</u>	<u>_____</u>	<u>_____</u>
<u>_____</u>	<u>_____</u>	<u>_____</u>
<u>_____</u>	<u>_____</u>	<u>_____</u>

Stages of Alteration:

None: Weak diagenesis creating micronodules of carbon with associated flame structures,
perpendicular to bedding (foliation 1), parallel with (foliation 2) axial planar cleavage.

Paragenesis:

The shale unit has undergone regional metamorphism which has caused dehydration with
fluid movement between layers of different organic concentrations, fluids transport
perpendicular to (foliation 1) bedding.

Petrographer: D. ParryDate: February 3, 1989Thin Section: CHole: -Depth: -Rock Name: Shale (Silvery Black)Location: Lower Earn, Outcrop C**Transparent Minerals:****%****Texture:**clays and graphite~92%sub parallel orientation of carbonaceousmaterial in argillaceous shalequartz>7%microcrystalline (< 10 microns) anhedralcommonly forms aggregates perpendicularto primary foliation (bedding)quartz3%bundles of acicular crystals < 0.1 m formpressure shadows perpendicular to bedding**Opaque Minerals:****%****Texture:**none0%**Notes:**V. weak crenulation to primary foliation (assumed bedding).Planes of weakness observed by fracture filling of mount epoxy at 60-62 degrees toprimary foliation.Siliceous flame structures probably caused by movement of carbonaceous or hematiticmaterial and silica being produced from reaction of during diagenesis.No diagram; sample is highly carbonaceous and difficult to sketch.Similar to sample A-1, but richer in carbon.**Paragenesis:**Deposition of argillaceous sediment, diagenetic dewatering of clays, breakdown of organicmaterial to carbon, movement of carbonaceous material and the production of pressureshadows. An event of folding causes second foliation, crenulation of bedding.

Alteration Minerals:

%

Texture(s):

none

Stages of Alteration:

none

Paragenesis:

Deposition of carbonaceous shale and interstitial rare quartz.

The carbonaceous shale has undergone only minor diagenesis and only weak folding on a regional scale.

Petrographer: D. Parry

Date: March '89

Thin Section: 9048

Hole: -

Depth: -

Rock Name: Phosphatic Chert

Location: Lower Earn Formation

Transparent Minerals: %

Texture:

carbonaceous chert 80%

fine lamellae, orientation of flame

structure is nearly perpendicular to bedding

quartz and amorphous silica 7%

anhedral-amorphous fracture filling

surround pyrite band like in

slide 9049 (diag. 2)

Opaque Minerals: %

Texture:

pyrite 1%

disseminated (framboidal)

pyrite 12%

band of pyrite is approximately 1.0 cm in

thickness; the band is subhedral to cubic

relatively homogeneous composed of pyrite

Diagram: PPL: X
Reflected Light

XPL: _____

Scale Bar: 4.5 mm

Diagram 1

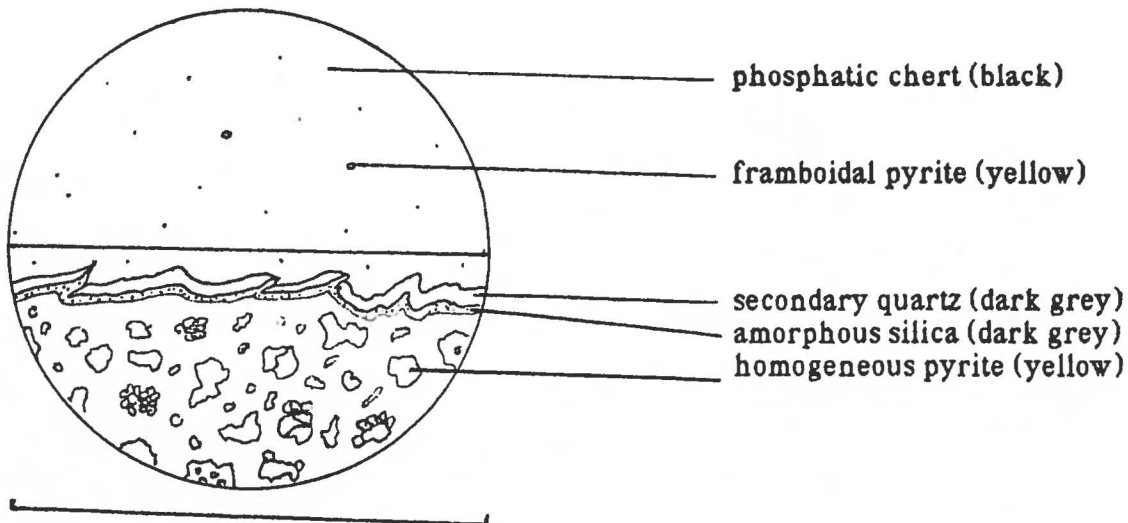


Diagram: PPL: X XPL: _____ Scale Bar: 4.5 mm

Diagram 2

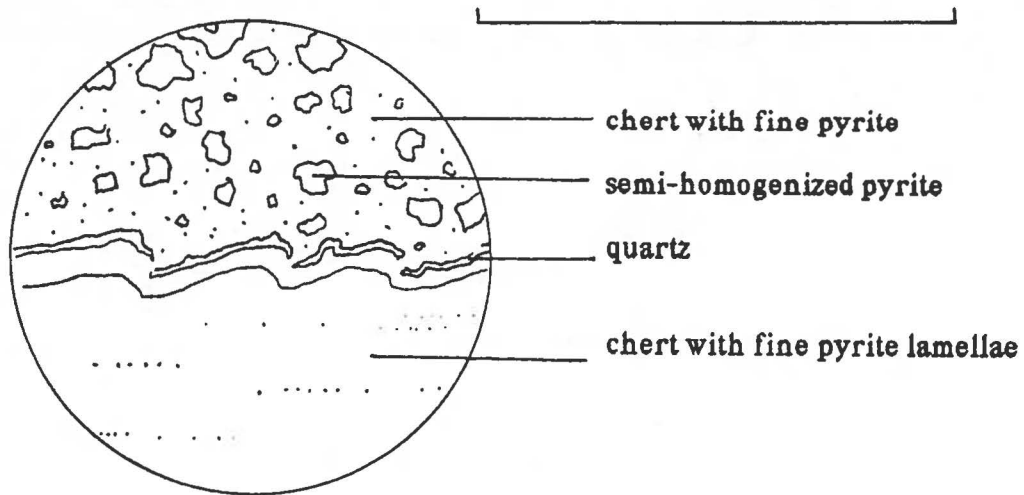


Diagram 3

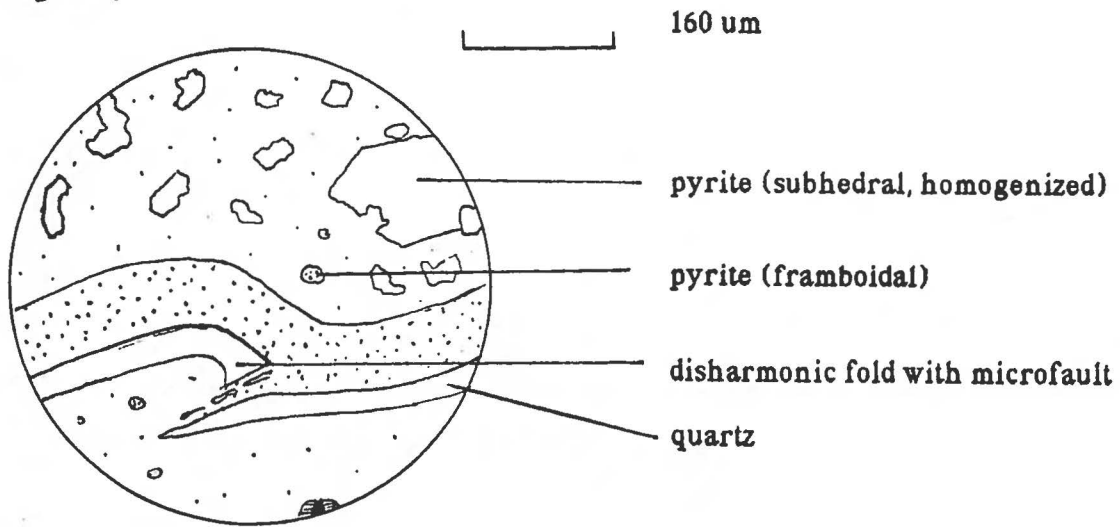
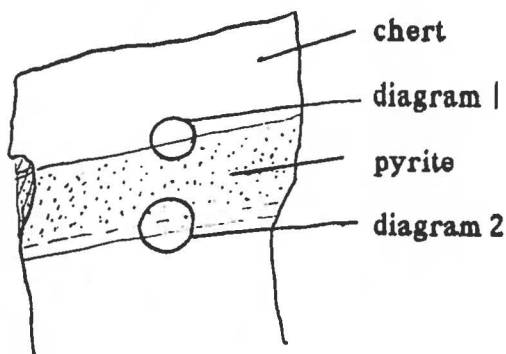


