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YKR INTERNATIONAL RESOURCES LTD.

INDUCED POLARIZATION SURVEY AT THE AUREX PROPERTY, McQUESTEN AREA, YUKON TERRITORY

093910

Andrew Davis

Location: 63° 52' N 135° 39' W
NTS: 105M13
Mining District: Mayo, YT
Date: September 25, 1998

This report has been examined by
the Geological Evaluation Unit
under Section 53 (4) Yukon Quartz
Mining Act and is allowed as
representation work in the amount
of $ 15,700.00.

for Regional Manager, Exploration and
Geological Services for Commissioner
of Yukon Territory.

SUMMARY

An induced polarization (IP) survey was conducted on the Aurex Property for YKR International Resources Ltd. to identify sulphide rich zones associated with gold occurrence within the Keno Hills Quartzite and Hyland Group rocks. The IP survey consisted of 4.25 line-km and was conducted over 4 days in August 1998.

A promising target was located on lines 0E and 100E. Both lines contain a broad, highly chargeable and resistive anomaly near 500N. The anomaly is estimated to be 100m wide, deep, and laterally extensive, extending from (and likely further to the west of) line 0E to beyond line 100E. Chargeability and resistivity values of the anomaly are in agreement with expected values for sulphide-rich silicic or calcareous auriferous host rocks. The anomaly is on strike with known showings of the area, striking approximately E-W and occurs within the same package of rocks, making it an important economic target.

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1.0 INTRODUCTION

This report describes induced polarization (IP) survey conducted on the Aurex Property for YKR International Resources Ltd. in the Mayo Mining District, Yukon Territory. The survey was conducted to identify auriferous zones of sulphide mineralization.

2.0 LOCATION AND ACCESS

The Aurex property is centred at approximately 63° 52' N 135° 39' W (Figure 1), approximately 30 km north of Mayo. Excellent access is provided to the property by the Silver Trail Highway.

3.0 PROPERTY

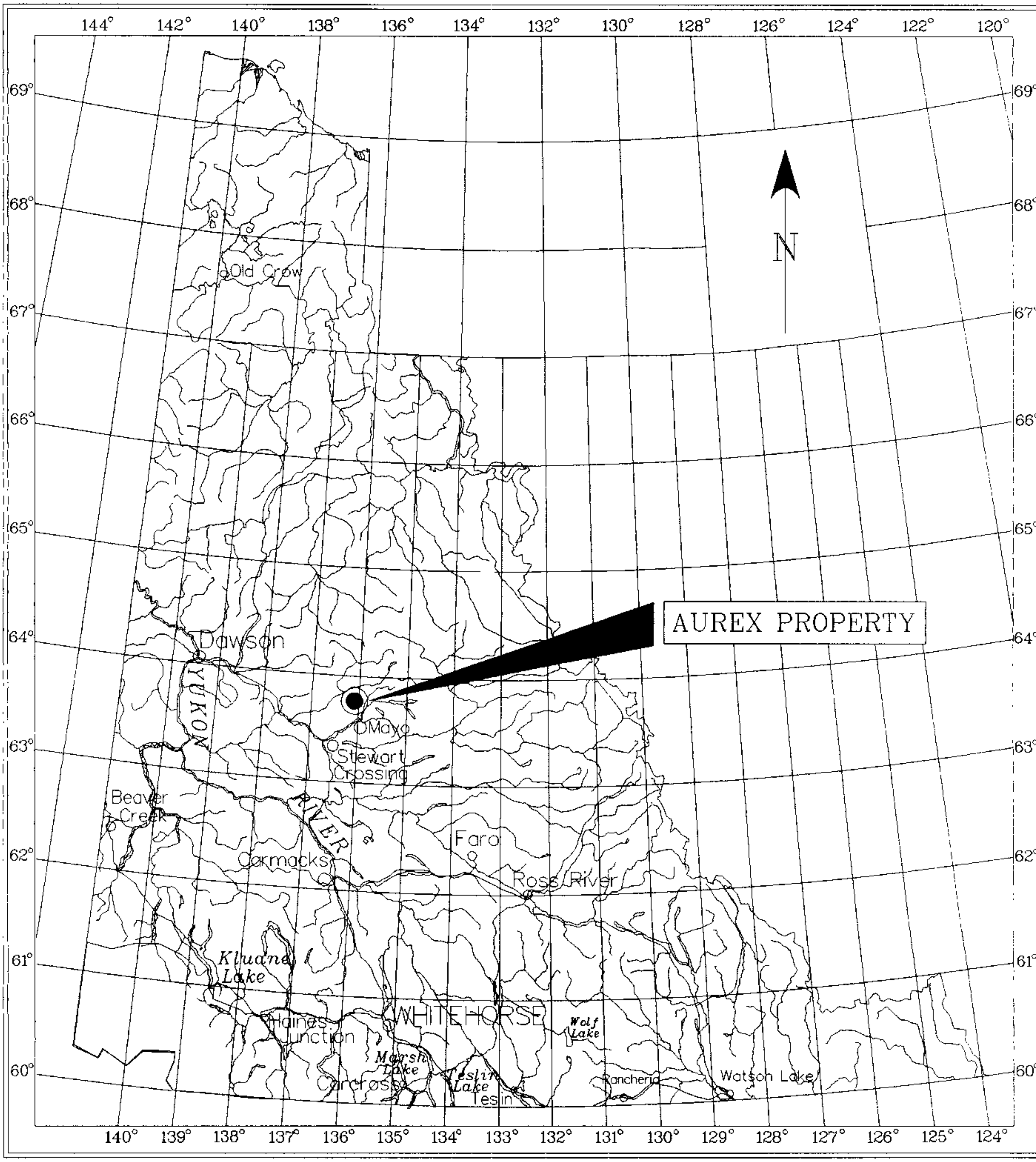
The Aurex property consist of 157 contiguous mineral claims, as listed in Table 1 below.

