

GEOCHEMICAL

REPORT

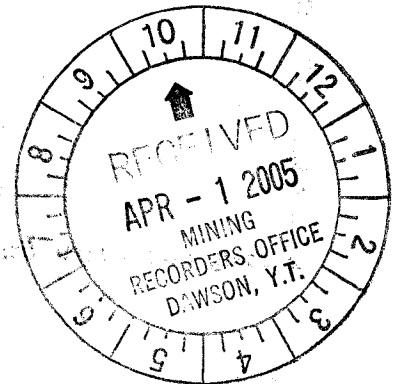
SIMBA 41- 214 CLAIMS

GRANT # YC21872 – YC30422

NTS # 116 C \ 09

LAT: 64' 03' N

LONG: 140' 25' W



DAWSON MINING DISTRICT

095283

AUTHOR OF REPORT SHAWN RYAN

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K. Perry

Mining Recorder
Dawson City Mining District

TABLE OF CONTENT

SUMMARY	p.3
1.0 INTRODUCTION	p.3
2.0 LOCATION AND ACCESS	p.3
3.0 PROPERTY DESCRIPTION	p.3
4.0 PHYSIOGRAPHY	p.3
5.0 REGIONAL AND PROPERTY GEOLOGY	p.4
5.1 REGIONAL GEOLOGY	p.4
5.2 PROPERTY GEOLOGY	p.4-5
6.0 WORK PROGRAM / METHODS	p.6
6.1 REGIONAL SILT SURVEY	p.6
7.0 INTERPRETATION	p.6
7.1 REGIONAL SILT SURVEY	p.6
8.0 RECOMMENDATION	p.7
9.0 REFERENCES CITED	p.7
10.0 QUALIFICATION	p.7
11.0 COST	p.8
Silt Gold Geochem Maps	Figure 1
Silt Copper Geochem Maps	Figure 2
Assay Data	Appendix

SUMMARY

The Simba Property was visited for three days during a property regional silt survey. A total of 143 silt were gathered by Scott Fleming, Mike Lindley, Issac Fage, and Tyson Foxcroft. A geological mapping program was undertaken to map the regional rock units found on the Shell Property. Chris Ash of Ash Consulting was hired with an assistant Lisa Peters to map the property. The mapping took place on June 29-30 and August 22 – 29, 2004. The silt survey outlined gold and copper anomalies found on the south slope of Simba Mountain with no geological explanation. The mapping outlined a new style of mineralization of quartz veins with visible gold found in saddle reef style setting. The Saddle reefs and the anomalous silt area will be targeted in the 2005 exploration program.

1.0 INTRODUCTION

This report describes a silt survey conducted on the Simba Property in the Dawson Mining District, Yukon Territory. The Simba Property host a large Algoma style Iron Formation with quartz veins that contain copper and visible gold. A total of 143 silts were taken around the property with the intention to define new targets on the property. The program was successful in outlining gold anomalies associated with copper on the south slopes of Simba Mountain.

2.0 LOCATION AND ACCESS

The Simba Property is centered at 64°03 N, 140°25 W in the central Yukon Territory, close to Alaska border. The Property is 75 kilometer north west of Dawson City. The only access is via helicopter from Dawson City. There also the Clinton Creek Road, a summer road that gets to within 9 kilometer of the property. It's a good spot to sling in supplies.

3.0 PROPERTY DESCRIPTION

The Simba Property consists of 488 full claims staked under the Yukon Quartz Mining Act in the Dawson Mining District. This report is for Simba 41-214 claims. The claims are located on NTS # 116 C / 09.

4.0 PHYSIOGRAPHY

The Simba Property is located in the Klondike Ecoregion. It straddles the southern Olgilvie mountain range to the north and the Tintina Valley to the south. Elevation on the Property range from 2500 ft to 4500 ft. The property vegetation is covered with spruce and aspen on the southern slopes, black spruce on the northern slope up to an elevation of 3700-4000 feet. The Property is covered with alpine tundra shrubs above 4000 feet.