Diagram 4

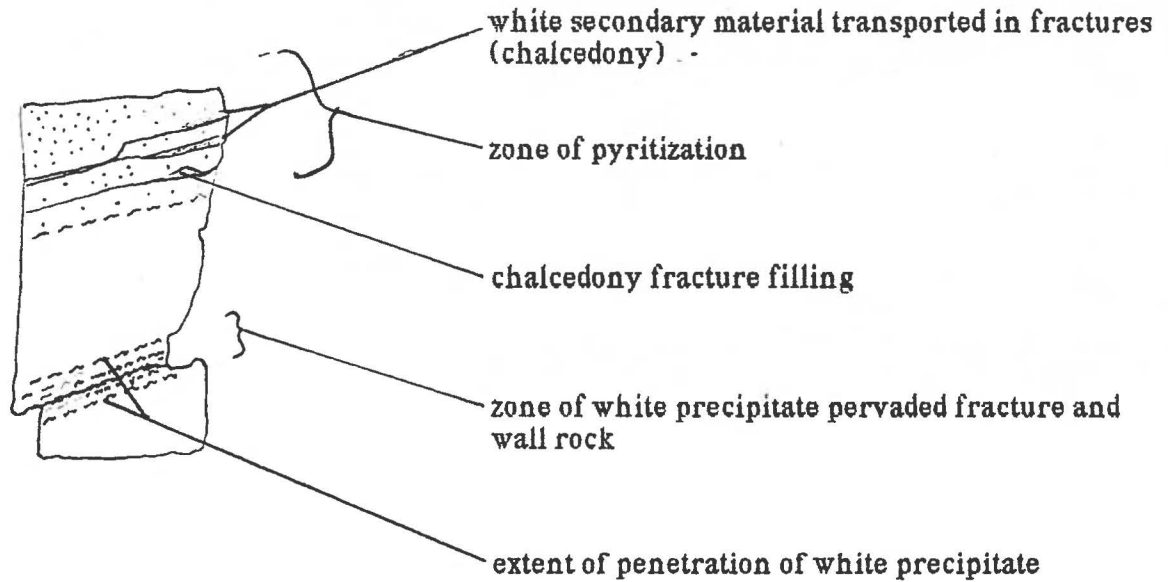
Hand Sample 9048

7 mm



9049

diag. 1

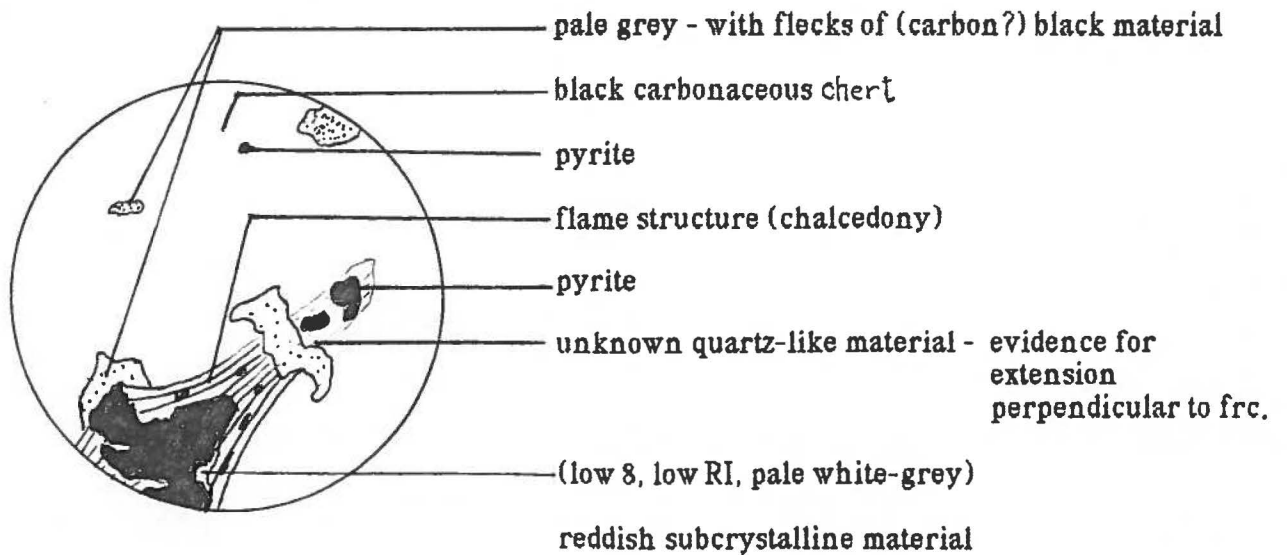


diag. 2

Sample 9049

Disseminated pyrite in carbonaceous chert
2 zones of chert parting - quartz precipitation.

ppl. transmitted 500X Field - 356 um



Alteration Minerals:	%	Texture(s):
<u>pyrite</u>	<u>< 0.5%</u>	<u>anhedral secondary</u>
<u>amorphous silica</u>	<u>5-8%</u>	<u>occasional replacement/etching of pyrite</u>
<u>single phase pyrite</u>	<u>3-4%</u>	<u>amorphous-framboidal, anhedral,</u>
<u>_____</u>	<u>_____</u>	<u>subhedral, subcubic, subdodecahedral</u>
<u>_____</u>	<u>_____</u>	<u>- occasionally pyritohedral</u>
<u>_____</u>	<u>_____</u>	<u>(homogenitization has occurred)</u>

Stages of Alteration:

No deformation or strong alteration, only weak diagenesis and cool fluid circulation has
occurred. Cubic pyrite has formed from framboidal pyrite. Diagenetic siliceous material
probably quartz infills fractures as last stage diagenesis.

Paragenesis:

Deposition of phosphatic-carbonaceous chert, minor diagenetic formation of pyrite ~1/2%.

Fracturing occurs, chert is etched and replaced by secondary pyrite and quartz.

Pyrite is diagenetically altered from microcrystalline framboidal pyrite to
semi-homogenized pyrite crystals.

Late weathering has oxidized pyrite to limonite on exposed surfaces.

Alteration Minerals:

%

Texture(s):

chalcedonic quartz

forms flame structures around pyrite

crystals

Stages of Alteration:

1) Diagenetic recrystallization of quartz forming chalcedonic flame structures.

2) Pyrite diagenetically grows from framboidal pyrite to sub-cubic homogenized pyrite.

Paragenesis:

Deposition of cherty shale with lamellae of pyrite.

Diagenetic growth of pyrite and chalcedonic flame structures.

Homogenization of framboidal pyrite to sub-cubic pyrite.

Petrographer: D. Parry

Date: March 26, 1989

Thin Section: Nick Hole 4 (R6) Hole: 4 Depth: 86.56 m

Rock Name: Carbonaceous Chert w/ Strain Filled Chalcedonic Quartz

Transparent Minerals:	%	Texture:
<u>chalcedonic quartz</u>	<u>1-2%</u>	<u>strained, interstitial between lamination</u>
<u>_____</u>	<u>_____</u>	<u>of carb. shale occasionally surrounding</u>
<u>_____</u>	<u>_____</u>	<u>pyrite as flame structures</u>
<u>carbonaceous chert</u>	<u>98-99%</u>	<u>finely laminar bedded</u>
<u>_____</u>	<u>_____</u>	<u>strongly carbonaceous</u>
Opaque Minerals:	%	Texture:
<u>pyrite</u>	<u>< 0.5%</u>	<u>framboidal - subhedral (< 75 um)</u>
<u>_____</u>	<u>_____</u>	<u>finely laminar beds of pyrite</u>

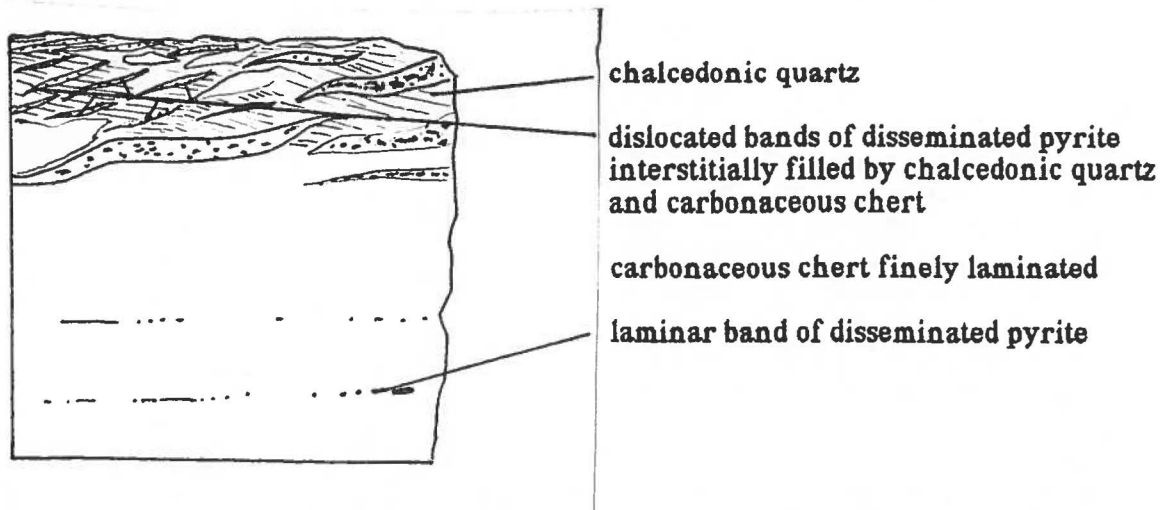
Paragenesis:

- (1) Deposition of carbonaceous chert and pyrite.
- (2) Dislocation by extension/shear & precipitation of formation fluids rich in silica
(chalcedonic quartz).