Table 1. Claim Data

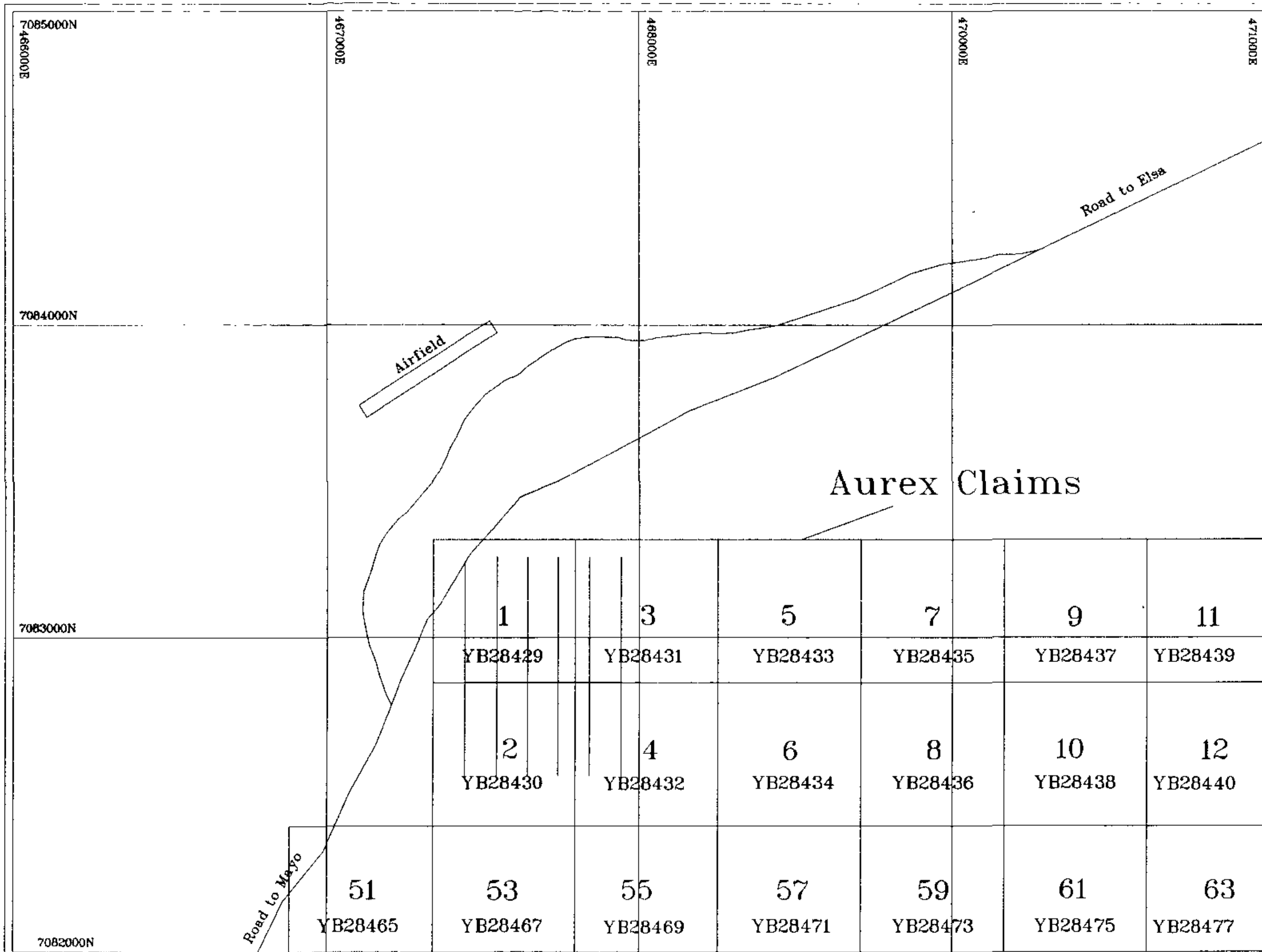
Claim Name	Grant Number	Expiry Date
Aurex 1-36	YB28429-464	21 Oct 1998
Aurex 51-86	YB28465-500	21 Oct 1998
Aurex 87-113	YB29367-392	21 Oct 1998
Aurex 114-171	YB29669-726	10 Mar 1999

4.0 GRID

The location of the survey grid is shown in Figure 2 relative to the boundaries of the property. Survey lines were cut and straight-chained (not slope corrected) with stations at 50m intervals with half-length, tagged survey lathe. Intermediate stations at 25m were marked with labeled flagging. A total of 4.25 line-km were covered during the IP survey.



YKR INTERNATIONAL RES.	AUREX PROPERTY
PROPERTY LOCATION	MINING DISTRICT: MAYO
	NTS:105 M/13 SCALE: 1: 6 000 000
	DRAWN BY CP
AMEROK GEOSCIENCES LTD.	DATE: 13 AUG 98 FIGURE: 1



YKR INTERNATIONAL RESOURCES LTD.

AUREX PROPERTY

(105 M/13)

IP SURVEY
GRID LOCATION

Figure 2.

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5.0 PROPERTY GEOLOGY

The geology of the Aurex Property has been well described by Davidson (1995) and the following summary is drawn from his report.

Rocks on the Aurex claims are of the Hyland Group, composed of sandstones, conglomerates and phyllites. On the Aurex property these rocks are mainly phyllitic quartz sericite schists which are in contact with the Keno Hill Quartzites to the north. Rocks follow the regional strike of 90° and dip 20-35° south. The contact is the Robert Service Thrust Fault, running approximately E-W.

Hyland Group rocks along the fault are highly deformed sedimentary units consisting of conglomerates, psammites, and phyllitic marbles. This package immediately above and below the Robert Service Thrust Fault has been interpreted as a structurally imbricated, isoclinally folded and faulted package.

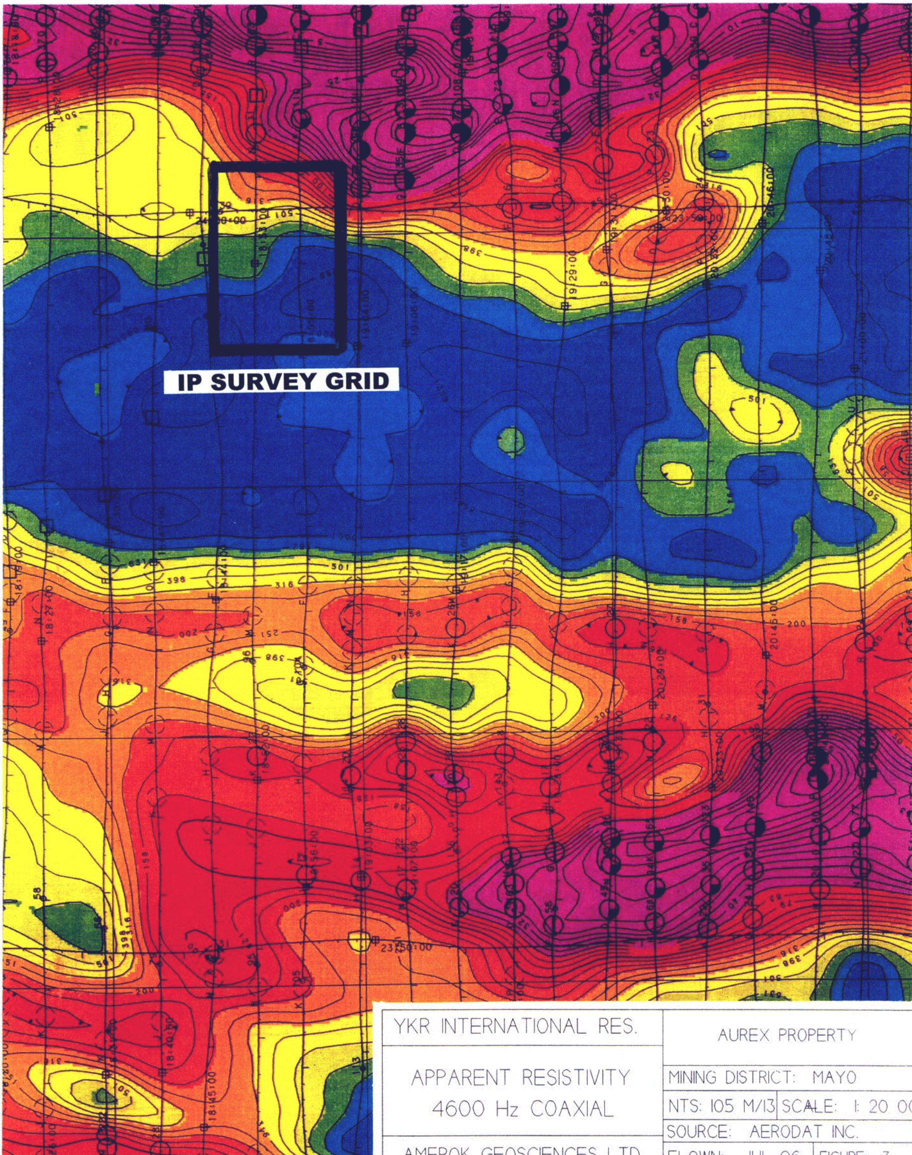
Also found on the claims are calc-silicate skarn/hornfels containing 1-5% very fine grained disseminated pyrrhotite/arsenopyrite and minor pyrite in contact with Hyland Group quartz sericite schist and phyllite.