5.0 REGIONAL AND PROPERTY GEOLOGY

5.1 REGIONAL GEOLOGY SETTING

The Simba BIF and related gold-quartz veins are contained in clastic sedimentary rocks of the Precambrian to Lower Cambrian Hyland Group, which is a component of the Selwyn Basin off-shelf succession. The Hyland Group currently comprises the lowest the lowest stratigraphic element of the Dawson Thrust Sheet, just north of the regionally extensive Tintina Fault Zone (Thompson et al. 1992).

5.2 PROPERTY GEOLOGY by Chris Ash, MSc, PGeo

The property geology includes three primary stratified units. A basal limestone sequence is overlain by an interval of interlayered sandstone, shale and gritty sandstone, which in turn overlain by a succession of intermediate volcanoclastic and epiclastic rocks. The middle clastic sedimentary unit hosts both the BIF and spatially associated gold-quartz veins. The entire section is folded into a shallow (10 to 30°) north-northeast plunging, anticlinal fold structure. Several small plugs of diorite in addition to minor felsite dikes, of unknown age are also present, but volumetrically minor.

The contorted surface trace of the BIF at the nose of the major fold structure suggests the development of related minor folding. Mapping along strike to the immediate south-southeast beyond the property boundary, in the same stratigraphic succession defines north-northeast trending folds with amplitudes of several 100 meters (Thomson, 1992), and supports this interpretation. If valid, this relationship would suggest that the folded nose of the iron formation would be a more prospective area for quartz vein formation due to the higher potential for development of local, dilatational settings within the hinge zone. This relationship remains to be fully established but may indirectly speak to the historic production of placer gold on Shell Creek, which has eroded through and drains the nose of the major fold structure.

Mineralization

Two distinct, yet apparently genetically related styles of epigenetic mineralization are noted. Low sulphide gold-quartz veins with visible gold represent the most obvious target, however, fine-grained clastic sedimentary rocks marginal to these veins are locally pervasively chloritized and Cu-enriched and may be of economic significance.

Au-Quartz Veins

Bull white, gold-bearing quartz veins, typically highly fractured, outcrop intermittently along the 6 to 7 km extent of the northern southwest-dipping limb and along the nose of the folded BIF. Large, 2-4 meter boulders of quartz are scattered 100 to 200 meters down-slope from the identified zone of gold-quartz vein outcroppings. Gold-quartz veins do not appear to be hosted by the BIF but are contained in fine- to medium-grained clastic sedimentary rocks immediately above and below the banded iron interval.

Two distinct stages of quartz are recognized within some of the veins. An earlier stage consists of pervasively fractured, bull-white quartz. A relatively younger and more often mineralized stage consists of quartz-carbonate-chlorite with trace to two percent Cu sulphides (\pm gold). This late stage quartz is typified by vuggy 6 to 10 cm wide veins with well-developed comb structures that usually form at the margins of the early stage quartz, or infill cavities with it.

Late stage veins contain from trace to two percent Cu sulphides occurring in 0.3 to 0.7 cm size clots, readily identified due to green/blue, malachite/azurite staining. A single thin section analysis of a mineralized quartz vein identified predominantly chalcocite (75%) with lesser bornite (25%), and trace amounts of pyrite, chalcopyrite and lead-bismuth tellurides (Wong, 2004).

The general distribution, continuity and grade of the gold-quartz vein system remains to be established.

Cu in Vein-marginal Chloritized Sediments

In some areas, fine-grained, foliated clastic sedimentary rocks marginal to the gold-quartz veins are pervasively chloritized and contain elevated Cu. Malachite staining is common along cleavage surfaces within these chloritized sediments. Three grab samples of the malachite-stained, chloritized sediments are reported to have returned assays indicating from 1.6 to 1.8% Cu. Chloritized, Cu-rich Vein-marginal chloritization with Cu enrichment appears to be associated with veins with late stage quartz.

Locally, the chloritized, Cu-stained sediments were noted to extend at least 1 to 2 meters away from the quartz vein margins, however, the limits could not be accurately established due to overburden. The extent and continuity of Cu-grade within these alteration haloes remains to be determined.