Diagram:

Scale: 6 cm = 8 mm

diagram of silicate/pyrite band in sample (R6) using binocular microscope



Alteration Minerals:

limonite

Notes:

amorphous iron oxide deposited during
formation of siliceous fibres

Stages of Alteration:

Clays of argillite undergo dewatering, due to compaction. Fluids move iron from pyrite and
other metals in shale to sites of deposition such as: fractures and tension gashes in argillite
in the vicinity of the cherty argillite/limestone ball member contact.

Paragenesis:

Argillaceous sediment deposited. Loading occurs, burial metamorphism and diagenesis
alter clays to mainly illite. Fractures open and allow movement of fluids into cavities
formed along chert-limestone ball member interface. Quartz and chalcedony are
precipitated onto walls of the cavities. Fluids deposit silica and 'clog' cavity. Further
shear movement reopens veinlets and fluids again migrate through the rock and
precipitate pyrite (probably) which later oxidizes to limonite.

Petrographer: D. Parry Date: _____

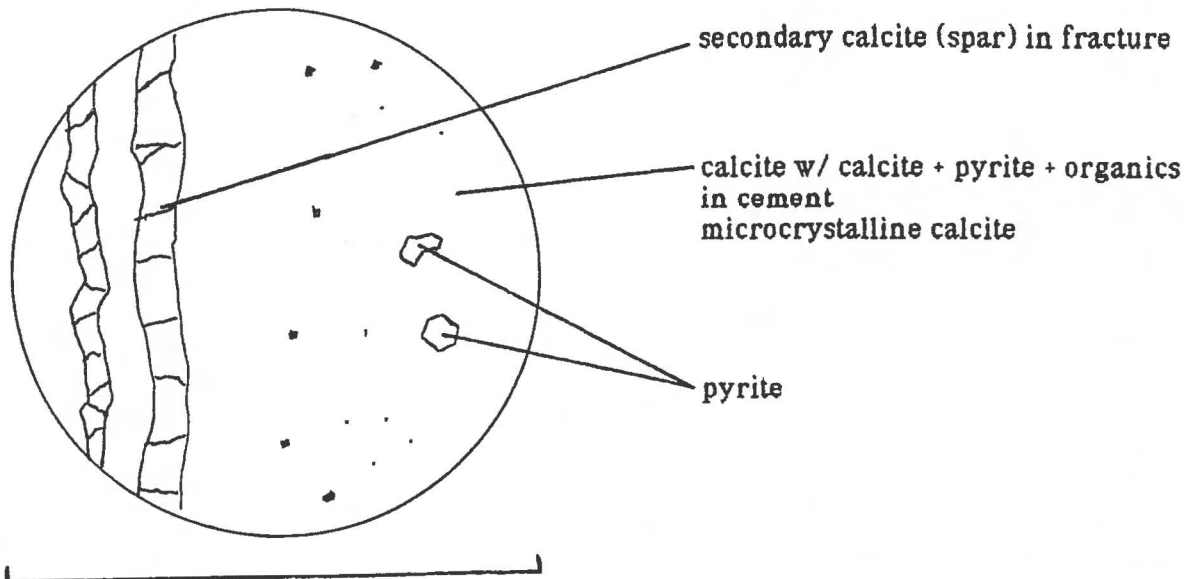
Thin Section: Nick Hole (R4) Hole: 4 Depth: 41.85 m

Rock Name: Disseminated Pyrite in Finely Laminated Limestone (Grey White Speckled Limestone)

Transparent Minerals:	%	Texture:
<u>calcite</u>	<u>92%</u>	<u>recrystallized, calcite cement, trapped</u>
_____	_____	<u>black organic flecks in cement, common</u>
_____	_____	<u>50 μm average grain size</u>
<u>organic material</u>	<u><2%</u>	<u>amorphous - trapped in cement</u>
_____	_____	_____

Opaque Minerals:	%	Texture:
<u>pyrite</u>	<u>~6%</u>	<u>subhedral to framboidal, hard to</u>
_____	_____	<u>distinguish - 120 μm max. grain size</u>
_____	_____	<u>- 2 μm max. grain size</u>
_____	_____	<u>- ~6 μm average grain size</u>

Diagram: PPL: X XPL: _____ Scale Bar: 1.77 mm



Alteration Minerals:	%	Texture(s):
<u>none</u>	<u></u>	<u></u>
<u></u>	<u></u>	<u></u>
<u></u>	<u></u>	<u></u>
<u></u>	<u></u>	<u></u>
<u></u>	<u></u>	<u></u>

Stages of Alteration:

Diagenetic crystallization of pyrite from framboids to cubes and dodecahedrons (rare)
usually large ~ 120 μ m.

Paragenesis:

Primary calcite deposition.

Diagenetic crystallization of pyrite (no sphalerite on margins).

Secondary calcite in fractures perpendicular to bedding (probably axial planar cleavage
w/ calcite ppt.).

Petrographer: D. Parry

Date: March 26, 1989

Thin Section: Nick Hole (R2)

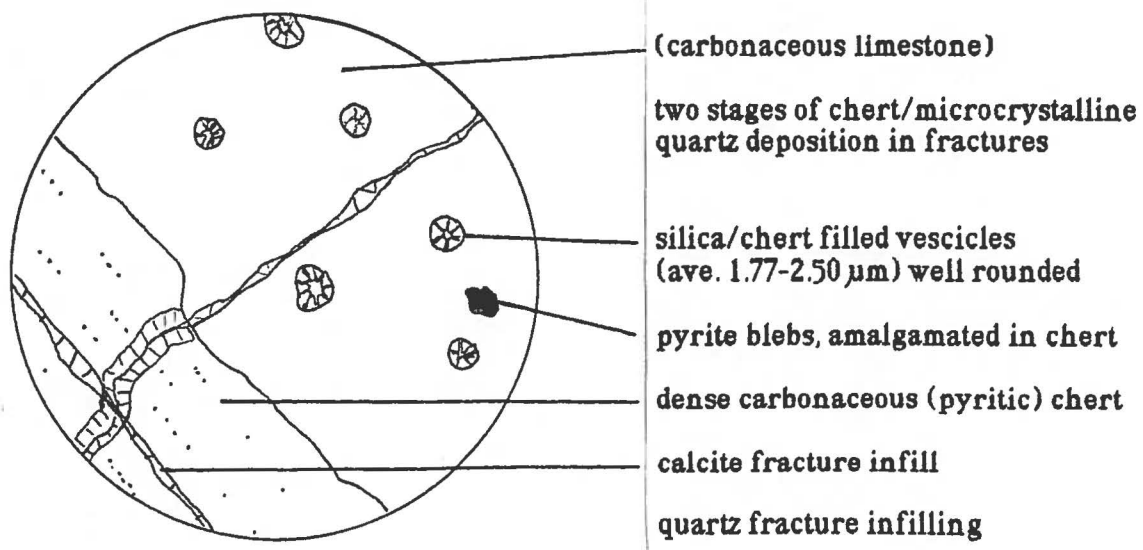
Hole: 4

Depth: 141.80 m

Rock Name: Crackle Brecciated Cherty Shale w/ Calcite (Fracture Filling Interstitial)

Transparent Minerals:	%	Texture:
<u>chert</u>	<u>10%</u>	<u>interstitial between dislocated lamellae</u> <u>of carbonaceous shale</u>
<u>calcite</u>	<u>~35%</u>	<u>interstitial in chert/shale clasts</u> <u>fracture filling</u>
<u>carbonaceous shale (lithic fragments)</u>	<u>20%</u>	<u>angular fragments w/ calcite fracture fill</u>
<u>carbonaceous chert</u>	<u>25%</u>	<u>forms pellets ~0.3 to 1.5 mm diameter</u> <u>subangular clasts separated by dislocated</u> <u>lamellae of carbonaceous shale</u>
Opaque Minerals:	%	Texture:
<u>pyrite</u>	<u>< 1%</u>	<u>forms aggregates of micro cubic pyrite</u>

Diagram: PPL: _____ XPL: X Scale Bar: 1.77 mm



Alteration Minerals:

%

Texture(s):

calcite

forms in fractures, in stages of dissolution

Stages of Alteration:

none visible

Paragenesis:

(1) Limestone deposited.