The intrusion causing the contact metamorphism has not been found in outcrop on the Aurex claims; intrusive rocks in the area are restricted to granitic and aplitic dykes and sills located along the northern property boundary.

Skarns appear to have formed from the contact metamorphism of limestone. Some are derived from calcareous schist and calcareous quartzite. Gold values are directly related to the abundance of sulphides, and the amount of sulphides is much higher in skarns than in quartz sericite schist and phyllite.

Economic targets are zones containing high concentrations of disseminated sulphides, such as the skarns described above. These rocks may have anomalous high chargeabilities, and are expected to be somewhat resistive.

Airborne EM and Magnetics surveys were flown in July 1996. Figure 3, included in this report is a reproduction of the apparent resistivity contour map for 4600 Hz coaxial EM from the 1996 survey conducted by Aerodat Inc.. Warm colours represent conductive units, while cooler colours represent resistive units. The IP survey grid boundary described in this report has been superimposed upon this map.



YKR INTERNATIONAL RES.	AUREX PROPERTY	
APPARENT RESISTIVITY 4600 Hz COAXIAL	MINING DISTRICT: MAYO	
	NTS: 105 M/13	SCALE: 1: 20 000
	SOURCE: AERODAT INC.	
AMEROK GEOSCIENCES LTD.	FLOWN: JUL 96	FIGURE: 3

6.0 PERSONNEL AND EQUIPMENT

The surveys were conducted by the following staff from Amerok Geosciences Ltd.:

Person	Position	Address
Andrew Davis	Crew chief	#8-201 Hanson St. Whitehorse YT Y1A 1Y3
Dan Hall	Technician	1 Bates Crescent Whitehorse YT Y1A 4T8
Gary Lee	Technician	Box 5348 Whitehorse YT Y1A 4Z2
Ron Stack	Technician	Site 20 Comp 86 RR1 Whitehorse YT Y1A 4Z6

The crew were equipped with the following instruments and equipment:

IP Transmitter:	Phoenix IPT-1 mated with 2.5 KW motor generator. Maximum output voltage: 1500 V / maximum output power approximately 2.2 KW. Spare Phoenix IPT-1 provided.
IP Receiver:	IRIS IP-10 digital 10-channel IP time domain receiver
Data processing:	486 laptop and HP-680C colour printer. Data processing with Geopak IPSECT software and proprietary data conversion software.
Other equipment:	6-conductor 50 m IP cables, stainless steel electrodes, 4 km wire, winders, VHF radios, F350 truck, line cutting equipment.

The crew spent a total of 4 days on the Property.

7.0 SURVEY SPECIFICATIONS

The IP surveys were conducted according to the following specifications:

Array:	Dipole-dipole
Dipole spacing:	25 m
Separations read:	n=1 to 6
Signal:	0.125 Hz / 50% duty cycle / reversing polarity
Receiver synch:	synchronization using n=1 dipole signal.
Signal sampling:	20 windows, Cole-Cole semi-logarithmic sampling over 2 s.
Measurements:	Vp - primary voltage prior to shutoff M_n - nth time slice chargeability (n=1 to 20) Mt - total chargeability Ro - apparent resistivity Sp - self potential Rs - electrode resistance
Noise threshold:	Standard deviation in Mt kept to $\leq$ 5 ms where possible. In the event that this was not possible, readings were repeated several times to ensure repeatability.
Stacking:	minimum 15 times, maximum 30 times for a single reading.

8.0 DATA

Copies of the corrected IP data in Geosoft format (.gsf) files are appended to this report on 3.5" diskettes. IP pseudosections are contained in Appendix B. Significant chargeability and resistivity anomalies are indicated by horizontal lines above the respective pseudosections, indicating the horizontal location of the tops of the source bodies. The anomalies are assigned a letter designator, from left to right across the pseudosections.

A small number of readings were deleted from the data set because they contained high apparent errors and readings were not repeatable in the field. These occurred at sites of very conductive ground, where very low primary voltages were recorded (< 1mV). Default (no data) values were inserted for these points and the software contoured around them.

9.0 IP INTERPRETATION PROCEDURES

The IP data was interpreted using a procedure sketched schematically in Figure 4. The numbers in the flow chart refer to information sheets in the company interpretation manual. Key features of the responses mentioned in these sheets are summarized below and are drawn from summaries and investigations by Telford *et. al.* (1990), Sumner (1985), Hanneson (1990), Hohmann (1990), and Coggon (1973).

The source field for the survey was a grounded current dipole with a spacing of 25m near a reading array of 25m dipoles. The receiving dipoles were separated from the current dipole by a variable spacing of 1 to 6 times the 25m dipole spacing. The source field from a grounded current dipole is symmetric about the midpoint of the pair and drops off dramatically with distance. There are no effects in the pseudosections which are primarily due to the source field.

9.1 Overburden responses

Overburden responses in a dipole-dipole survey show up as a flat-lying layer in the pseudosection. The depth to the boundary between layers of different resistivity or chargeability can be estimated as 1.5 to 2.0 times separation at which the gradient between the two layers is the greatest. This inevitably leads to an overestimation if the dipole spacing is large relative to the thickness of the layer. In some cases, the overburden response is not visible as a separate resistivity anomaly but is apparent as a flat lying layer of lower chargeability - usually only down to n=1. This is attributed to oxidation or leaching of chargeable minerals or graphite from bedrock near the surface or to the absence of chargeable minerals in overburden. Features which appear to be limited in depth extent are referred to as depth limited in this report.

9.2 Two dimensional versus three dimensional responses

Responses were interpreted as two dimensional (i.e. extending along strike to some extent) unless otherwise stated. If a target is in fact three dimensional and is interpreted as being two dimensional, the contrast between the host and target properties will be underestimated.

9.3 Apex location and width

Targets which are less than one half a dipole spacing (i.e. 12.5 m) will produce single slash responses. The apparent dip of the single slash response *does not* indicate the dip of the feature but merely indicates which electrode was closer to the source. A thin target may also produce a symmetric two-slash response if it is centred at an electrode site. The width of the source body was considered to be definitely less than 12.5 m if a single slash anomaly was encountered and to be at least 12.5 m if a symmetric response were encountered. It is difficult to discriminate between a 12.5 and 25 m wide target response if the response is symmetric and the author has chosen to err on the wide side. If the response at the shortest separation is wider than one dipole, this is an indication that the source body is also wider than one dipole. The width of the response at the shortest separation was used to determine the width of the source body in most case; in certain circumstances, however, the response was compared with model responses to determine the source width. The solid lines in the pseudosections and on the anomaly maps show the horizontal location of the top of the source bodies and the apparent width of the target. The error in apex location is conservatively estimated ± 1 dipole (25 m).

9.4 Depth to top

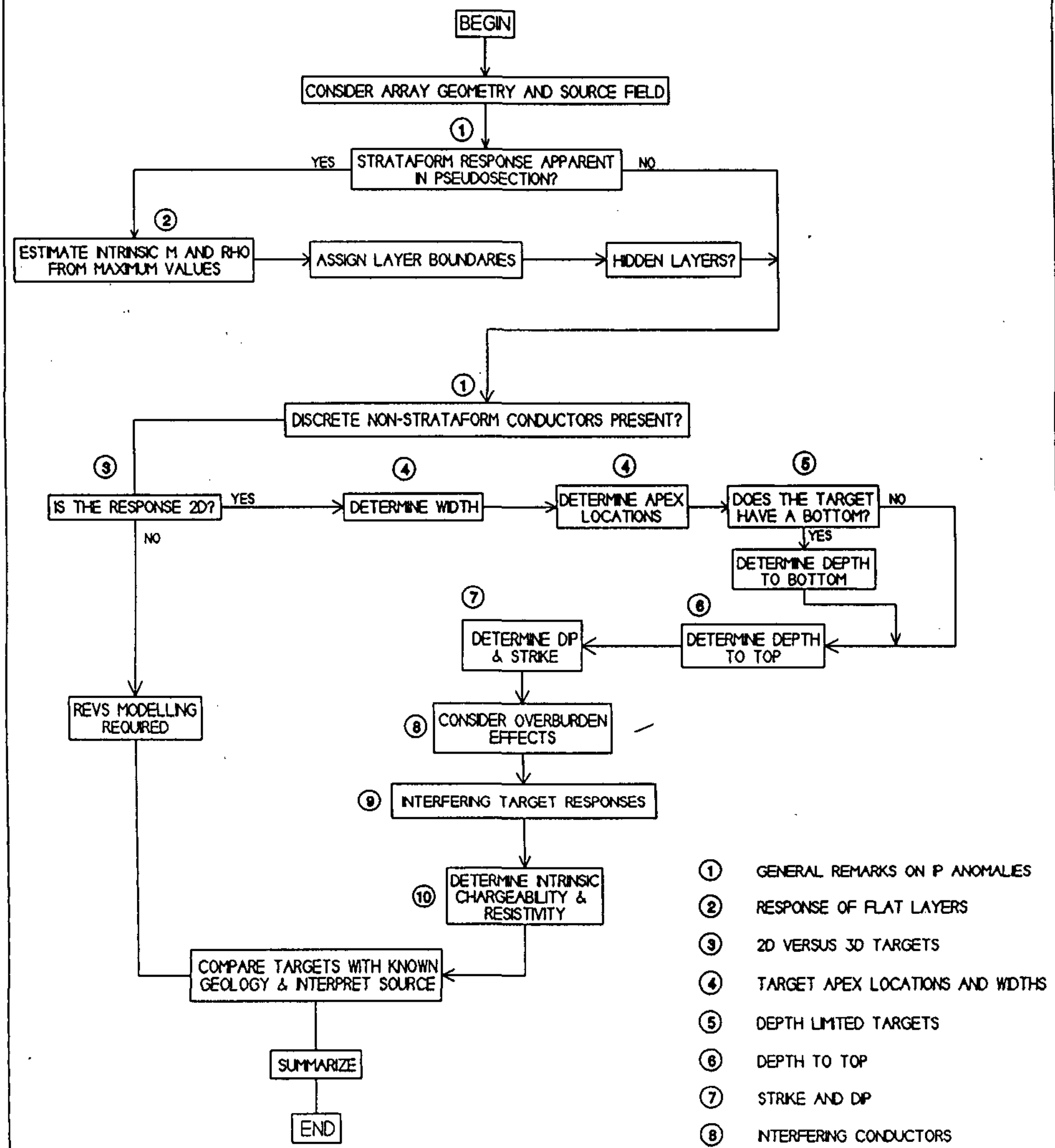
The depth to the top of a source body is generally indicated by the shortest separation at which the response if visible. Thus a target at a depth of 25 m would be expected to produce some response at n=1 but a target with a top at 50 m would generally not be visible at n=1.

9.5 Dip direction

The dip direction and dip of a source body are difficult to estimate with dipole-dipole data. Dip must be estimated using both the resistivity and chargeability data because the dip direction will be different depending upon whether the chargeable target is more or less resistive than the host. If the target is more resistive than the host, the dip in the chargeability pseudosection will be in the same direction as the target. If the target is less resistive than the host, the apparent dip will be opposite the true dip. At a dipping contact, the steepest gradient in a resistivity section dips in the opposite direction to the true dip of the contact. Estimates of dip direction are difficult or impossible to make where targets of alternating resistivity are adjacent to each other.

9.6 Target resistivity and chargeability

Estimates of true or intrinsic target chargeability and resistivity can be made once the



IP interpretation flow-chart.

interpreter has some idea of the target dimensions. In general, for a given resistivity and chargeability contrast, the target response will decrease as the target dimensions decrease. In addition, the amplitude of the chargeability contrast will be affected by the resistivity contrast. Targets which are very resistive or very conductive will show much lower apparent chargeabilities relative to true chargeability.

A three dimensional target (e.g. a sphere) will produce an anomaly with a maximum apparent chargeability which is at best 30% of the true chargeability response. If the target is two dimensional, the maximum apparent chargeability is 50% of the true chargeability unless the target is thin in which case the maximum apparent chargeability will be up to 40% of the true intrinsic chargeability.

Estimates of the true chargeability and, to a lesser extent, resistivity can be used to estimate the probable source of an anomaly. Chargeabilities are largely determined by the bulk concentration of chargeable minerals such as sulphides or graphite. It is difficult to discriminate between the two. Rules of thumb cited by Sumner (1976) and Hohman (1990) relate chargeable mineral content to recorded IP parameters:

$$1\% sulphides \cong 3\% PFE \cong 20 ms \cong 10 mrad$$

There are wide variances between the sulphide content predicted by these relations and the actual sulphide content. These arise from the effect of electrical resistivity on measured chargeability. Rocks which are highly resistive (few current paths) or very conductive (too many current paths) will exhibit lower than predicted apparent chargeability and estimates of chargeable mineral content will err on the low side. In addition, estimates of sulphide content based on chargeability must account for background chargeability due to clay minerals.

10.0 IP SURVEY RESULTS

Disseminated sulphide or graphite mineralization produce high chargeability responses. If the concentrations are very high, they may significantly lower rock apparent resistivity as well. Silicification or potassic alteration can raise bedrock apparent resistivity by blocking current paths, even if significant sulphide concentrations are present. In effect, chargeability is governed by the surface area of chargeable minerals present and the apparent resistivity is controlled by the extent to which conductive minerals are electrically connected. Consequently, high chargeability anomalies indicate the presence of disseminated sulphides or graphite. Low apparent resistivities may indicate the presence of clay alteration, significant sulphide concentrations or graphite which enhance the overall electrical conductivity of the bedrock. Graphite is known to occur in this environment, and many of the anomalies located are likely due to its presence. Characteristic responses of graphitic rocks are high chargeability values coupled with very low resistivity values.

All lines show a permafrost layer in both the chargeability and resistivity pseudosections. In the chargeability section this appears as a zone of suppressed chargeability at the n=1, or first separation, layer of the section; in the resistivity pseudosection, permafrost appears as increased resistivity, again at the n=1 layer. Both features are flat-lying.

A geologic contact is also present in all sections. The contact divides less resistive rocks to the north and more resistive rocks to the south and is marked by what appears to be a graphite containing unit. The contact strikes approximately E-W and is located at approximately 750N along the lines. Descriptions of the data, anomalies and interpreted source parameters for each line follow.