(2) Carbonaceous chert deposited with diagenetic pyrite.

(3) Dissolution of limestone.

(4) Brecciation of limestone and chert (probably collapse due to dissolution).

(5) Deposition of silica in fractures.

(6) Deposition of calcite in fractures.

(7) Deposition of silica (2nd stage of silica).

(8) Crystallization of subhedral to euhedral microcrystalline pyrite.

Petrographer: D. Parry

Date: March 26, 1989

Thin Section: Nick Hole (R5)

Hole: 4

Depth: 80.74 m

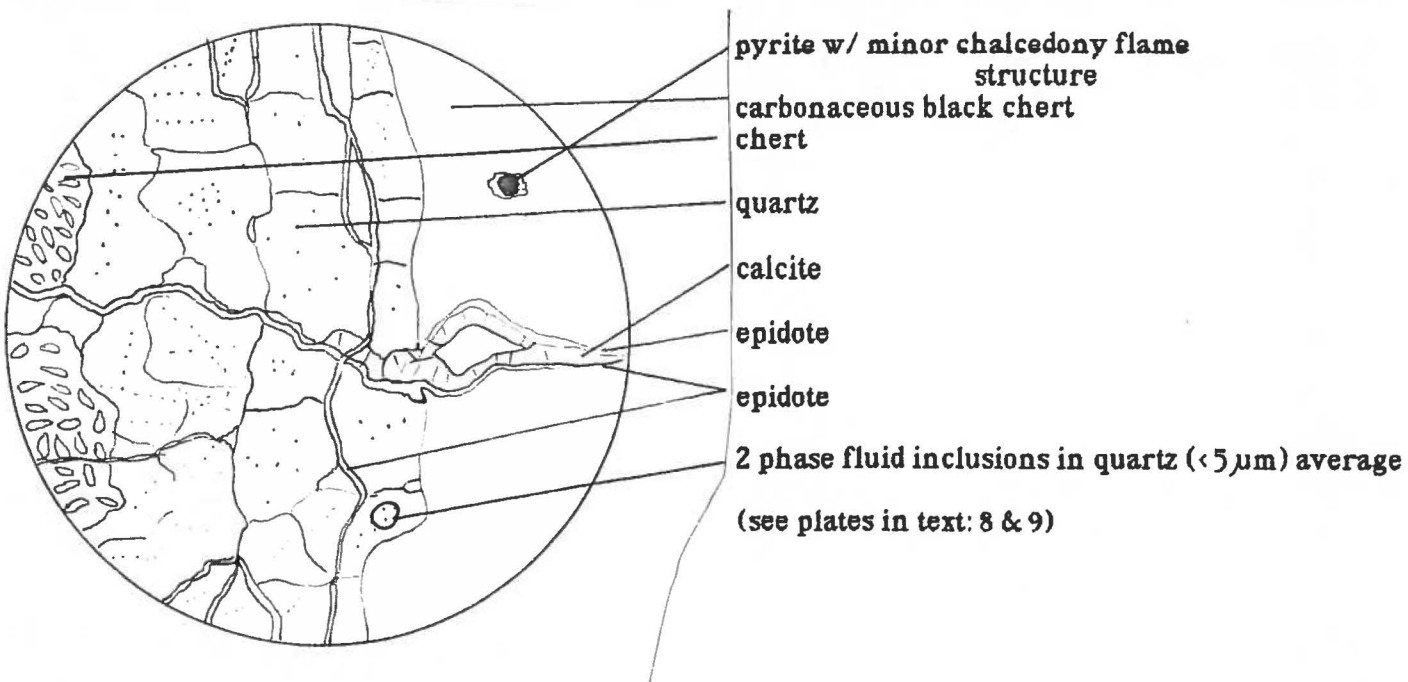
Rock Name: Quartz-Calcite Supported Lithic Breccia

Transparent Minerals:	%	Texture:
<u>quartz</u>	<u>23%</u>	<u>anhedral-subhedral, ave. grn. ~ 0.75 mm diameter</u> <u>subhedral (fractured)</u> <u>(lithic frag. (15-20%))</u>
<u>chert</u>	<u>30%</u>	<u>anhedral, ave. gr. size ~ 0.65 mm dia. (remobilized (10-15%))</u>
<u>calcite</u>	<u>18%</u>	<u>subhedral; ave. gr. size ~ 0.50 mm dia.</u> <u>(max. size 8 mm x 8 mm)</u>
<u>carbonaceous shale fragments</u>	<u>25%</u>	<u>angular lithic fragments (ave. size 2.5 mm x 6 mm)</u> <u>(min. size ~ 0.25 mm dia.)</u>
<u>epidote</u>	<u>4%</u>	<u>anhedral, fracture filling</u>
Opaque Minerals:	%	Texture:
<u>pyrite</u>	<u>< 1%</u>	<u>subhedral (near cubic) in the</u> <u>cherty-shale fragments only</u>

Diagram: PPL: _____

XPL: X

Scale Bar: 4.5 mm/f.v.



360 μ m

calcite rhombs

chert w/ lithic frags

epidote fracture filling

Plate A
(PPL:)

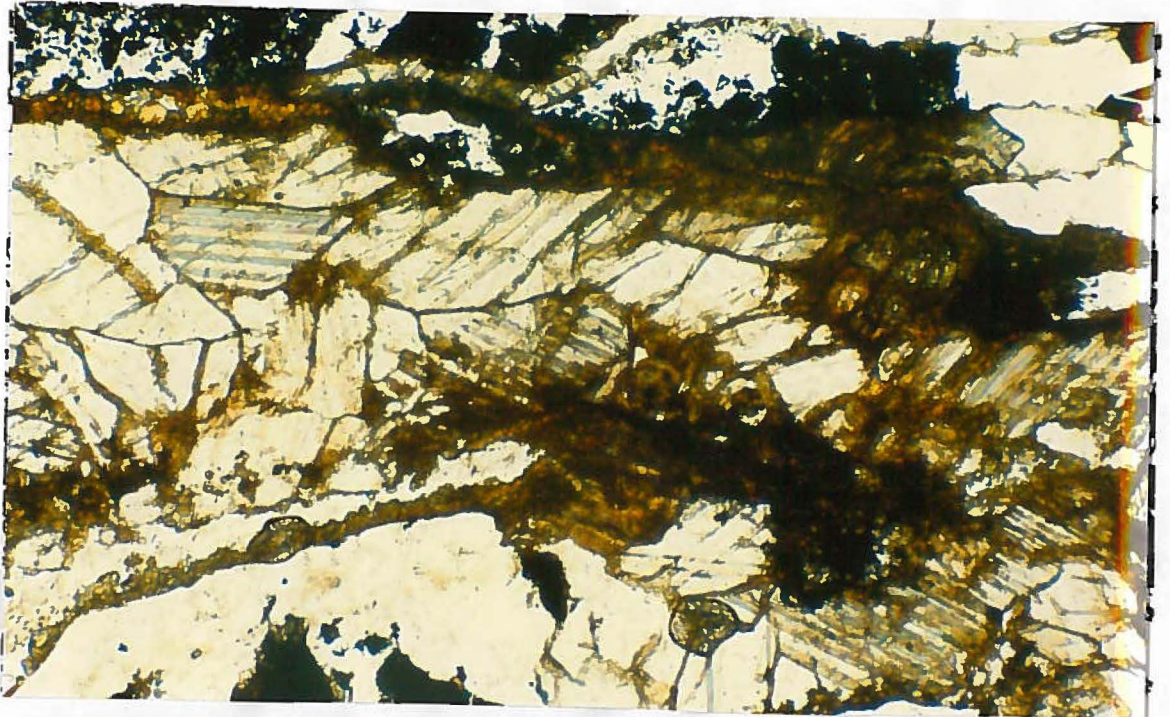
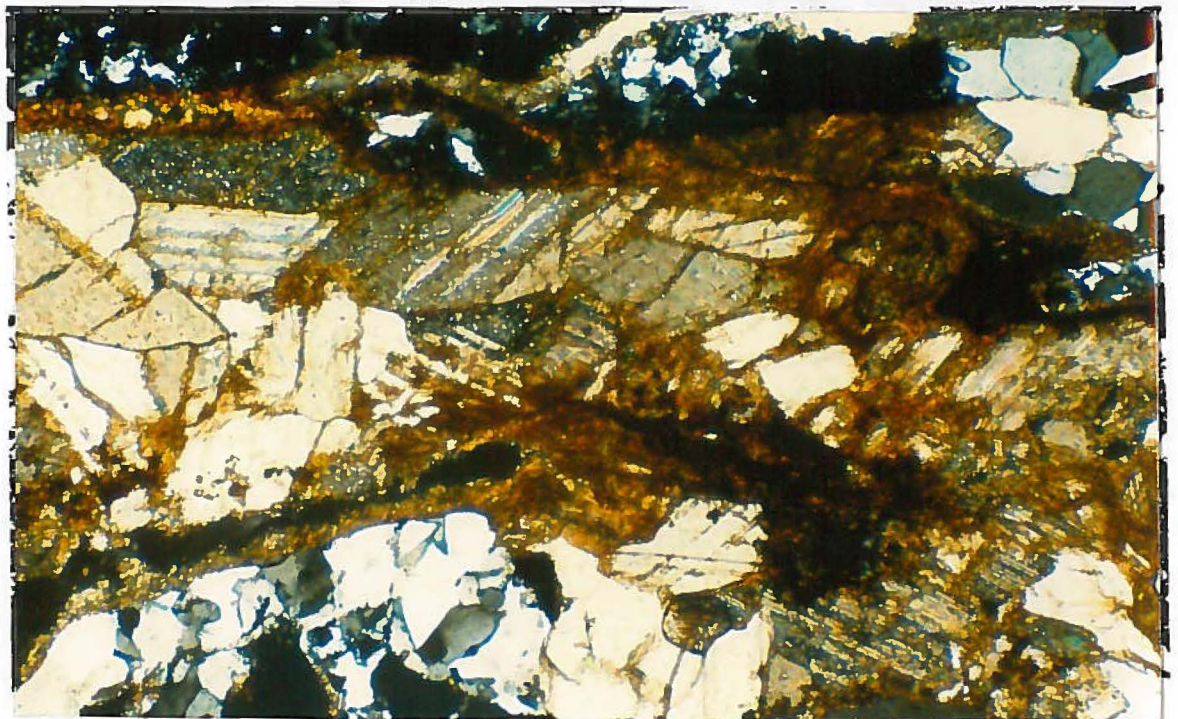


Plate B
(XPL:)



Alteration Minerals:

quartz, chert, epidote, calcite

Notes:

by determining the cross cutting

relations the paragenesis of the

deposition of alteration minerals was

determined

Stages of Alteration:

Fracture/fault and introduction of hydrothermal fluid.

Deposition of quartz, 2nd movement of fluid pathway.

Deposition of chert + quartz, 3rd movement of fluid pathway.

Deposition of epidote and calcite.

(See Plates, A & B)

Paragenesis:

(1) Fracturing of host rock.

(2) Injection of quartz rich fluid.

(3) Precipitation of quartz along lithic fragments and host rock contacts.

(4) Fracturing of quartz supported breccia.

(5) Precipitation of chert and later fracture in fluid pathway.

(6) Precipitation of epidote.

(7) Precipitation of calcite.

(8) Precipitation on epidote.

Petrographer: D. ParryDate: January 1, 1989Thin Section: Outcrop A-1 Hole: -Depth: -Rock Name: Mineralized VaesiteLocation: Outcrop A-1 (Olfert Creek)

Transparent Minerals:	%	Texture:
<u>carbonaceous and/or</u>	<u>~88%</u>	<u>amorphous bands which are present</u>
<u>phosphatic chert</u>	<u>_____</u>	<u>between layers of sulfides (also forms</u>
<u>quartz</u>	<u>2%</u>	<u>nodules)</u>
<u>_____</u>	<u>_____</u>	<u>microcrystalline to subcrystalline forms</u>
<u>unidentified brown material</u>	<u>< 1%</u>	<u>in cherty nodules and is exsolved</u>
<u>probably bitumen</u>	<u>_____</u>	<u>from pyrite bands</u>
<u>_____</u>	<u>_____</u>	<u>amorphous masses commonly form</u>
Opaque Minerals:	%	Texture:
<u>pyrite and melnikovite</u>	<u>43%</u>	<u>parallel to bands of pyrite/melnikovite/</u>
<u>_____</u>	<u>_____</u>	<u>and vaesite</u>
<u>_____</u>	<u>_____</u>	<u>forms subcolliform bands, subhedral</u>
<u>_____</u>	<u>_____</u>	<u>aggregates of homogenized pyrite</u>
<u>vaesite</u>	<u>13%</u>	<u>interstitially in bands of homogenized</u>
<u>_____</u>	<u>_____</u>	<u>pyrite framboids</u>
<u>_____</u>	<u>_____</u>	<u>the vaesite is finely intergrown with the</u>
<u>_____</u>	<u>_____</u>	<u>pyrite</u>
<u>sphalerite</u>	<u>4%</u>	<u>commonly forms as intergrown with</u>
<u>_____</u>	<u>_____</u>	<u>pyrite and vaesite, predominantly on</u>
<u>_____</u>	<u>_____</u>	<u>either top or bottom of each layer</u>

See plates in text, numbers 10 through 16.

Alteration Minerals:	%	Texture(s):
<u>quartz (microcrystalline)</u>	<u>< 1%</u>	<u>Is exsolved from the chert and pyrite</u>
<u>_____</u>	<u>_____</u>	<u>layers. It forms siliceous fibres in some</u>
<u>_____</u>	<u>_____</u>	<u>samples probably formed from effect of</u>
<u>_____</u>	<u>_____</u>	<u>soft sediment deformation</u>
<u>_____</u>	<u>_____</u>	<u>_____</u>

Stages of Alteration:

Diagenetic growth of pyrite framboids and other sulfides become incorporated.

Soft sediment deformation causes fine lamellae to be disturbed with pyrite-vaesite layers
conforming to harder chert nodules.

Paragenesis:

Diagenetic growth of pyrite framboids at sediment water interface.

Incorporation of nickel into pyrite amalgamations and bands in the form of vaesite.

Sphalerite is incorporated into the pyrite during homogenetization, typically on only one
side of the bands.

Formation of cherty nodules, possibly by biogenesis.

Deposition of cherty sediment in fine rhythmic beds between the pyrite-vaesite layers.

APPENDIX B

SUMMARY OF MINERALIZED ZONE ASSAY DATA

**APPENDIX B
MINERALIZED ZONE ASSAY DATA**

Location	Ni %	Zn PPM	Pb PPM	Mo PPM	Ba PPM	As PPM	Pt PPB	Pd PPB	Au PPM	Ag PPM
Outcrop A	*5.8	8000	-	-	20000	-	-	-	-	-
	5.5	9586	-	1454	90	1254	540	210	58	10.1
	4.6	138	-	2650	3000	2990	310	150	50	7.8
Outcrop B	*2.9	2530	-	-	2800	-	-	-	-	-
	3.6	5710	-	1265	40	2560	280	120	86	9.2
Outcrop D	+4.4	772	-	1485	60	2260	5	<2	6	12.6
DDH 88-1 (1801)	2.9	6590	76	1570	<10	1985	<62	62	-	6.8
DDH 88-3 (1035)	5.0	10040	130	3920	170	3930	434	155	-	12.2
	$\frac{-}{\bar{x}}$ 4.34 %	$\frac{-}{\bar{x}}$ ~5420 ppm								

NOTES:

- Archer, Cathro (1988)
 - * Cominco (1987)
 - + Strongly Oxidized
 - () Assay Number
- All data supplied by (Archer, Cathro & Associates Ltd., (1981))

APPENDIX C

DRILL LOGS

APPENDIX D

ASSAY DATA FROM DRILL LOGS



Chemex Labs Ltd.

Analytical Chemists * Geochemists * Registered Assayers
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 BRITISH COLUMBIA, CANADA V7J-2C1
 PHONE (604) 984-0221

RCHER CATHRO & ASSOC. (1981) LTD.