Line 0E

The contact is located at 775N. High background chargeabilities are apparent throughout this line, and may indicate the presence of disseminated sulphide mineralization or graphite within the host rocks.

Anomaly A is a strong chargeability anomaly with subtle expression in the resistivity pseudosection. The broad pantleg anomaly extends from 400N to 550N and shows chargeabilities of up to 75mV/V. The apex of the anomaly is at 512N. Depth to top is estimated to be 75m, with no apparent bottom (deeper than 150m). Dip is difficult to estimate; however, the shape of the anomaly and pattern of resistivity suggest a dip to the south. The presence of a similar anomaly on line 100E may indicate a 2D source, striking approximately E-W. High chargeabilities may represent the presence of sulphide mineralization at depth.

Line 100E

The contact is located at approximately 750N. There are three discrete anomalies within the chargeability pseudosection, each with corresponding anomalies in the resistivity section. A fourth anomaly is apparent in the resistivity section, with no corresponding chargeability signature. High background chargeabilities are apparent throughout this line, and may indicate the presence of disseminated sulphide mineralization in the host rocks.

Chargeability anomaly A and resistivity anomaly A occur beneath the interval 450N to

550N, and are centred at 500N and 467N respectively. Depth to top is estimated to be 50m with no apparent bottom (deeper than 150m). Chargeabilities increase with depth, suggesting a deep source. Dip is impossible to estimate. The characteristics of this anomaly are very similar to anomaly A of line 0E, and this may be the 2D extension of a single E-W striking source. High chargeabilities may represent the presence of sulphide mineralization at depth.

Chargeability anomaly B, and resistivity anomaly C are centred at 725N. The triangular pattern, especially visible in the resistivity section suggests a thin, steeply dipping layer of very conductive and chargeable rock; likely graphite. The width of the source is estimated to be less than 12.5m .

Chargeability anomaly C and resistivity anomaly D are located between 787N and 800N. The single slash anomaly indicates a thin, conductive and chargeable source estimated to be less than 12.5m wide. Dip is impossible to estimate. High chargeabilities and low resistivities are indicative of graphite. Similar patterns are found on lines to the E, suggesting an E-W striking source.

Resistivity anomaly B is a single slash moderate resistive high with apex at 512N. This pattern is representative of a thin, resistive source less than 25 m wide, extending from near surface to below the depth of penetration (deeper than 150m).

<u>Line 200E</u>

The contact is located at 737N. One chargeability anomaly and two resistivity anomalies are present in these pseudosections. High background chargeabilities south of the contact may indicate the presence of sulphide mineralization in the host rocks.

Chargeability anomaly A and resistivity anomaly A are located with apices at 775N and 737N respectively. Chargeability anomaly A is a moderate, deep source, chargeability high, occurring within a complex system of resistivity lows. Resistivities of under 100 ohm-m present here, are indicative of a graphitic source. The pattern of the resistivity anomaly suggests a thin, steeply dipping, conductive source with estimated width of 12.5m or less located between 725N and 737N. A similar anomaly pattern is visible on neighbouring lines, suggesting an E-W striking thin bed.

Resistivity anomaly B occurs at 825N to beyond 925N. Very low resistivities are likely due to a graphite source. Depth to top is estimated to be between 50m and 100m.

Line 300E

The contact is located at 725N.

Chargeability anomaly A is located between 525N and 600N and is characterized by enhanced chargeabilities in otherwise low-chargeability ground. The broad extent indicates a wide source. Depth to top is estimated from the chargeability section to be 75m to 100m. Dip of this feature is impossible to estimate.

Chargeability anomaly B and resistivity anomaly B are centred at 750N, on or immediately north of the contact. The pattern of this anomaly suggests a thin conductive bed. Asymmetry within the chargeability section may indicate a dip to the south. The source is likely graphite.

Resistivity anomaly B neighbours the contact, and appears to be a very conductive mass of rock suggesting the presence of sheet like-graphite in this transition zone. Depth to top is estimated to be 50m to 100m. Higher resistivity values near the surface show the presence of more resistive rock or permafrost above this feature.

Line 400E

There are four discrete chargeability anomalies and four corresponding resistivity anomalies. The pattern of chargeabilities and resistivities is more complex than in this section than in the previous lines; however, the same general pattern of more resistive rocks to the south and less resistive rocks to the north is repeated. The contact is estimated to be near 675N.

Chargeability anomaly A is a highly chargeable anomaly at 525N, from a deep, broad source extending across four dipole spacings. Depth to top is estimated to be 100m. Resistivity anomaly A is a corresponding resistivity high. Complex patterns in the sections - negative IP responses and a high-low resistivity coupling - may indicate that the source is offline, or more likely, a thin, very conductive source at depth, likely graphite.

Chargeability anomaly B, and resistivity anomaly B are located with apex at 575N. The general pattern of this anomaly indicates a source greater than 25m wide. Depth to top is estimated to be 25 to 50m. Depth to bottom is estimated at 100m. Geometry of the anomaly suggests a south dipping slab. The source may be graphite or sulphide mineralization.

Chargeability anomaly C, is a single slash chargeability high located at 700N.

Resistivity anomaly C is a corresponding resistivity low at this slash. The pattern of these anomalies indicate a thin, chargeable, conductor - possibly graphite or sulphide mineralization.

Chargeability anomaly D occurs at the end of the section, beginning at 825N and extending past the end of the line. Very low resistivity values in corresponding resistivity anomaly C are indicative of a graphitic source. Negative chargeabilities are present, which are sometimes associated with graphite. Depth to source is approximately 75m. The geometry of the anomaly suggests a flat lying source with limited depth. A similar, but stronger anomaly is noticed at the north end of line 500E, which may indicate an E-W striking feature.

Line 500E

The approximate location of the contact is 750N.

Chargeability anomaly A is a very broad, subtle anomaly, extending from 550N to 650N, and centred at 587N. A moderate resistivity low corresponds to the chargeable high. The shape of the anomaly indicates a wide, possibly depth limited source. Estimated depth to top is 50 to 75m. Depth bottom is estimated to be 150m. The response suggests a flat lying or moderately dipping source, possibly graphite.

Chargeability anomaly B and resistivity anomaly B are located at depth and centred at 787N. Both anomalies extend over several dipole spacings. The complex pattern of high positive and negative chargeabilities combined with very low resistivity values, suggests a thin interbed of conductive chargeable mineralization; most likely graphite. Depth to top is estimated to be 100m. This feature appears to be south-dipping.

Chargeability anomaly C and resistivity anomaly C extend from approximately 850N to beyond the end of the line and are centred at 950N. The pattern and magnitude of the anomalies suggest the presence of graphite. Depth to top is estimated to be 25 to 50m. Dip is impossible to estimate. This anomaly appears to be a 2D extension of a similar anomaly at the end of Line 400E, indicating an E-W striking feature.

11.0 DISCUSSION

Gold values on the Aurex property are directly related to the abundance of sulphides. The amount of sulphides is much higher in skarns than in quartz sericite schist and phyllite. Silicic and calcareous source rocks containing high concentrations of sulphide mineralization would be expected to present excellent IP targets, exhibiting both high chargeabilities and moderate to high resistivity values.

Stratigraphically, known showings of the region occur within conformable beds of the Hyland Group and Keno Hill Quartzites, in close proximity to the Robert Service Thrust Fault. Regional strike is approximately E-W and dip is to the south.