BOX 4127
 WHITEHORSE, Y.T.
 Y1A 3S9

Project : NDU-NICK
 Comments :

Page # :
 Tot. # : 1
 Date : 29-SEP-88
 Invoice # : I-8824185
 P.O. # : NONE

CERTIFICATE OF ANALYSIS A8824185

SAMPLE DESCRIPTION	PREP CODE		Al	Ag	As	Ba	Bc	Bi	Ca	Cd	Co	Cr	Cu	Fe	Ga	Hg	K	La	Mg	Mn	Mo
			%	ppm	ppm	ppm	ppm	ppm	%	ppm	ppm	ppm	ppm	%	ppm	ppm	%	ppm	%	ppm	ppm
1576	205	238	0.43	1.8	30	340	< 0.5	< 2	1.02	16.0	8	254	45	1.85	< 10	< 1	0.14	10	0.08	114	59
1577	205	238	0.64	1.0	20	640	< 0.5	4	8.61	7.0	3	152	34	1.03	20	2	0.20	< 10	0.19	109	41
1578	205	238	0.28	1.2	25	360	< 0.5	< 2	0.16	13.0	3	183	70	1.23	< 10	1	0.09	< 10	0.03	75	41
1579	205	238	0.35	1.2	30	360	< 0.5	< 2	0.43	6.0	6	167	48	1.40	< 10	1	0.12	10	0.04	93	64
1580	205	238	0.36	0.6	35	310	< 0.5	< 2	1.73	6.0	7	110	54	1.72	< 10	< 1	0.12	20	0.07	171	85
1581	205	238	0.13	0.8	5	280	< 0.5	18	>15.00	0.5	1	23	17	0.20	20	< 1	0.03	< 10	0.49	118	6
1582	205	238	0.25	0.8	15	260	< 0.5	4	8.45	2.5	2	123	15	0.70	10	1	0.08	< 10	0.16	77	41



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Project : NDU-NICK
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Page # : 1-B
 Tot. # : 1
 Date : 29-SEP-88
 Invoice # : I-8824185
 P.O. # : NONE

CERTIFICATE OF ANALYSIS A8824185

SAMPLE DESCRIPTION	PREP CODE		Na	Ni	P	Pb	Sb	Sc	Sr	Ti	Tl	U	V	W	Zn
			%	ppm	ppm	ppm	ppm	ppm	ppm	%	ppm	ppm	ppm	ppm	ppm
1576	205	238	< 0.01	178	1440	16	5	1	119	< 0.01	< 10	< 10	207	< 5	1130
1577	205	238	0.01	96	>10000	12	5	4	674	< 0.01	< 10	10	304	10	530
1578	205	238	< 0.01	147	410	26	5	< 1	43	< 0.01	< 10	< 10	160	< 5	986
1579	205	238	< 0.01	209	510	16	< 5	< 1	55	< 0.01	< 10	< 10	176	< 5	484
1580	205	238	< 0.01	303	560	8	< 5	1	142	< 0.01	< 10	< 10	298	< 5	530
1581	205	238	0.01	19	220	6	< 5	< 1	563	< 0.01	< 10	< 10	77	< 5	71
1582	205	238	< 0.01	134	350	16	< 5	1	364	< 0.01	< 10	< 10	225	< 5	265



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Y1A 3S9

Project : NDU NICK
Comments :

Page 1 of 1-A
Tot. Pgs. : 2
Date : 20-SEP-88
Invoice # : I-8823286
P.O. # : NONE

CERTIFICATE OF ANALYSIS A8823286

SAMPLE DESCRIPTION	PREP CODE	Al %	Ag ppm	As ppm	Ba ppm	Be ppm	Bi ppm	Ca %	Cd ppm	Co ppm	Cr ppm	Cu ppm	Fe %	Ga ppm	Hg ppm	K %	La ppm	Mg %	Mn ppm	Mo ppm
1802	205 238	0.72	0.4	10	280	0.5	2	0.09	1.5	7	84	34	1.58	< 10	< 1	0.21	10	0.10	61	57
1803	205 238	0.96	0.6	15	190	0.5	< 2	0.57	2.0	51	102	64	1.68	< 10	< 1	0.28	10	0.12	58	54
1804	205 238	0.82	0.2	20	180	0.5	2	0.95	1.5	18	81	47	2.09	< 10	< 1	0.24	10	0.10	40	76
1805	205 238	0.82	0.4	15	110	0.5	4	0.09	1.0	8	85	33	1.54	< 10	< 1	0.23	10	0.10	38	54
1806	205 238	1.00	0.6	20	80	0.5	4	0.10	1.0	10	84	41	2.06	< 10	< 1	0.32	10	0.17	76	59
1807	205 238	0.65	0.6	15	130	0.5	2	0.14	1.0	6	98	40	1.50	< 10	< 1	0.20	10	0.10	57	55
1808	205 238	0.67	0.4	10	220	0.5	2	0.50	1.0	4	112	30	1.20	< 10	< 1	0.20	10	0.20	62	33
1809	205 238	0.75	0.6	15	100	0.5	< 2	0.17	1.0	6	102	33	1.62	< 10	< 1	0.23	10	0.09	59	58
1810	205 238	0.98	0.6	20	100	0.5	< 2	0.42	1.0	7	107	33	1.85	< 10	< 1	0.31	20	0.15	90	60
1811	205 238	0.87	0.6	10	140	0.5	4	0.41	1.5	5	89	33	1.51	< 10	< 1	0.26	10	0.17	107	40
1812	205 238	0.87	0.6	20	140	1.0	< 2	0.41	5.0	8	108	51	1.81	< 10	1	0.28	10	0.16	155	48
1813	205 238	1.08	0.6	20	130	1.0	< 2	0.38	2.5	7	91	37	1.79	< 10	< 1	0.33	10	0.22	85	50
1814	205 238	0.84	0.6	20	300	0.5	2	0.20	4.0	5	122	25	1.15	< 10	< 1	0.25	10	0.21	54	47
1815	205 238	1.22	1.2	40	120	1.0	< 2	0.20	33.0	8	141	75	1.65	< 10	< 1	0.34	20	0.13	66	56
1816	205 238	1.28	< 0.2	20	500	1.0	2	2.25	15.5	7	156	37	1.20	< 10	1	0.35	20	0.86	121	37
1817	205 238	1.17	0.2	30	230	1.0	4	1.39	9.5	9	126	48	1.60	< 10	< 1	0.30	20	0.30	88	63
1818	205 238	0.77	0.4	15	480	0.5	4	0.09	5.5	15	150	22	1.04	< 10	< 1	0.13	< 10	0.07	39	35
1819	205 238	0.64	0.2	10	500	0.5	2	0.11	1.5	16	154	17	0.81	< 10	< 1	0.11	< 10	0.05	34	37
1820	205 238	0.70	0.4	65	330	0.5	4	0.18	3.0	10	136	27	1.53	< 10	< 1	0.17	20	0.08	49	144
1821	205 238	1.13	< 0.2	35	340	1.0	2	2.61	2.0	9	138	55	2.06	< 10	1	0.33	30	0.16	205	88
1822	205 238	0.65	< 0.2	15	680	0.5	4	7.66	1.5	6	142	48	1.10	< 10	< 1	0.20	< 10	0.19	168	46
1823	205 238	0.35	< 0.2	5	380	< 0.5	< 2	2.99	1.5	2	188	37	0.54	< 10	< 1	0.10	< 10	0.07	59	26
1824	205 238	1.19	< 0.2	25	480	1.0	< 2	2.57	2.5	12	124	26	1.96	< 10	1	0.39	20	0.24	261	74
1825	205 238	0.30	< 0.2	5	320	< 0.5	10	>15.00	0.5	2	86	65	0.35	< 10	< 1	0.07	< 10	0.49	103	15
1826	205 238	0.86	< 0.2	10	610	0.5	4	11.55	1.5	6	101	23	0.92	< 10	2	0.29	< 10	0.28	121	48
1827	205 238	0.53	< 0.2	10	390	0.5	8	>15.00	0.5	3	80	54	0.44	< 10	< 1	0.15	< 10	0.49	78	23
1828	205 238	0.93	< 0.2	20	570	0.5	2	8.74	1.5	4	96	21	0.93	< 10	2	0.29	< 10	0.32	63	50
1829	205 238	1.21	< 0.2	25	350	1.0	2	10.60	3.0	6	111	32	1.40	< 10	2	0.40	< 10	0.31	92	70
1830	205 238	0.20	< 0.2	< 5	370	< 0.5	12	>15.00	< 0.5	2	17	52	0.17	< 10	< 1	0.05	< 10	0.52	87	8
1831	205 238	1.47	0.2	10	130	1.5	< 2	0.94	2.0	37	118	13	4.57	< 10	2	0.37	20	0.43	955	51
1832	205 238	1.31	0.4	25	40	1.0	2	0.11	2.5	10	75	37	2.57	< 10	< 1	0.35	10	0.24	76	57
1833	205 238	1.09	0.4	15	110	1.0	< 2	0.07	4.0	7	91	31	1.59	< 10	< 1	0.27	10	0.21	43	43
1834	205 238	1.00	0.2	20	100	1.0	< 2	0.03	5.0	9	95	35	1.76	< 10	< 1	0.26	10	0.10	42	61
1835	205 238	1.03	< 0.2	20	140	1.0	< 2	0.47	2.0	8	84	42	1.77	< 10	< 1	0.30	10	0.09	47	60
1836	205 238	0.72	< 0.2	20	270	0.5	4	0.97	1.0	8	56	48	1.74	< 10	< 1	0.12	10	0.93	140	56
1837	205 238	0.92	< 0.2	20	290	0.5	2	0.39	1.0	8	44	46	1.65	< 10	< 1	0.14	10	0.81	102	56
1838	205 238	0.62	< 0.2	20	120	0.5	< 2	0.10	1.0	6	64	33	1.64	< 10	< 1	0.17	10	0.08	39	55
1839	205 238	0.67	0.2	20	120	0.5	< 2	0.26	1.0	7	49	39	1.81	< 10	< 1	0.20	10	0.07	65	74
1840	205 238	0.61	0.2	20	110	0.5	< 2	0.22	1.0	5	78	28	1.52	< 10	< 1	0.19	10	0.09	67	55
1841	205 238	0.82	0.2	15	120	0.5	< 2	0.38	3.5	6	58	37	1.48	< 10	< 1	0.25	10	0.17	91	49