A promising target is seen in lines 0E and 100E. Both contain a broad, highly chargeable and somewhat resistive anomaly centred near 500N. Chargeabilities in excess of 40mV/V have been recorded. Strike of this feature is approximately E-W, in agreement with regional trends and known showings in the area. Dip is difficult to estimate. The anomaly extends more than 200m from 0E across 100E, but becomes less intense at line 200E. Width is estimated to be 100m to 125m. Depth to top is estimated to be 50m to 75m. The anomaly may extend to the west of the survey grid.

This highly chargeable horizon is considered a viable economic target due to its intrinsic chargeability and resistivity values, its apparent location within the Hyland Group rocks, stratigraphically above the graphitic schist of the Keno Hills Group, its width, and lateral dimensions. The signature of this anomaly fits a sulphide rich, moderately resistive, conformable target.

A contact between highly conductive rocks in the north of the grid and more resistive rocks to the south has been identified on all lines. The marker bed of this contact is suspected to be a thin graphite bed. The contact strikes approximately E-W and is located between 700N and 775N on all lines. Results of the survey show that high chargeability anomalies north of the contact are most likely from graphitic sources within the graphite schist, while those anomalies south of the contact are less likely from graphite sources but instead from sulphide mineralization. This interpretation is based upon study of local geology and comparison of chargeability and resistivity pseudosections.

The IP survey results differentiate three distinct units within the grid based upon geophysical responses. North of the contact at approximately 750N is a graphite bearing unit, presumed to be the graphitic schist of the Keno Hills Group. South of the contact, west of line 400E is a second unit that is moderately resistive and chargeable containing the very chargeable horizon described above. Higher than average background chargeabilities may indicate the presence of disseminated sulphides in this unit. The third unit is not as well defined but is seen south of the contact and east of line 300E. It exhibits both moderate chargeability and moderate to low resistivity, with more complex local features throughout. Sites of high chargeability and low resistivity within this third unit may be due to the presence of graphite.

12.0 CONCLUSIONS

The results of the induced polarization survey conducted on the Aurex Property Grid suggest the following conclusions:

a. An E-W striking contact at approximately 750N separates graphitic rocks to the north from less graphitic rocks to the south.

b. Rocks south of the contact can be grouped into two distinct units on the basis of IP response. The apparent contact is located between lines 300E and 400E.

c. Rocks south of the contact at 750N and west of 400E show promise as auriferous host rocks, exhibiting high background chargeabilities and high resistivities and contain one broad and chargeable horizon.

d. The broad and chargeable horizon strikes E-W, extends at least 200m from line 0E to beyond 100E, and fits the target model of a sulphide-rich silicic or calcareous auriferous host rock.

13.0 RECOMMENDATIONS

The following recommendations are made based on the conclusions of this work:

a. A ground total magnetic field survey should be conducted on the grid to further evaluate the potential of the auriferous host rocks. Gold has been associated with pyrrhotite, and a magnetic survey may help to discriminate sulphide sources from graphite sources.

b. Geochemical surveying and prospecting should be conducted on the anomaly at 500N of lines 0E and 100E.

Respectfully submitted,
AMEROK GEOSCIENCES LTD.

Andrew Davis
Geophysicist

References Cited

Coggon, J.H. (1973). A comparison of IP electrode arrays. Geophysics Vol. 38, No 4. pp 737-761.

Davidson, G.S. (1995). Summary Report on the Aurex Property.

Hanneson, J.E. (1990) A model for interpreting IP/resistivity data from areas of steep dip and thin overburden. in: Fink, J.B. *et. al.* (ed.) Induced polarization: applications and case histories. Tulsa: Society of Exploration Geophysicists.

Hohmann, G.W. (1990) Three-dimensional IP models. in: Fink, J.B. *et. al.* (ed.) Induced polarization: applications and case histories. Tulsa: Society of Exploration Geophysicists.

Johnson, I.M. (1990) Spectral IP parameters derived from time domain measurements. in: Fink, J.B. *et. al.* (ed.) Induced polarization: applications and case histories. Tulsa: Society of Exploration Geophysicists.

Sumner, J.S. (1976) Principles of Induced Polarization for Geophysical Exploration. New York: Elsevier.

Telford, W.M., L.P. Geldart and R.E. Sheriff (1990) Applied Geophysics (2nd Edition) New York: Cambridge University Press.

APPENDIX A. CERTIFICATE

I, Andrew Richard Davis, with residence address in Whitehorse, Yukon Territory do hereby certify that:

1. I have completed my fourth year of studies in Geophysics.

2. I have been actively involved in mineral exploration the Northern Cordillera since June 1998.

3. I conducted the IP survey and interpreted the results with the assistance of Mike Power, P. Geoph. .

4. I have no interest, direct or indirect, nor do I hope to receive any interest, direct or indirect, in YKR International Resources Ltd. or any of its properties.

Dated this 23rd day of September, 1998 in Whitehorse, Yukon.

Respectfully Submitted,

Andrew Davis

APPENDIX B. PSEUDOSECTIONS

APPENDIX C. STATEMENT OF COSTS

Mobilization / demobilization	$2,300.00
IP Survey and line cutting	$9,622.00
Report	$1,900.00
GST	$967.54
Total	**$14,789.54**

LINE : 0 E

INDUCED POLARIZATION SURVEY

DIPOLE-DIPOLE ARRAY

A NA A

DEPTH POINT

N = 1, 2, 3, 4, ...

"A" SPACING = 25.0 METRES

YKR INTERNATIONAL RES.

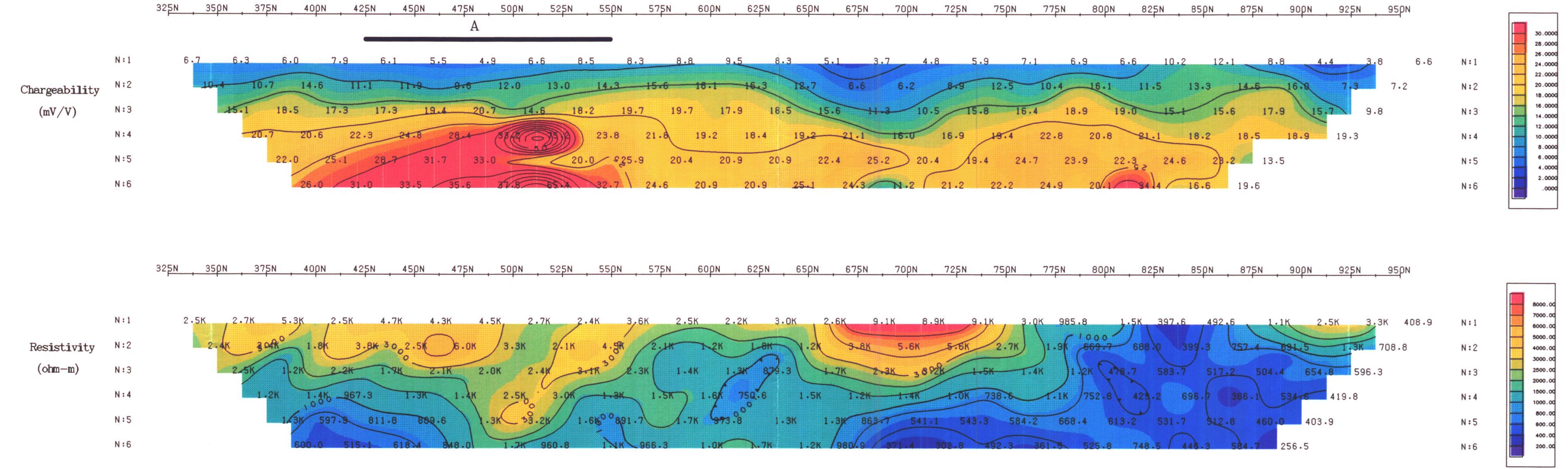
AUREX PROPERTY

NTS 105 M/13

DATE : 12 AUG 98	REF : 98-22

SCALE = 1: 1600

AMEROK GEOSCIENCES LTD.