CERTIFICATION : _____



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Project : NDU NICK

Comments :

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Tot. : 2
Date : 20-SEP-88
Invoice # : I-8823286
P.O. # : NONE

CERTIFICATE OF ANALYSIS A8823286

SAMPLE DESCRIPTION	PREP CODE	Na %	Ni ppm	P ppm	Pb ppm	Sb ppm	Sc ppm	Sr ppm	Ti %	Tl ppm	U ppm	V ppm	W ppm	Zn ppm
1802	205 238	0.01	110	230	14	< 5	1	29	0.01	< 10	< 10	110	< 5	208
1803	205 238	0.01	122	200	12	< 5	2	30	0.01	< 10	< 10	207	< 5	292
1804	205 238	0.01	142	230	14	< 5	1	45	0.01	< 10	< 10	134	< 5	206
1805	205 238	0.01	104	250	10	< 5	2	36	0.01	< 10	< 10	117	< 5	181
1806	205 238	0.01	124	150	10	< 5	3	19	0.01	< 10	< 10	141	< 5	215
1807	205 238	0.01	107	220	6	< 5	1	32	< 0.01	< 10	< 10	139	< 5	231
1808	205 238	0.01	74	360	2	< 5	1	84	< 0.01	< 10	< 10	153	< 5	154
1809	205 238	0.01	116	310	10	< 5	1	36	0.01	< 10	< 10	165	< 5	196
1810	205 238	0.01	118	390	14	5	2	84	0.01	< 10	< 10	250	< 5	223
1811	205 238	0.01	100	370	10	5	2	56	0.01	< 10	< 10	274	< 5	238
1812	205 238	0.01	122	170	14	5	2	38	0.01	< 10	< 10	446	< 5	738
1813	205 238	0.01	123	690	14	5	2	73	0.01	< 10	< 10	491	< 5	386
1814	205 238	0.01	134	200	4	5	1	26	0.01	< 10	< 10	383	< 5	366
1815	205 238	0.01	173	600	16	15	3	44	0.01	< 10	< 10	638	< 5	1850
1816	205 238	0.01	132	2130	10	5	3	231	0.01	< 10	< 10	576	< 5	982
1817	205 238	0.01	172	470	6	10	3	129	0.01	< 10	< 10	451	< 5	710
1818	205 238	0.01	190	600	2	5	1	188	< 0.01	< 10	< 10	254	< 5	774
1819	205 238	0.01	188	440	4	5	1	121	< 0.01	< 10	< 10	177	< 5	551
1820	205 238	0.01	902	610	6	5	1	55	< 0.01	< 10	< 10	208	< 5	489
1821	205 238	0.01	295	560	4	5	3	98	0.01	< 10	< 10	735	< 5	431
1822	205 238	0.01	155	640	10	5	3	260	< 0.01	< 10	< 10	576	< 5	244
1823	205 238	0.01	83	280	6	< 5	< 1	147	< 0.01	< 10	< 10	261	< 5	181
1824	205 238	0.01	249	440	12	5	3	108	0.01	< 10	< 10	1020	< 5	445
1825	205 238	0.01	47	220	2	< 5	1	592	< 0.01	< 10	< 10	279	< 5	100
1826	205 238	0.01	144	340	12	< 5	3	410	0.01	< 10	< 10	762	< 5	246
1827	205 238	0.01	75	180	8	< 5	2	358	< 0.01	< 10	< 10	528	< 5	120
1828	205 238	0.01	153	170	14	5	3	262	0.01	< 10	< 10	820	< 5	241
1829	205 238	0.01	231	350	8	5	3	347	0.01	< 10	< 10	969	5	409
1830	205 238	0.01	26	190	2	< 5	< 1	597	< 0.01	< 10	< 10	157	< 5	55
1831	205 238	0.01	331	500	2	< 5	3	68	0.01	< 10	< 10	910	10	754
1832	205 238	0.01	133	300	12	< 5	3	33	0.01	< 10	< 10	227	< 5	182
1833	205 238	0.01	111	330	6	< 5	3	49	0.01	< 10	< 10	254	< 5	216
1834	205 238	0.01	127	230	10	< 5	3	61	0.01	< 10	< 10	163	< 5	201
1835	205 238	0.01	120	2230	10	< 5	3	475	0.01	< 10	< 10	241	< 5	206
1836	205 238	0.01	118	290	14	< 5	3	129	< 0.01	< 10	< 10	120	5	221
1837	205 238	0.01	120	270	6	< 5	3	65	< 0.01	< 10	< 10	141	< 5	202
1838	205 238	0.01	108	490	14	< 5	1	56	< 0.01	< 10	< 10	154	< 5	184
1839	205 238	0.01	139	750	18	< 5	2	78	< 0.01	< 10	< 10	151	< 5	196
1840	205 238	0.01	98	240	8	< 5	1	37	< 0.01	< 10	< 10	177	< 5	194
1841	205 238	0.01	110	460	14	5	2	60	0.01	< 10	< 10	364	5	506

CERTIFICATION : _____



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CERTIFICATE OF ANALYSIS A8823286

SAMPLE DESCRIPTION	PREP CODE		Al	Ag	As	Ba	Be	Bi	Ca	Cd	Co	Cr	Cu	Fe	Ga	Hg	K	La	Mg	Mn	Mb
			%	ppm	ppm	ppm	ppm	ppm	%	ppm	ppm	ppm	ppm	ppm	%	ppm	ppm	%	ppm	%	ppm
1842	205	238	0.52	< 0.2	15	410	0.5	< 2	1.21	4.0	12	155	56	1.76	< 10	< 1	0.15	10	0.11	290	24
1843	205	238	0.75	< 0.2	15	90	0.5	< 2	1.17	2.5	5	94	31	1.77	< 10	< 1	0.24	10	0.15	88	47
1844	205	238	0.69	< 0.2	10	530	0.5	6	8.31	11.0	5	94	30	1.01	< 10	< 1	0.21	< 10	0.22	100	37
1845	205	238	0.50	< 0.2	20	250	0.5	< 2	0.56	9.5	6	85	39	1.35	< 10	< 1	0.13	20	0.05	50	50
1846	205	238	1.15	< 0.2	35	100	1.0	< 2	0.22	3.5	12	130	62	2.13	< 10	2	0.30	20	0.12	108	92
1847	205	238	0.14	< 0.2	10	360	< 0.5	10	>15.00	< 0.5	8	17	37	0.96	< 10	< 1	0.04	< 10	0.42	224	12
1848	205	238	0.70	< 0.2	15	410	0.5	< 2	2.75	2.0	14	91	35	2.10	< 10	< 1	0.21	10	0.15	270	64
1849	205	238	0.12	< 0.2	< 5	260	< 0.5	10	>15.00	< 0.5	3	17	39	0.29	< 10	< 1	0.02	< 10	0.38	177	4
1850	205	238	0.71	< 0.2	15	510	0.5	< 2	6.84	1.5	6	111	19	1.17	< 10	< 1	0.23	< 10	0.23	98	51
1851	205	238	0.16	< 0.2	< 5	340	< 0.5	10	>15.00	< 0.5	5	40	13	0.50	< 10	< 1	0.03	< 10	0.42	201	7
1852	205	238	0.34	0.2	15	320	< 0.5	2	2.14	2.0	7	100	59	1.24	< 10	< 1	0.10	20	0.07	136	54
1853	205	238	0.57	< 0.2	10	500	0.5	2	3.91	1.5	5	229	43	1.01	< 10	< 1	0.17	< 10	0.13	96	44
1854	205	238	0.50	< 0.2	10	480	0.5	6	10.05	1.5	5	108	13	0.77	< 10	< 1	0.17	< 10	0.16	94	32
1855	205	238	0.62	< 0.2	10	520	0.5	2	8.57	2.0	4	67	42	0.87	< 10	1	0.20	< 10	0.21	74	50
1856	205	238	0.70	< 0.2	20	530	1.0	2	4.33	2.0	24	71	38	2.86	< 10	< 1	0.19	< 10	0.23	440	56

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CERTIFICATE OF ANALYSIS A8823286

SAMPLE DESCRIPTION	PREP CODE		Na	Ni	P	Pb	Sb	Sc	Sr	Ti	Tl	U	V	W	Zn
			%	ppm	ppm	ppm	ppm	ppm	ppm	ppm	%	ppm	ppm	ppm	ppm
1842	205	238	0.01	98	70	6	5	1	130	< 0.01	< 10	< 10	303	< 5	718
1843	205	238	0.01	112	360	10	5	2	96	0.01	< 10	< 10	360	5	369
1844	205	238	0.01	105	2780	8	5	3	568	< 0.01	< 10	< 10	344	5	601
1845	205	238	0.01	128	1820	8	5	1	76	< 0.01	< 10	< 10	238	< 5	646
1846	205	238	0.01	406	690	10	5	3	87	0.01	< 10	< 10	683	< 5	568
1847	205	238	0.01	55	190	6	5	1	485	< 0.01	< 10	< 10	133	10	171
1848	205	238	0.01	224	450	6	< 5	1	132	< 0.01	< 10	< 10	577	5	470
1849	205	238	0.01	21	250	< 2	< 5	1	952	< 0.01	10	< 10	115	< 5	72
1850	205	238	0.01	165	430	10	< 5	2	297	< 0.01	< 10	< 10	617	< 5	265
1851	205	238	0.01	32	230	< 2	< 5	< 1	489	< 0.01	< 10	< 10	133	< 5	87
1852	205	238	0.01	162	240	4	< 5	1	103	< 0.01	< 10	< 10	306	< 5	327
1853	205	238	0.01	149	650	6	5	1	197	< 0.01	< 10	< 10	596	< 5	242
1854	205	238	0.01	109	210	2	< 5	2	364	< 0.01	< 10	< 10	447	< 5	219
1855	205	238	0.01	144	280	6	< 5	2	323	< 0.01	< 10	< 10	631	< 5	271
1856	205	238	0.01	281	440	4	< 5	2	182	< 0.01	< 10	< 10	520	< 5	649

CERTIFICATION : _____

Elevation Drill contractor Logged by Total depth
 Coordinates Hole started completed Core size
 Dip Target
 Azimuth

Depth m	% Recov	Visual Log	Struct	Lithology	Alteration	Sulphide & Alteration Mineralogy						Sample Number	Assay Interval	Assay Results					
														Ni PPm	Co PPm	Mo PPm	Zn PPm	Pb PPm	As PPm
	100%		51°	48.15 - BLACK-GREY, CHERY ARGILLITE WITH 2 TO 20 cm THICK BEDS. 3 TO 10% PERVASIVE SILICATE STRINGERS.															
55	98%			56.62 - 57.05 - LIMESTONE BALL WITH SILICATE STRINGERS CONTINUED FROM CHERY ARGILLITE TAKU INTERVAL															
60			76°	57.05 - 57.38 - 3cm of SILICATE WITH CHEAT BRACIA FRAGMENTS. 57.08 - SILICEOUS SHALE WITH 5mm TO 15cm BEDS, 1% SILICATE STRINGERS															
65	100%		70° 71° 72°	65.29 - INTERVAL COMPOSED OF ZONES OF SILICATE WITH BLACK CHERY BRECCIATED FRAGMENTS OR BLACK CHERY SHALE WITH PERVASIVE SILICATE VEINLETS PARALLEL TO AND OBLIQUE TO BEDDING								1576 A	66.15 - 69.23	178	8	59	1130	1.8	30
70			69°	69.91 - 70.58 - 60cm OF LIMESTONE BALL WITH A 7mm ACICULAR WHITE SILICATE BAND. 70.58 - 76.92 - BRECCIATED CHERT								1577 A	69.23 - 72.31	96	3	41	530	1.0	20
												1578 A	72.31 - 75.4	147	3	41	986	1.2	25
75				76.92 - 78.46 - SILICEOUS SHALE WITH 1% SILICATE 78.46 - 78.56 - 10cm OF LIMESTONE BALL								1579 A	75.4 - 77.38	209	6	64	484	1.2	30
												1580 A	77.38 - 79.4	303	7	85	530	0.6	35
80				78.56 - 79.54 - CALCAREOUS SHALE 79.54 - 80.0 - LIMESTONE BALL 80.0 - 80.62 - CALCAREOUS SHALE 80.62 - 81.55 - LIMESTONE BALL 81.55 - 81.85 - CALCAREOUS SHALE								1581 A	79.4 - 81.5	19	1	6	71	0.8	5
												1582 A	81.5 - 85.35	134	2	41	265	0.8	15
85			63°	81.85 - 82.16 - LIMESTONE BALL 82.16 - 82.50 - CALCAREOUS SHALE 82.50 - 82.90 - LIMESTONE BALL 82.90 - 85.35 - SILICEOUS SHALE, BRECCIATED AT TOP AND BOTTOM, CARBONATE MATRIX															

Elevation

Drill contractor CARON

Logged by F. ANDERSON

Total depth 144.5 m (474')

Coordinates

Hole started

completed

Core size NQ

Dip - 45°

Target

Azimuth 025°

- CHERTY ARGILLITE
- LIMY SHALE
- N/S HORIZON
- BRECCIATED CHERT
- LIMESTONE BALL MARKER
- SHALE
- SILICEOUS SHALE

Depth m	% Recov	Visual Log	Struct	Lithology	Alteration	Sulphide & Alteration Mineralogy						Sample Number	Assay Interval	Assay Results			
	95%			CHERTY ARGILLITE AND SILICEOUS SHALE, LAMINATIONS 5 mm TO 10 cm													
5	99%			6.7 - BLACK LAMINATED CHERTY ARGILLITE, 5 TO 25 cm THICK BEDS													
10	100%			11.37 - CHERTY ARGILLITE BEDS, 1 TO 8 cm THICK, SEPARATED BY UP TO 5 mm THICK SHALE LAYERS. UP TO 2 mm PY LAMINAE.													
15	99%			12.09-13.23 - WHITE AND GREY SPECKLED LIMY CORE. 13.23 - CHERTY ARGILLITE BEDS 0.5 TO 1 cm THICK, THIN SHALE OR SHALE/SILICATE LAYERS.													
20	100%			20.57 - MASSIVE, LAMINATED (2 TO 8 mm) CHERT, 1% PY LAMINAE.													
25	100%			24.38 - GREY, BANDED SILICEOUS SHALE, BEDS 2 TO 45 cm. AMOUNT OF SULPHIDES LESS.													
30	100%			27.95 - BLACK, LAMINATED (1 TO 3 cm) CHERTY ARGILLITE WITH 1 cm SILICATE BANDS AND MUDSTONE BANDS.													
35																	
40	100%			41.54-42.28 - GREY/WHITE SPECKLED LIMESTONE 42.28-43.46 - BLACK CHERTY ARGILLITE BEDS 1 TO 15 cm THICK 43.46 - 43.89 - SILICEOUS SHALE, PALER THAN ARGILLITE 43.89 - CHERTY ARGILLITE, BEDS ARE 2 TO 10 cm THICK, < 1% SILICATE BANDS 47.55 - SILICEOUS SHALE, LIGHTER GREY THAN ARGILLITE AND NON-LAMINATED													

Elevation	Drill contractor	Logged by	Total depth
Coordinates	Hole started	completed	Core size
Dip	Target		
Azimuth			

Depth	% Recov	Visual Log	Struct	Lithology	Alteration	Sulphide & Alteration Mineralogy							Sample Number	Assay Interval	Assay Results				
				50.29 - BLACK, CHESTY ARGILLITE BEDS, WEAKLY LAMINATED, 0.5 TO 5cm THICK 1% PY BANDS															
55	100%			53.34 - LAMINATED ARGILLITE BEDS, 1 TO 8cm THICK, UP TO 5mm SHALE LAYERS, 1% PY BANDS															
60	100%		70°	61.57 - WEAKLY LAMINATED, BLACK, CHESTY ARGILLITE WITH 5 TO 20cm THICK BEDS. PY BANDS HAVING 10cm SPACING															
65				62.79 - LAMINATED BLACK TO GRAY BLACK ARGILLITE/SILICEOUS SHALE WITH 5 TO 10cm THICK BEDS. PY BANDS, HAVING 15cm SPACING AND SILICATE BANDS															
70	100%			69.5 - LAMINATED BLACK TO GRAY BLACK ARGILLITE/SILICEOUS SHALE WITH 1 TO 13cm THICK BEDS. NO SILICATE BANDS.															
75				78.03 - WEAKLY LAMINATED, BLACK CHERT BEDS, 5 TO 30cm THICK BEDS. MINOR SILICATE BANDS, UP TO 5mm THICK.															
80	100%			80.68-81.08 - CARBONATE MATRIX SUPPORTING BRECCIATED CHERT CLUSTS. 81.08 - WEAKLY LAMINATED CHESTY ARGILLITE WITH 2 TO 20cm THICK BEDS. <1% SILICATE LAYERS															
85				86.56 - SAME AS 81.08 BUT WITH >1% WHITE SILICATE															
90			90°																
95	100%			94.79-95.70 - LIMEY SHALE, SILICA RICH 95.40 - HARD, GREYISH CHESTY ARGILLITE WITH 2 TO 12cm THICK BEDS. <1% SILICATE AND PY LAMELLAE.															