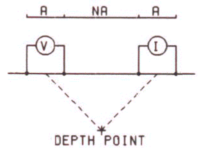


Dwh ①

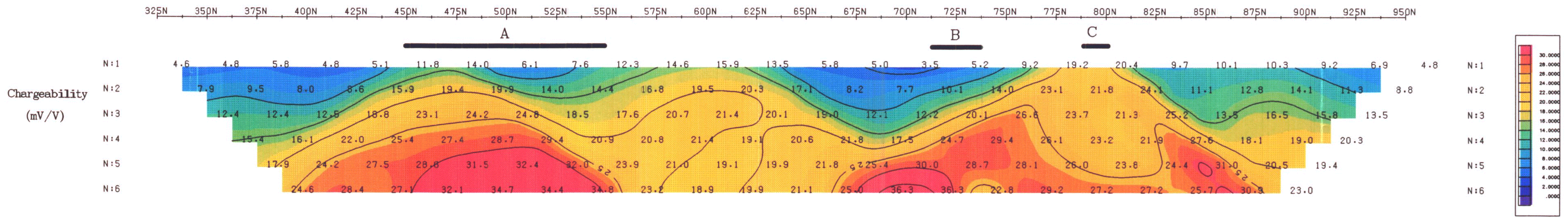
LINE : 100 E

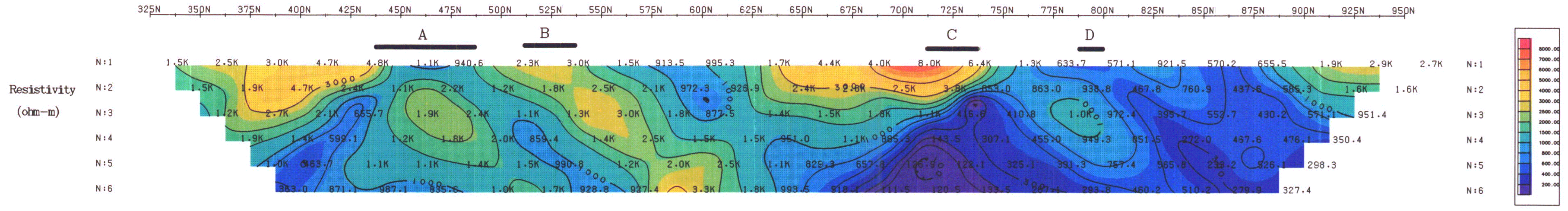
INDUCED POLARIZATION SURVEY

DIPOLE-DIPOLE ARRAY



N = 1, 2, 3, 4, ...
"A" SPACING = 25.0 METRES





YKR INTERNATIONAL RES.

AUREX PROPERTY

NTS 105 M/13

DATE : 12 AUG 98	REF : 98-22

SCALE = 1: 1600

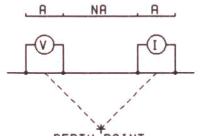
AMEROK GEOSCIENCES LTD.

Dwg (2)

LINE : 200 E

INDUCED POLARIZATION SURVEY

DIPOLE-DIPOLE ARRAY



DEPTH POINT

N = 1, 2, 3, 4, ...

"A" SPACING = 25.0 METRES

YKR INTERNATIONAL RES.

AUREX PROPERTY

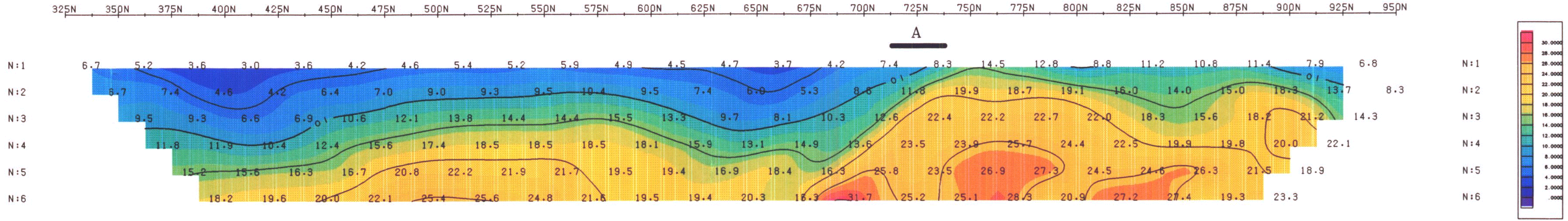
NTS 105 M/13

DATE : 12 AUG 98 | REF : 98-22

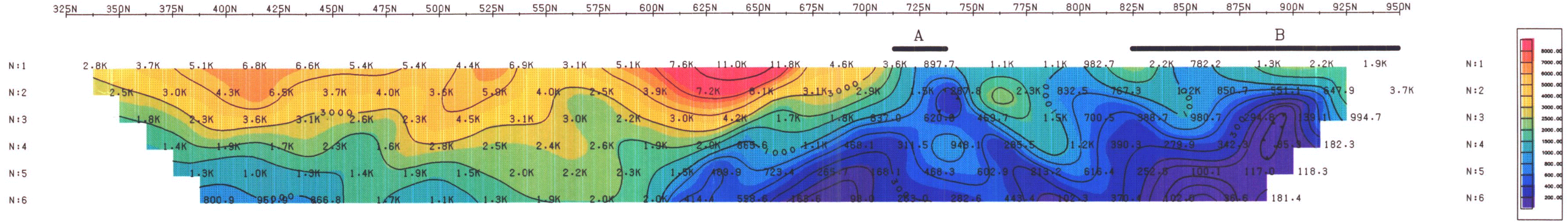
SCALE = 1: 1600

AMEROK GEOSCIENCES LTD.

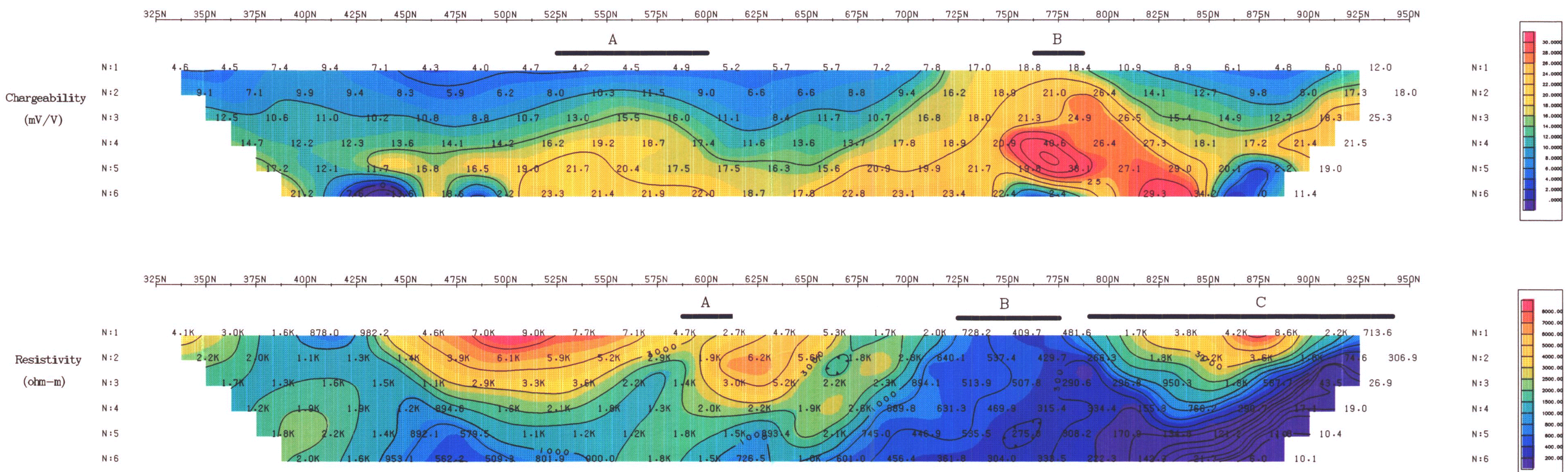
Chargeability (mV/V)



Resistivity (ohm-m)



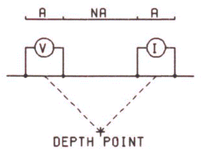
DWG (3)



LINE : 300 E

INDUCED POLARIZATION SURVEY

DIPOLE-DIPOLE ARRAY



N = 1, 2, 3, 4, ...
"A" SPACING = 25.0 METRES

YKR INTERNATIONAL RES.

AUREX PROPERTY

NTS 105 M/13

DATE : 12 AUG 98	REF : 98-22

SCALE = 1: 1600

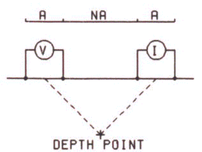
AMEROK GEOSCIENCES LTD.

Dwg 4

LINE : 400 E

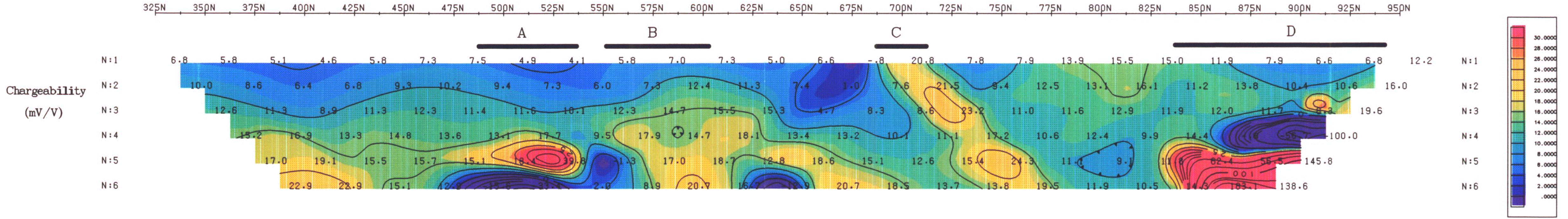
INDUCED POLARIZATION SURVEY

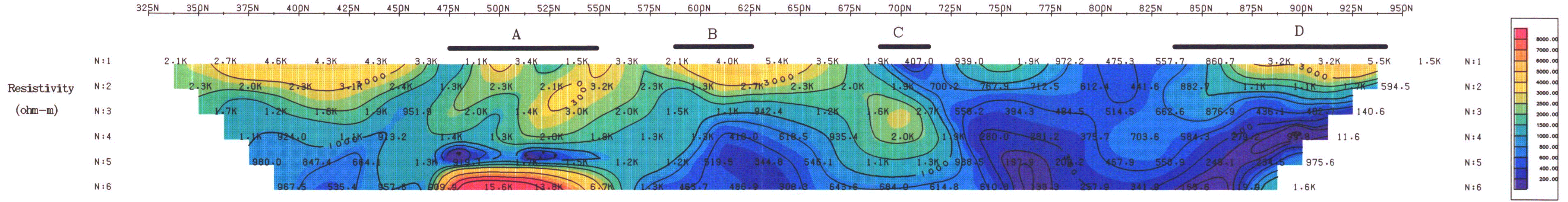
DIPOLE-DIPOLE ARRAY



N = 1, 2, 3, 4, ...

"A" SPACING = 25.0 METRES





YKR INTERNATIONAL RES.

AUREX PROPERTY

NTS 105 M/13

DATE : 12 AUG 98	REF : 98-22

SCALE = 1: 1600

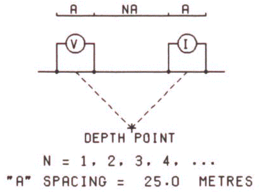
AMEROK GEOSCIENCES LTD.

DWG (5)

LINE : 500 E

INDUCED POLARIZATION SURVEY

DIPOLE-DIPOLE ARRAY



YKR INTERNATIONAL RES.

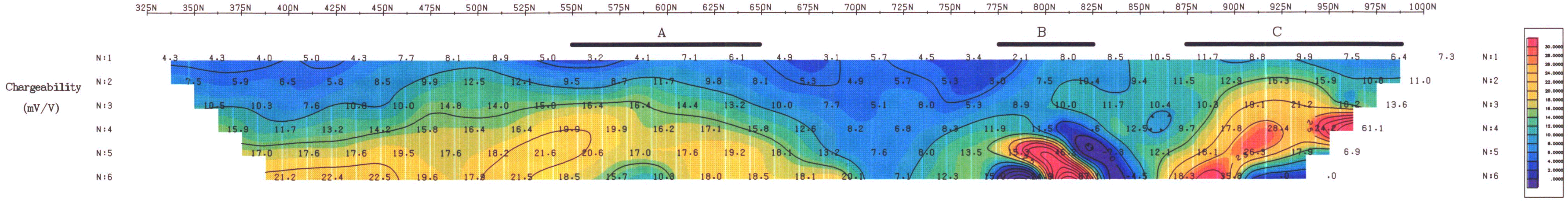
AUREX PROPERTY

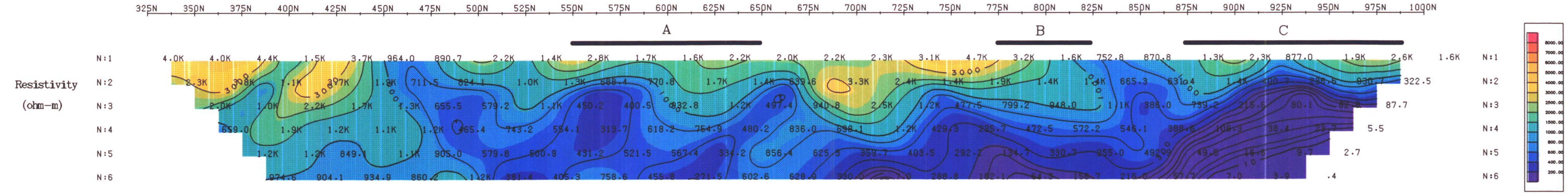
NTS 105 M/13

DATE : 12 AUG 98 | REF : 98-22

SCALE = 1: 1600

AMEROK GEOSCIENCES LTD.





Dwg 6